

GH 3197

A GEOHYDROLOGICAL INVESTIGATION IN THE VICINITY OF
SCHWEIZER-RENEKE, WESTERN TRANSVAAL

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N.B. These maps were originally drawn to the scale of 1:50 000 (A0 size) and were reduced to a scale of approximately 1:150 000 (A3 size) for convenience. Copies of the original maps can be obtained from the Information Section on request.

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1. INTRODUCTION

1.1 Objective

The primary objective of this investigation was to ascertain whether alternative sources of ground water of a potable quality could be obtained for use by the municipality of Schweizer-Reneke within a reasonable distance of the town.

Such sources could be used either to partially supplement the existing sources of municipal water in Wentzel Dam or to provide the bulk of the municipal water during periods of severe surface water shortage.

Areas of favourable ground water occurrence were to be delineated but at this stage no production boreholes were to be developed for immediate use by the municipality.

1.2 Location of study area

The main study area is confined within a 15 km radius of Schweizer-Reneke, an area of about 700 km^2 . The total area investigated is located within the lines of latitude $27^{\circ}02'$ and $27^{\circ}24'$ south, and within lines of longitude $25^{\circ}10'$ and $25^{\circ}32'$ east. Map coverage is provided by the 1:50 000 topo-cadastral sheets 2725 AA (Amalia); 2725 AB (Schweizer-Reneke); 2725 AC (Diewedraai); 2725 AD (Koosfontein); 2725 BA (Glaudina) and 2725 BC (Sewefontein). Farms surveyed are listed alphabetically in Appendix 1.

1.3 Previous work

Prior to this investigation the geology of the Schweizer-Reneke district was well known but little co-ordinated information was available on geohydrological conditions.

The Schweizer-Reneke area was geologically mapped by Van Eeden in 1934-35. The result of the mapping were published by the Department of Mines in 1962 on one of the 1:125 000 series of geological maps namely 2725A. An explanatory booklet (Van Eeden, 1963), with a short section on ground water, accompanied the published map.

The basic geohydrological information was provided by De Villiers (1961) from his experience of siting boreholes in the south-west Transvaal between 1949 and 1956. De Villiers proved that the electrical resistivity method contributed considerable improvement to the traditional techniques, namely dowsing, in borehole site selection on granite and quartzite but with less significant results being obtained from lavas and Karoo mudstones. The few chemical analyses that were published indicated high nitrate concentrations from sampled ground waters both in the granite (30 mg/l) and in the lava (80 mg/l) in the Schweizer-Reneke district.

Subsequent borehole siting assignments by the Geological Survey in the Schweizer-Reneke district were undertaken by Gordon-Welch (late 1950's) and Erasmus (mid-1960's). Few of the 30 borehole sites selected were subsequently drilled.

Water samples collected from municipal boreholes in Schweizer-Reneke in April 1980 were analysed and reported on by the Division of Chemical and Biological Services (1980). Organic pollution was inferred as the ground water was found to have a high concentration of nitrite and nitrate, plus a high faecal coliform bacteria count. Information on municipal boreholes drilled in 1979 was supplied by Cahier (1980).

Steyn (1980) conducted a short investigation into the ground water potential of part of the Schweizer-Reneke municipal area. Five boreholes sites were selected and reported on. The borehole Ss 11 (G32304) was at that stage incorporated into the municipal borehole field. (Borehole Ss 20 (G32303) should also now be included.)

Information on boreholes previously drilled by the Boring Services subdivision of the Department of Water Affairs was provided on DW56 completion forms.

Details on inflows, releases and evaporation losses from Wentzel Dam were obtained from the Division of Hydrology and details on various yield analyses were extracted from Planning Division File No. P330.

Complete aeromagnetic coverage was available from east-west flights flown in 1977 and 1978.

1.4 Background of investigation

Wentzel Dam with full supply capacity $5,29 \times 10^6 \text{ m}^3$ (1979) is usually capable of providing adequate supplies of surface water for municipal consumption in Schweizer-Reneke ($0,53 \times 10^6 \text{ m}^3$ in 1979) in addition to supplying downstream irrigators ($0,6 \times 10^6 \text{ m}^3$ in 1979). However as a result of reduced inflow into the dam due to subnormal rainfall over the dam catchment during the preceding two years the dam storage at the end of 1979 was reduced to 20% of the full supply capacity. In anticipation of the situation worsening the municipality commissioned the development of a borehole field in the south-western portion of the town for the abstraction of ground water. By May 1980 a number of boreholes had been equipped and connected to the town water distribution system. Abstractions from Wentzel Dam were discontinued in May 1980.

To meet the town's municipal water requirements, namely $0,53 \times 10^6 \text{ m}^3$ per annum, it was planned to abstract ground water from this borehole field at an average rate of 22,6 l/s based on a pumping schedule of 18 hours per day. Declining borehole yields resulted in the visit of Steyn and Venter of this Division in August of that year. They noted that the continuous abstraction of ground water in this area was resulting in seriously declining ground water levels and established an observation borehole to monitor this decline. An additional abstraction borehole was established adjacent to the borehole field in an attempt to restore the required pumping

rate. By November 1980 the pumping rate had declined to about 14 l/s when pumped for 20 hours per day and a water supply crisis was developing in Schweizer-Reneke. Fortunately heavy rains at the end of November resulted in the recovery of the dam level and surface water again became available from Wentzel Dam.

Since December 1980 there has been no serious municipal water supply problem in Schweizer-Reneke. However in view of the town's growing municipal water requirements and the possibility of further 'dry' periods occurring in the future this current geohydrological investigation was commissioned to identify other favourable water-bearing formations in reasonably close proximity to Schweizer-Reneke which could be capable of supplying all, or part, of the municipal water requirements.

It is envisaged that surface water and ground water be utilized conjunctively as conditions dictate. In other words during periods of above average rainfall and runoff of surface water that only surface water be utilized from Wentzel Dam, and that aquifers be allowed to recover. During periods of subnormal rainfall and surface runoff that ground water be used to meet the municipal water requirements. To this end a variable yield analysis on Wentzel Dam was undertaken by Langhout (Planning Report No. 330) utilizing 47 years of dam inflow and rainfall records. Three different scenarios were analysed. During the worst scenario a shortfall of surface water over a continuous period of 12 months would have to be supplied from ground water sources at a continuous pumping rate of 28,5 l/s. This would be sufficient to meet the municipal water requirements up to about the year 1990.

1.5 Approach of investigation

Initially the investigation was confined to land owned by the municipality. This was later expanded to an area confined within a 15 km radius of Schweizer-Reneke and detailed investigations were concentrated in more favourable areas. Accessibility and proximity to existing infrastructure was a consideration.

The initial exploration phase was aimed at producing a base map from a stereoscopic examination of areal photographs and from an interpretation of aeromagnetic, geological and topographic maps. The complied map effectively enable target areas to be delineated which could justify receiving more attention later. These areas were later subjected to detailed geological mapping and geophysical exploration, and exploration drilling was undertaken at some favourable sites.

To gain an insight of the regional ground water movement and quality, a ground water contour map was complied from ground water level measurements made during a restricted borehole survey and a ground water quality map was compiled from the results of chemical analyses undertaken on ground waters sampled in the area.

To estimate the possible potential of various water-bearing formations a number of aquifer tests were also undertaken.

The approach adopted was not to develop production boreholes at this stage for immediate use by the municipality but rather to delineate relatively limited areas which provided favourable conditions for the successful exploitation of ground water. Based on this report, and on other factors, the municipality would be in a better position to decide which areas could most economically be developed for the recovery of ground water for augmentation of the present municipal water supply.

2. MUNICIPAL WATER SUPPLY

2.1 General

Water for municipal consumption in Schweizer-Reneke has since the 1950's been supplied from Wentzel Dam apart for short periods. It is expected that for the foreseeable future the dam will continue to provide the bulk of the municipal water requirements but not the entire requirements. This implies that supplementary supplies of municipal water will have to be supplied during periods of subnormal rainfall over the catchment when the dam is at a low level, ideally from ground water sources.

In 1980 the municipality had to provide domestic water for a total population of about 10 700 (Steyn, 1980). A population growth rate of about 3,5% is predicted for Schweizer-Reneke. The annual increase in rate of municipal water consumption is predicted to be about 5,2%. The annual predicted municipal water consumptions are shown in Table 1 below up to the year 2015.

TABLE 1: Water requirements to the year 2015

Calender year	Water demand (10^6 m^3)	C.P.R. 1/s	Calender year	Water demand	C.P.R. 1/s
1980	0,55*	17,3	2000	1,52	48,2
85	0,71	22,5	05	1,95	61,8
90	0,91	28,8	10	2,51	79,5
95	1,18	37,4	15	3,24	102,7

(*not met)

(C.P.R. = continuous pumping rate)

As yet there is very little industrial development and the largest water consumers are the two schools and the hospital. During 1979 the latter utilized 2,7% of the total municipal supply.

The marked seasonal demand for municipal water is clearly evident in Table 2 below.

TABLE 2: Monthly consumption pattern as a percentage of the annual water consumption

January	13	July	4
February	11	August	5
March	8	September	10
April	5	October	12
May	4	November	12
June	3	December	13

If the trends and seasonal demands continue it can be expected that by the year 2000 the peak demand for municipal water would require a continuous pumping rate of almost 75 l/s.

2.2

Availability and reliability of records

Available records on the quantities of surface water entering and being released from Wentzel Dam are incomplete. Useable records only exist for the period 1964 to 1981. Even for this period the record is incomplete. Table 3 and Fig. 15 show therefore only the minimum quantities of water released from the dam.

2.3

Wentzel Dam

Wentzel Dam is an earth wall dam constructed across the Harts River at a natural damming site between two low quartzite ridges. The dam has an upstream catchment of $9\ 220\ km^2$ over which a mean annual precipitation of 565 mm has been measured. The mean annual runoff from the catchment is estimated to be $37,0 \times 10^6\ m^3$. The present full supply capacity of the dam is estimated at $5,29 \times 10^6\ m^3$. At FSC (gauge plate reading 6,1 m) the area flooded behind the wall is 2 760 hectares. At FSC the maximum depth of water in the dam is 8,1 m and the average depth 1,9 m. Evaporation losses are very high estimated at $3,4 \times 10^6\ m^3$ annually for the period 1963/64 to 1979/80.

FIG. 15
G.H.P. 5656

BREAKDOWN OF WATER SUPPLIED BY
WENTZEL DAM OVER THE PERIOD
1964 to 1981

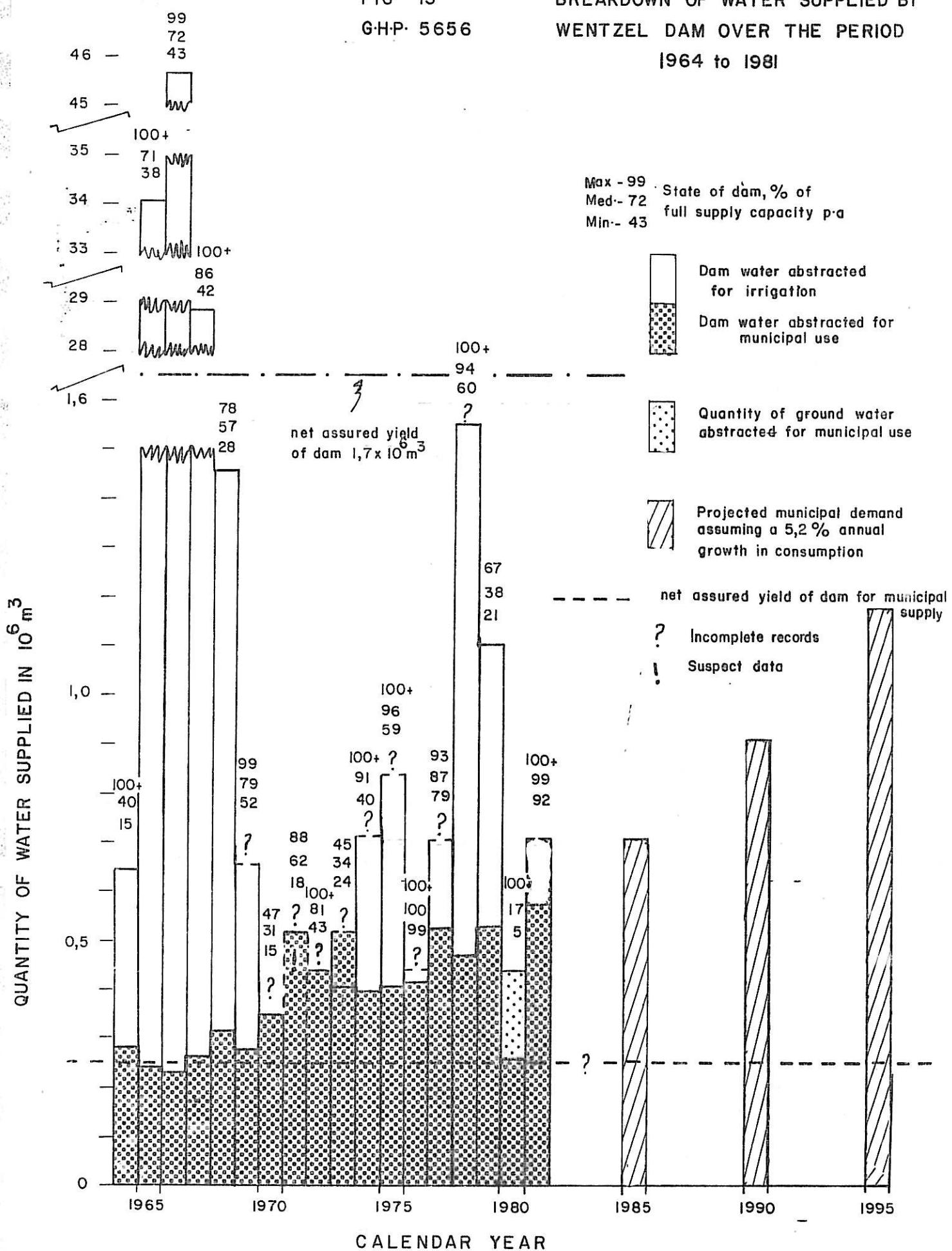


TABLE 3: Water releases from Wentzel Dam

Calender year	Quantity of water released 10^6 m^3	Number of months	Quantity of water released 10^6 m^3	Schweizer-Reneke metered for irrigation Municipality consumption 10^6 m^3
1964	0,366	3	0,280	no record
1965	3,165	12	0,240	no record
1966	4,337	11	0,230	no record
1967	2,621	10	0,261	no record
1968	1,144	9	0,311	no record
1969	(0,379)	7	0,276	no record
1970	no record	no record	0,348	0,348
1971	no record	no record	0,517!	0,343
1972	no record	no record	0,436	0,391
1973	(0,109)	(2)	0,404	0,404
1974	(0,318)	(9)	0,394	0,394
1975	(0,436)	(6)	0,401	0,405
1976	(0,028)	(1)	0,411	0,412
1977	(0,178)	(2)	(0,421)	0,523
1978	(1,088)	(8)	0,465	0,497
1979	(0,573)	(6)	0,528	0,529
1980	NIL	NIL	0,254	0,431*
1981	0,139	3	0,511**	no record

() incomplete records

! suspect record

* difference made up by ground water

** not including December 1981

A recent sounding survey undertaken by this Department indicated that the original volume of the dam has been significantly reduced as a result of sedimentation within the dam.

The dam was built at the request of the Transvaal Irrigation Board and apparently became fully operational on 1934. The original purpose of the dam was to supply via two canals (see Enclosure 1) water to the irrigation scheme situated downstream adjacent to the Hart River. In the mid-1950's surface water from the dam was also being used by the Schweizer-Reneke municipality to meet local demand. To safeguard the interest of the municipal water consumers a minimum low level was imposed below which no water was to be released for irrigation. The level was recently re-established at a gauge plate reading of 4 m corresponding to a dam volume of $1,39 \times 10^6 \text{ m}^3$ or 26% of the FSC. This minimum level still applies.

The quantities of water released annually from Wentzel Dam are given in Table 3 from 1964 and diagrammatically in Fig. 15.

2.4 Municipal water supply

Available records indicate that Wentzel Dam adequately supplied municipal water from 1964 to May 1980. During this period varying quantities of surface water were released for irrigation (Table 3 and Fig. 15).

For the first time on record abstraction from the dam was discontinued in May 1980 due to the poor quality of the water remaining in the dam. From May 5 to November 1980 Schweizer-Reneke was totally dependent on ground water (see section 1.4). During this period of seven months a municipal water shortage developed as only $0,177 \times 10^6 \text{ m}^3$ of ground water could be recovered from the shallow weathered granite near town. The borehole field consisting of six and later seven boreholes was located within an area of 15 hectares (for locations see Enclosure 1 and Steyn, 1980). The limited borehole records available are presented in Table 4.

TABLE 4: Borehole data, Schweizer-Reneke Municipality

Borehole number	Date drilled	Depth (m)	Casing length (m)	First water strike	Tested yield (l/s)*	Initial pumping rate (l/s)	Pumping period hours/day
			Ø 165 mm	(m)			
Ss2	1979	22,3	12	12,8	17,5	7,5	18
Ss3	1979	22,8	12	13,0	16,1	5,0	18
Ss5	1921	36,6	?	?	?		pumped occasionally
Ss6	1979	36,6	?	?	11,4	3,8	18
Ss7	1979	39,6	?	?	3,8	2,6	18
Ss8	1979	39,6	?	?	10,1	3,7	18
Ss11	1980 (G32304)	40,0	?	?	2,8	2?	pumped occasionally

(*Tested yields are for pumping period of 8 hours in December 1979.)

Ss 11 was drilled and tested for 9 hours in July 1979 and integrated into the borehole field.

The aquifer failed to supply the anticipated quantity of municipal water for the following reasons:

- (i) no recharge to the aquifer took place from the nearby Harts River, and
- (ii) the storage for ground water within the weathered granite was insufficient to maintain a constant pumping rate of 17 l/s.

With the dam again filling and overflowing in December 1980 water could again be abstracted for municipal consumption. The borehole field is not at present in use but is maintained in a state of readiness. Ground water level measurements indicate that by April 1981 ground water levels had again recovered to normal.

2.4.1 Total supply available

The Division of Planning has calculated that the net assured yield of Wentzel Dam is $1,7 \times 10^6 \text{ m}^3$ annually, equivalent to a continuous pumping rate of 54 l/s. However if downstream irrigators utilized their full and rightful allocation only $0,25 \times 10^6 \text{ m}^3$ per annum would be available for municipal consumption by Schweizer-Reneke, equivalent to a continuous pumping rate of about 8 l/s. Reference to Table 3 indicates that the Schweizer-Reneke municipality has abstracted considerably more than $0,25 \times 10^6 \text{ m}^3$ per annum for many years and that this state of affairs is expected to continue in the future particularly as more irrigators relinquish their quotas of irrigation water.

Extended pumping of ground water during 1980 demonstrated that only about $0,2 \times 10^6 \text{ m}^3$ is available to the municipality from the existing borehole field. If no competition was experienced from private boreholes to the south of the borehole field an additional $0,1 \times 10^6 \text{ m}^3$ of ground water may become available for municipal use.

If use is made of both Wentzel Dam and the borehole field the minimum quantity of water available annually to the municipality would be

$$(0,25 + 0,20) \times 10^6 \text{ m}^3$$

$$= 0,45 \times 10^6 \text{ m}^3$$

However it might be expected that for most years at least $0,5 \times 10^6 \text{ m}^3$ could be supplied from the dam plus the $0,2 \times 10^6 \text{ m}^3$ from ground water giving a total of available water of $0,7 \times 10^6 \text{ m}^3$, sufficient to about the year 1985.

In the case of total expropriation of the rights of downstream irrigators a potential supply of $1,9 \times 10^6 \text{ m}^3$ ³ annually could be made available to the municipality sufficient to about the year 2005 assuming projections of water consumptions to be correct.

It is apparent that in the event of subnormal runoff into the dam again occurring, repeating the crisis condition of 1980, that the ground water available for exploitation at the municipal borehole field would be totally inadequate.

Further it is apparent that the estimates made previously of the quantities of ground water available for use at the existing borehole field were grossly in error.

2.4.2 Alternative solutions

As a potential water supply problem exists in Schweizer-Reneke several possible solutions can be proposed as to how the municipal water supply can be augmented. These are listed below, not in order of importance;

- (i) increase the net assured yield of Wentzel Dam by raising the wall,
- (ii) reduce the quantity of water released from Wentzel Dam to downstream irrigators, ultimately by expropriating all rights of irrigators,
- (iii) improve the runoff characteristics of the Harts River catchment above Wentzel Dam,
- (iv) pipe water from Bloemhof Dam,
- (v) incorporate Schweizer-Reneke in the Orange Free State Goldfields water scheme,
- (vi) use treated sewage effluent,

(vii) the development and utilization of available ground water resources in the area.

It is to this latter solution that this report is addressed, being the most feasible and immediate solution to the potential municipal water shortage in Schweizer-Reneke.

3. PHYSIOGRAPHY

3.1 Topography and drainage

The area investigated is part of the Highveld region with elevations varying between 1 270 and 1 400 m above M.S.L. The major drainage artery is the Harts River which conveys the only perennial surface water flow in the area. The general configuration of the topography might have been initiated during the Karoo glacial period.

Van Eeden (1963, p.5) has described the landscape as part of:

"..... a great, undulating plain diversified with low rounded hills or rises and hollows which differ little from one another and from the general plain in height."

Despite this low natural relief two major drainage basins exist in the study area, namely C330 (part of the Harts River system) and C325 (part of the Vaal River system), which are separated by a sometimes pronounced surface water divide. This ridge, trending approximately north-east to south-west (see Enclosure 2), forms the topographic high of the region with a maximum elevation of 1 397 m on Panfontein 58 HO. South and south-east of this divide the ground surface slopes very gently southwards with a gradient of 8×10^{-4} on De Park 87 HO.

In the north-west of the area a very poorly defined drainage divide with an elevation of approximately 1 360 m separates the tributaries of the Dry Harts River from those of the Harts River to the south.

The Harts River draining south-westwards is characterised by a uniformly gentle gradient of approximately 1×10^{-3} . The actual fall in elevation is from 1 313 m in the north-east to 1 270 m in the south-west over a distance of some 41 km.

The valley of the Harts River within the lava areas tends to be relatively narrow; slopes towards the river are of the order of 3×10^{-3} . A gap some 300 m wide has been cut through the ridge formed by quartzites of the Bothaville Formation, north-east of Schweizer-Reneke where the Harts River manifests the phenomenon of superimposed drainage. Downstream of this gap the flood plain of the river basin becomes considerably wider, up to 1 000 m in places, as the Harts River meanders across the plain underlain by granite.

Throughout this area there is a clear relationship between geology, topography and drainage. East of Schweizer-Reneke Karoo mudstones overlie lava and form a high flat plain which has a very poorly defined drainage network and is thus subject to occasional water logging. This area is characterised by the development of numerous shallow pans or depressions some of which are incised up to 10 m below the local topography. Immediately east of Schweizer-Reneke, where the Karoo formations are absent, lava forms a better defined ridge but again surface drainage is poor.

The quartzite produces a small escarpment with the scarp slope often rising some 20 to 30 m above the surrounding plain. This ridge gives rise to several dip slope streams e.g. Hamburg Stream and occasional strike streams which follow the contact between the sedimentary rock and the overlying lava.

It is noticeable that the granite forms more rolling landscape which is well drained by numerous tributaries of the Harts River. (See minor catchments on Enclosure 2.)

Small hills rising some 20 m above the local landscape e.g. on Palachooma 64 HO, are caused by outliers of quartzite resting on granite.

3.2 Climate and vegetation

Schweizer-Reneke falls within the northern steppe climatic zone of Southern Africa (Sn.). This zone is characterised by very cold winters and excessively hot summers with precipitation occurring principally as the result of thunderstorms in summer and autumn. Rainface station No. 397/581 has been operating in Schweizer-Reneke since 1953. From these records, for the hydrological years 1953/54 to 1980/81, the mean annual rainfall has been calculated as 509,5 mm, with a standard deviation of 136,6 mm. Total precipitation plotted as both bar graphs and as a curve of cumulative departure from the rainfall mean in Fig. 16 confirms the variability of the rainfall. Subnormal rainfall in two consecutive years 1978/9 and 1979/80 was responsible for the water shortage experienced in the town during 1980.

The natural vegetation cover of Kalahari Thornveld invaded by Karoo has been extensively modified by man. However, natural vegetation types persist on quartzite, lava and some gravel outcrops e.g. on parts of Mimosa 61 HO, Rietput 60 HO and Schweizer-Reneke Townlands 62 HO. Elsewhere the original vegetation has been all but cleared and replaced by grasslands capable of supporting a higher livestock density or by arable lands. The main crop is maize with some groundnuts being grown on the more friable soils. Where water for irrigation is available lucerne, wheat and market vegetables are grown.

4. GEOLOGY

4.1 General

The distribution of the geological formations outcropping in the vicinity of Schweizer-Reneke is depicted in Enclosures 1 and 2 while the relationship between formations is shown by the cross-section in Fig. 17. Lithostratigraphic terminology is according to S.A.C.S. 1980.

FIG. 16

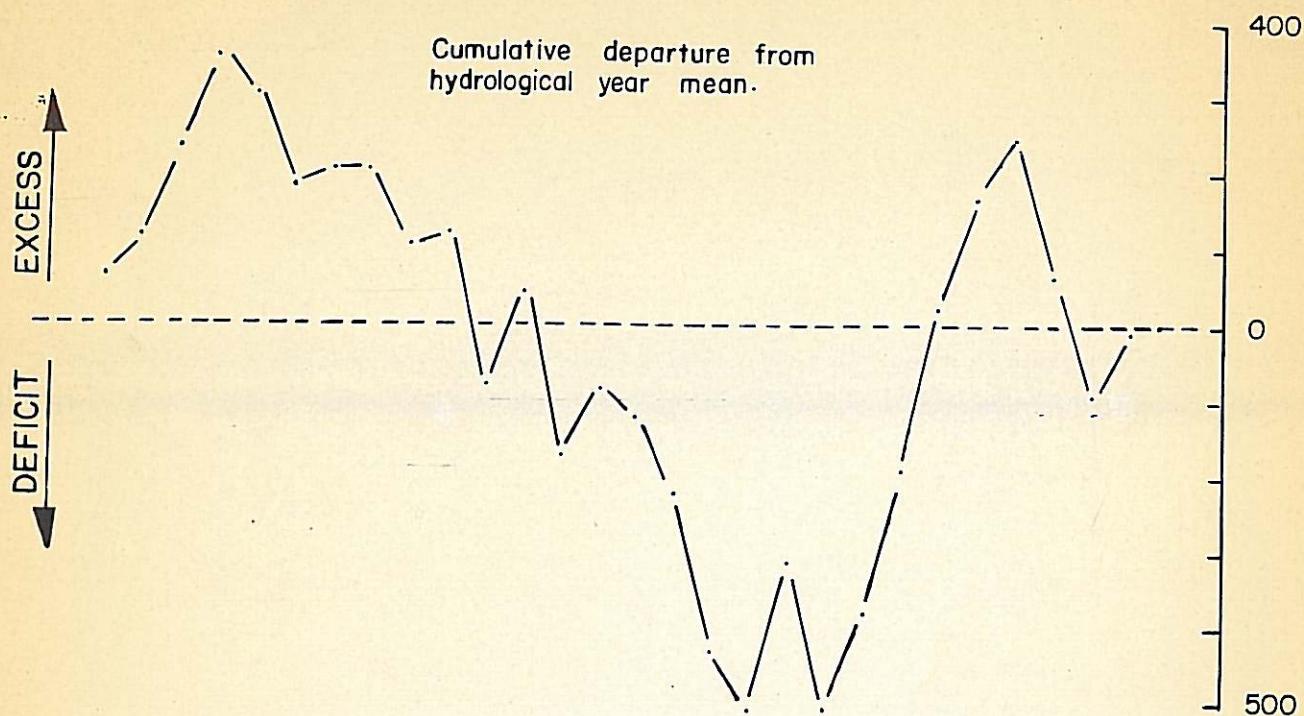
CHARACTERISTICS OF RAINFALL STATION

G.H.P 5657

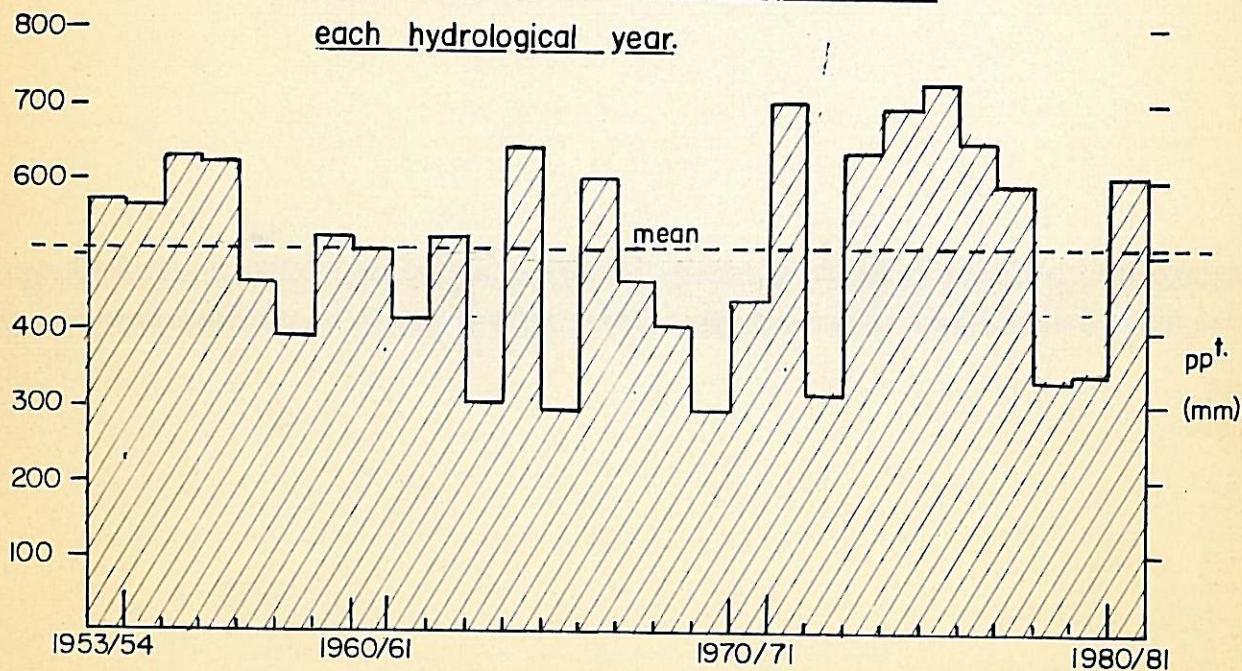
No. 397/581, SCHWEIZER-RENEKE POLICE

STATION FOR THE HYDROLOGICAL YEARS

1953/54 TO 1980/81.



Histogram showing total precipitation in each hydrological year.



N.B. Hydrological year is from October to September.

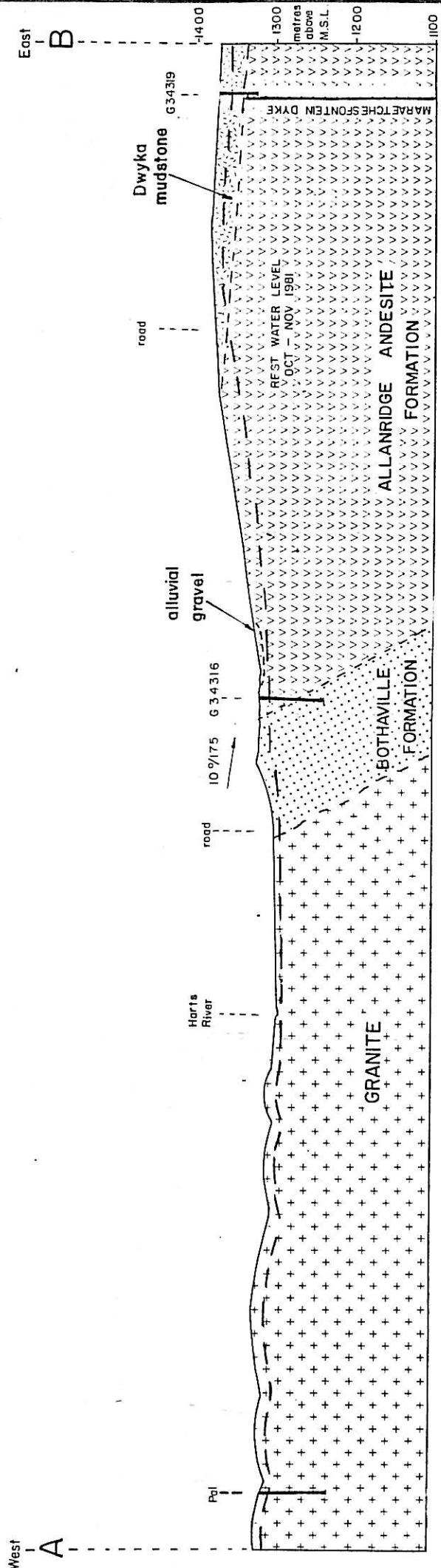
station altitude: 1280m

mean precipitation = 509,5mm

standard deviation 136,6mm

FIG. 17
G.H.P. 5658

GEOHYDROLOGICAL PROFILE OF SECTION LINE A-B



The Archaean basement complex comprises rocks of the Swazian era that have been intruded and metamorphosed by younger siliceous magma. In the area studied older sedimentary sequences were deposited on the uneven, eroded floor of this basement and are now arranged quaquaversal to the central dome. The regional dip is fairly shallow being only 6 to 12° while local dips of up to 20 or 30° are caused by small scale antiforms and synforms. Extruded apparently conformable onto the sedimentary sequence were numerous, thick, mainly andesitic lava flows. During and after this period of volcanic activity tectonic movement commenced. This was responsible for the development of faulting and fracturing along north-south and north-east to south-west trending lines. Pre-Karoo diabase dykes were emplaced along some of the zones of structural weakness.

After subsequent intense glaciation in Karoo times sedimentary rocks, mainly argillaceous, were deposited unconformably across the exposed irregular rock floor of lava. These rocks are horizontally bedded indicating only minor earth movement since, however some fracturing and faulting with accompanying dyke intrusion, again with a dominant north-south trend, has occurred.

A long period of erosion then followed which has resulted in the present rock distribution, the accumulation of superficial deposits and the formation of soil.

4.2 Kraaipan Group

The main outcrop of this group occurs some 3 to 4 km west of the studied area on the farms Abelskop 75 H0 and Goudplaats 96 H0. Dynamically and thermally metamorphosed sediments and lava constitute this group. These rocks are highly folded, faulted and fractured. The principal rock types developed are various kinds of schists (mainly chlorite-schists), quartzite (partly magnetic), chert, jaspelite, banded ironstone, amphibolite and lava.

Within the studied area small outcrops (xenoliths) of amphibolite and quartzite occur within the granite, e.g. on Mooilaagte 91 HO. Previous workers have assigned these to the Kraaipan Group. Similar rock types crop-out on Hamburg 82 HO and were also encountered in drilling on both Hamburg and Mimosa 61 HO.

The presence of sub-cropping Kraaipan Group rocks might be indicated by north-south trending intense magnetic gradients both north of Schweizer-Reneke and west of longitude $25^{\circ} 15' E$ as indicated on the aeromagnetic map.

4.3 Archaean granite gneiss

These rocks occur west and south-west of Schweizer-Reneke. Borehole records and the presence of small inliers of granite indicate that granite subcrops below a thin cover of Ventersdorp Supergroup north of Schweizer-Reneke.

The granite generally exhibits only weak foliation, although occasionally strong foliation does occur. The granite gneiss varies in colour from grey-green to pink depending on the proportion of pink, orthoclase feldspar phenocrysts present. The rock is medium-to very coarse-grained with occasional aplitic and pegmatic phases. The major minerals invariably present are orthoclase feldspars, quartz and plagioclase feldspars. Muscovite, biotite, hornblende and pyroxenes also occur though their importance as rock-forming minerals is very variable. Occasionally they constitute up to 10% of the rock while at other times they are all but absent.

Feldspar have generally been altered to sericite, calcite, or chlorite. Within the zone of intense weathering generally less than 15 m, the ferromagnesian minerals and to a lesser extent the feldspars have been decomposed to form various clay minerals.

Within the granite gneiss are thin layers (xenoliths) of micaceous schists, pyroxenite-amphibolite and recrystalline quartz (e.g. on Hamburg 82 HO).

4.4 Ventersdorp Supergroup

Due to progressive overlap on the eroded, uneven Archaean basement only the uppermost formations of the Ventersdorp Supergroup outcrop. These are the lithostratigraphic units of the Bothaville and Allanridge Andesite Formations. Deep drilling indicates that younger lavas of the Platberg Group underlie the Bothaville Formation eastwards in the main basin e.g. borehole Sb 1 on Springbok 191 HO (Van Eeden 1963). North of Schweizer-Reneke, only a thin cover of Bothaville and Allanridge Andesite Formation is present covering a palaeo-topographic high.

4.4.1 Bothaville Formation

This formation forms a conspicuous topographic feature as it is composed of relatively resistant, dipping, sedimentary rocks. Although the Bothaville Formation is unconformable on the granite, the unconformity has not been observed in the field. The topmost contact with the overlying lava is sharp, but in certain localities e.g. G34315 (Fig. 10) intercalations of lava and quartzite occur indicating more or less contemporaneous lava effusions and sedimentation phases. (Van Eeden 1963.)

The thickness of this formation is probably in the order of 100 m. Borehole logs presented by Van Eeden (1963) show that a maximum drilled thickness of 124 m, dip about 10°, was encountered on Springbok 191 HO. During the present investigation the formation was never fully penetrated and the maximum thickness drilled was only 49 m, dip 6°, in G34314.

Van Eeden (1963) observed that the basal succession is characterised by conglomerates and grits.

Examination of rock exposures and drill cuttings demonstrate that the formation is composed principally of fine- to medium-grained arenites. The colour of these rocks is extremely variable ranging

from milky-white through grey-green to lilac and purple-red. The degree of cementation is also extremely varied, in some cases the rock is an extremely hard quartzite while in other cases cementing material has disappeared and the rock is a friable sandstone. These changes in the degree of cementation occur rapidly vertically e.g. see Fig. 9 and less sharply horizontally, e.g. contrast Fig. 11 and Fig. 12.

The contrast in the relative hardness of these rocks may be partly explained by bedding joints and the effect of weathering. Quartzite units are generally well bedded, unit thickness generally 0,6 m, and demonstrate cross bedding with occasional ripple marks. The more friable zones are generally thinly bedded and occur both as thin units some 50 mm thick between quartzite bands, and as thicker units up to 2 m thick. Well cemented, fine-grained sandstones and siltstones also occur and have been locally quarried for floor and roofing tiles.

4.4.2 Allanridge Andesite Formation

This formation is principally composed of many separate flows of andesitic lava. The thickness of this succession increases rapidly down dip of the outcrop of the underlying sediments. On Springbok 191 HO this lava formation is some 480 m thick (Van Eeden 1963).

The lava is generally andesitic but occasionally rhyolitic. Thin bands of tuff and agglomerate are preserved between some lava flows.

Fresh lava ranges in colour from light grey-green through green to dark grey-blue and from very fine-grained to coarse-grained. Large amygdales of chalcedony and jasper, sometimes epidotised, are locally common. Black and greenish phenocrysts, up to 2 mm also occur.

Joints and fractures are often filled with silica, calcite and epidote. Weathering is extensive and fairly deep even where lava

apparently crops out e.g. G34316 which was sited on outcrop. The weathering process produces abundant clay. Where more than 15 m of Karoo rock is present, the underlying lava is generally fresh and unweathered.

4.5 Karoo Sequence

Rocks of Karoo age (assigned to the Dwyka Formation and Ecca Group by previous workers) unconformably overlie much of the Allanridge Andesite Formation. The outcrop pattern preserved is the remnant of a much more extensive cover. Due to the easily erodable nature of these argillaceous rocks no exposure of Karoo rocks have been observed in the studied area. Information on the distribution of the Karoo must therefore be inferred from borehole records, aeromagnetic maps and aerial photographs. Thus the map presented in Enclosure 1 and 2 is subject to revision and must be treated with circumspect.

Drilling records indicate that on Geluk 56 HO the thickness of the Karoo rocks is some 50 m so that the lava/Karoo contact is at approximately 1 310 m above M.S.L. The thickness of Karoo encountered in G34319 was 33 m (contact at 1 350 m) and in G34318 was 6 m (contact at 1 330 m). These elevations indicate that the eroded surface of the lava is highly irregular and thus the thickness of the Karoo deposits should also be highly variable.

The formation encountered during supervised drilling was a monotonous sequence of greyish occasionally purplish mudstone which in part contain calcareous concretions and are finely laminated. Interbedded with the argillaceous rocks are rare thin sandstone layers. No diamactite was observed.

4.6 Superficial deposits

Much of the solid geology of the area is masked by varying thicknesses of superficial deposits which include soil, sand alluvium, river-terrace gravels and calcrete.

4.6.1 River-terrace gravels

Proved deposits have been shown separately in Enclosures 1 and 2. These deposits are found along the Harts River valley and old drainage courses. As these gravels have been extensively worked for their diamond content good exposures occur. Diggings show that the gravel usually directly overlies bedrock and that the gravel thickness is generally less than one metre but occasionally thicker, e.g. at Schweizer-Reneke. The gravel consists of mainly poorly size-sorted pebbles within a clay and silt matrix. Principal pebble types are fairly well rounded chert, chalcedony and jasper, derived from the lavas of the Ventersdorp Supergroup. Numerous pebbles and boulders of eluvial lava also occur. Occasional pebbles of granite and banded ironstone, when found within the lava subcrop, may be derived from the weathering of Karoo rocks. Stone artifacts have been found within the gravels on Hamburg 82 HO and Maraetchesfontein 54 HO.

4.6.2 Calcrete

Major occurrences of outcropping calcrete have been indicated in Enclosures 1 and 2. The calcrete occurs in the form of hardpan calcrete, less than one metre thick, and as massive calcrete, up to 4,5 m thick, in river valleys and pans. This hard rock is generally white to cream in colour, fine-grained, pseudo-bedded and contains occasional fractures.

Downward migration of calcium rich solutions has caused the local calcification of underlying gravels and weathered bedrock.

4.7 Lineations

The interpretation of aerial photographs and aeromagnetic maps revealed the presence of some 280 km of traceable lineaments. These lineations represent structural features, i.e. faulting and fracture zones, some of which have been intruded by magmatic material.

The compass rose plot presented in Fig. 18 demonstrates the predominately north-south alignment of these features. The major lineation directions are $170^\circ - 180^\circ$ and $000^\circ - 010^\circ$ with another well developed set at $030^\circ - 040^\circ$. Although east-west lineations are all but absent, an airborne magnetic survey with north-south flight paths may have revealed the presence of more features in this direction.

Exploratory drilling has been carried out on four of these features with the following results:

- three represent dykes of which two are magnetic and the other not, (Maraetchesfontein 54 HO and Schweizer-Reneke Townlands 62 HO), and
- one represents locally magnetic amphibolite-pyroxenite (Hamburg 82 HO).

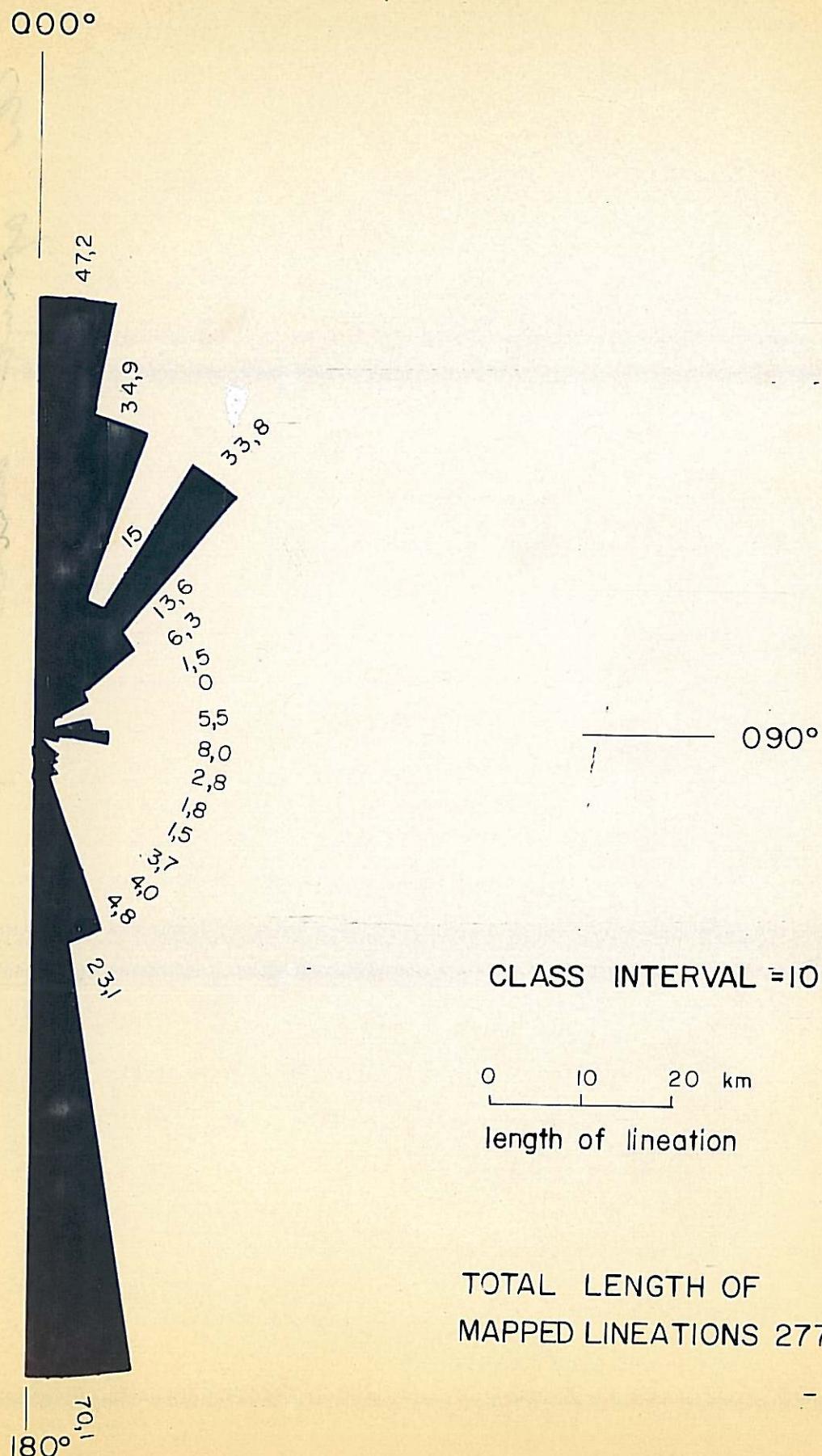
For convenience of description two dykes have been locally named i.e. Townland Dyke and the Maraetchesfontein Dyke.

5. GENERAL GEOHYDROLOGY

Information for this section was obtained from the following sources:

- records of the Boring Division of the Directorate of Water Affairs;
- Steyn and Venter (1980);
- partial borehole survey conducted by Bekker, Hoffmeyr and Taylor during August-December 1981 (borehole survey forms in Appendix 1);
- logging of 21 specifically drilled exploration boreholes, September-December 1981 (borehole profiles presented in Appendix 2 and Figs. 1 to 14).

FIG. 18 COMPASS ROSE PLOT OF VISIBLE LINEAR
GHP 5659 FEATURES PRESENT ON AERIAL PHOTOGRAPHS
AND AEROMAGNETIC MAPS



5.1 Water-bearing properties of aquifers

The rural population is dependent almost exclusively on ground water for domestic and stock watering purposes. In favourable areas ground water is used for irrigation (see para. 5.2). During 1980 ground water was used to augment the municipal water supply of Schweizer-Reneke.

Conclusions on the water-bearing properties of the aquifers present are based on the survey of 212 boreholes in the area and 21 exploration boreholes drilled in the present investigation (details summarised in Table 5).

Thirty-six private boreholes have yields reported by their owners in excess of 5 l/s (see Table 6) although only nineteen have been pumped at rates greater than 5 l/s (see Table 7). During 1981 only twelve boreholes were regularly being pumped in excess of 5 l/s. Five government boreholes drilled in 1981 had yields greater than 5 l/s.

Yields = blow yields except for pumping yields as indicated as * ; 1 = levelled

TABLE 6: BOREHOLES WITH REPORTED YIELDS GREATER THAN 5 l/s CLASSIFIED ACCORDING TO FORMATION
 N.B. () = reported water strike (m)

AQUIFER	R	Rb-R	Rb	Ra-Rb	Ra	Karoo-Ra	Karoo	Total
5 to 9,9 1/s	5 - Ma 61/2 (25-) Ma 61/3 (25-) Ma 61/10 (25-) Ss 20 (30-) G34312 (13 to 26)	2 - Gn 2 (?) S1 2 (37-)	1 - Je 1 (42)	3 - Ma 61/9 (36-) Rt 12 (30-) G34315 (17, . 33 + 60)	3 - Lt 49/2 (30-) Lt 49/4 (30-) G34308 (50)	4 - Dk 1 (23) Mn 12 (17) Sk 8 (58-) Sk 9 (?) !	None	18
10 to 14,9 1/s	7 - J1 3 (42) Ma 61/1 (25-) Pa 1 (36) Rt 1 (30-) Ss 4 (14 to 21) Ss 5 (36-) Ss 8 (39-)	None	None	None	2 - Ln 1 (58) G34307 (71)	3 - Kl 2 (60) Nt 86/2 (30) Sk 5 (37-)	None	12
15 to 19,9 1/s	5 - Hg 1 (18) Pa 5 (20) Ss 2 (13 to 21) Ss 3 (13 to 23) G34310 (11 to 22)	1 - Gn 1 (40-)	1 - Rt 2 (?)	None	None	2 - Sk 3 (46-) Sk 4 (58-)	None	9
20 to 24,9 1/s	None	None	None	None	None	1 - Kl 1 (70)	None	1
Over 25 1/s	None	None	None	None	None	1 - Sk 1 (58)	None	1
TOTAL	17	3	2	3	5	11	0	41

TABLE 7: BOREHOLE THAT HAVE BEEN PUMPED IN EXCESS OF 5 l/s CLASSIFIED ACCORDING TO FORMATION
 N.B. * boreholes pumped regularly in 1981 (total = 12)

AQUIFER	R	Rb-R	Rb	Ra-Rb	Ra	Karoo-Ra	Total
5 to 9,9 1/s	8 - Ma 61/1* Ma 61/2* Ma 61/3* Ma 61/10* Rt 1* Ss 2 Ss 3 Ss 20	1 - Gn 2*	1 - Rt 2	1 - Ma 61/9	2 - Ln 1 Lt 49/2	1 - Mn 12*	14
10 to 14,9 1/s	1 - Hg 1	1 - Gn 1*	None	None	None	2 - Kl 2* Sk 1*	4
Over 15 1/s	1 - Pa 5*	None	None	None	None	None	1
TOTAL	10	2	1	1	2	3	19

5.1.1 Water-bearing status of the Archaean granite gneiss (R)

Within the area studied this formation can be generally regarded as a good aquifer capable of supplying adequate yields from shallow depths for domestic and stock watering purposes. Problems of water supply from this formation are experienced where extensive outcrops of unweathered granite occur, principally on the townlands and in the south-west of the area.

Of the 87 boreholes investigated 43% had yields reliably assessed as greater than 1,0 l/s (see Table 8) while 21% yielded more than 5 l/s.

Groundwater is usually abstracted from zones of weathering and fracturing at shallow depths (see Tables 5, 6 and 8). When weathered granite is either absent or unsaturated deeper boreholes (over 30 m) have been drilled with little success. The lack of success is due to fractures being generally absent at depths greater than 30 m. Locally quartz veins and pegmatites encountered in deeper drilling, may be fractured and may yield useful water supplies. Near surface fractured amphibolite-pyroxenite bands may improve the chances of obtaining adequate supplies of ground water. Experience gained during the drilling programme indicated that dyke contact zones do not appear to provide particularly favourable drilling sites.

5.1.2 Water-bearing status of the Bothaville Formation (Rb)

Primary porosity is virtually zero within these arenites thus the storage and movement of ground water is controlled by the presence of secondary openings, namely interconnected joints and fractures. Available borehole information suggests that the upper Bothaville arenites are better ground water yielders than the lower Bothaville arenites. This change in aquifer status is related to the fewer joint and fracture openings present in the lower part of the formation. The relative abundance of bedding plane joint and fracture openings in the upper sequence is tentatively correlated with deep chemical weathering (up to 80 m) accentuating differences between thinly and well-bedded arenaceous units.

Of the 11 boreholes drilled within the Bothaville Formation, 45% had yields in excess of 1 l/s and 18% greater than 5 l/s, while three boreholes were drilled well below the ground water level, without striking water. The drilling programme indicated that several individually low yielding water interceptions are generally made by boreholes drilled within this formation.

TABLE 8: BOREHOLE STATISTICS

AQUIFER	R	Rb/R	Rb	Ra/Rb ^{***}	Ra	Karoo/Ra	Karoo
No. of boreholes investigated	87	8	11	11	61	55	None located that have been drilled only within Karoo Formation
No. of boreholes in use	63	8	3	7	46	38	
Average depth of borehole (m)	31 (54)	43 (3)	41 (7)	61 (9)	53 (23)	50 (23)	
Average yield of boreholes 1/s*	2,9 (74)	3,9 (8)	1,6 (10)	3,5 (11)	0,7 (41)	2,4 (50)	
No. yielding 1 l/sec + as % of total no. of boreholes	37	5	5	8	12	24	
No. Yielding 1 l/sec + as % of total no. of boreholes	43	63	45	73	20	44	
Average depth of water interceptions (m) **	23 (30)	38 (2)	32 (3)	42 (7)	33 (21)	38 (12)	
Average depth of water level below collar (m)	7,7 (44)	10 (3)	13,5 (5)	12 (6)	10 (33)	6,5 (34)	

Average based on number of returns in brackets

* Windpump yields taken as 0,1 l/s, unless other information available and pumping yields of boreholes used rather than tested yield.

** In cases where multiple water interception reported deepest water interception used.

*** Average depth at which contract penetrated is 35 m (6 returns).

The highest yielding borehole Rt 2 was drilled into a fracture zone. It is expected that boreholes drilled into dyke contact zones and lineaments may have improved yields.

5.1.3 Water-bearing status of the Allanridge Andesite Formation (Ra)

Boreholes drilled exclusively within this formation generally have poor yields (see Table 8). Sixty-one boreholes have been investigated, of which only 20% yield more than 1,0 l/s and only 8% in excess of 5 l/s.

The water-bearing zone generally exploited in this formation is the weathered zone which is generally sufficiently permeable, at depths between 15 and 30 m, to produce yields of 0,1 to 1,0 l/s. Where the weathered zone is absent, unsaturated or produces insufficient ground water deeper drilling has often been successful in locating increased quantities of ground water in deep fractures e.g. borehole Ln 1-13 1/s at 58 m and Ma 61/6 - 2,5 l/s at 71,5 m. Some boreholes penetrating fracture zones associated with lava-dyke contacts have encountered good yields e.g. G34308 blow yield 6 l/s.

5.1.4 Water-bearing status of the Karoo Sequence

Within the area studied no boreholes were located that were drilled exclusively within the Karoo Sequence. It is expected that the mudstones are very poor aquifers and that their principal importance is their ability to store ground water, which under pumpage will leak into the underlying lava aquifer.

5.1.5 Water-bearing status of the alluvial deposits

Most of these deposits are not of importance as aquifers as they are thin and usually lie above the present zone of saturation. However the superficial deposits do function as efficient ground water recharge areas for the underlying water-bearing formations.

5.1.6 Influence of formational contacts on water-bearing status

Borehole statistics (Table 8) suggest that boreholes penetrating one of the three major lithological contacts viz arenite-granite, lava-arenite and mudstone-lava have generally higher average yields than those boreholes that are drilled exclusively in one formation. For this reason the water-bearing properties of these contact zones have been considered separately from the principal water-bearers enumerated above. These formational contact zones are discussed below:

Bothaville Formation - Archaean granite gneiss (Rb-R)

This contact zone seems to be a good water bearer. Of the eight boreholes investigated five yield more than 1 1/s and three in excess of 5 1/s.

Due to favourable topography the arenite-granite contact has been successfully exploited north of Schweizer-Reneke but elsewhere access to suitable drilling positions on this contact is difficult.

Allanridge Andesite Formation - Bothaville Formation (Ra-Rb)

The eleven recorded boreholes drilled through this contact have been fairly successful of which eight yield over 1 1/s and three in excess of 5 1/s.

The average depth at which the lithological contact has been penetrated is 35 m, however the depth of penetration does not seem to greatly influence borehole yields. Borehole logs, Fig. 9 to 12, illustrate that ground water has not been struck directly on the lithological contact.

Karoo Sequence - Allanridge Andesite Formation (Karoo-Ra)

As up to 50 m of Karoo mudstone overlie much of the lava subcrop the mudstone has had to be penetrated to obtain ground water supplies for domestic and stock watering purposes.

In many cases good success has been achieved. Fifty-five boreholes are recorded of which 44% yield more than 1 l/s and 20% in excess of 5 l/s.

These good overall results have been attained despite the absence of weathered underlying lava in areas with more than about 15 m of Karoo cover. Successful boreholes are due to the penetration of interconnected fracture networks within the otherwise fresh lava. The deepest depth of fracturing recorded is at borehole K1 1 which intercepted some 23 l/s at a depth of 70 m.

TABLE 9: DETAILS ON PRINCIPAL GROUND WATER CONSUMERS, 1980/81 SEASON

Farm	Owner	Distance from Schwizer- Renkele (Km)	Pumping boreholes	Pumping rate (1/s)	Pumping period (hours)	Total quantity of water abstracted 10^6 m^3	Irrigated area (hectares)	Number of irrigation scheme	Formation
Groenpan			Gn 1	13,8	- 2 900				
23 HO	Le Roux	14,5	Gn 2	6,0	variable	0,144 plus	8,6	No	Rb/R
			Gn 3	3,7	variable				
Kameelkuil	Van der Merve	18,0	Kl 2	12,5	4 960	0,223	17	No	Karoo/Ha
88 HO									
Lot 7 49 HO	Deale	6,5	Lt 49/2	6,0	2 000?	0,072	3	No	Ra
			Lt 49/4	4,0					
Mimosa 61 HO	Van Rensburg		Ma 61/1						
		1,0	Ma 61/2	combined					
			Ma 61/3	25,0 *	1 300 *	0,117 * plus	25,7	Yes	R
			Ma 61/10	plus					(Principal consumer)
Maraatches-									
fontein 54 HO	Viljoen	12,5	Mn 12	8,8	2 090	0,066	2,6	No	Karoo/Ha
Palachooma	Visser		Pa 4	2,5	~	3 440	0,031)	Yes	
64 HO	(now Roos)	5,0	Pa 5	15,0	(12,5) **	4 770	0,258)	12	(not used since R 1979)
			Pa 6	3,1		3 440	0,038)		
Rietput 60 HO	Reyneke	0,5	Rt 1	10,0 (6) **	1 100	0,040	10	Yes	R
									(not used since 1979)
Springbok	L.T.R. Holdings	17,5	Sk 1	14,4	2 190	0,114	6,5	No	Karoo/Ha
191 HO									

* Actual pumping rate declined from mid 1980 to end 1980 and pumping period extended.

** Pumping rate declined from mid 1980 to end 1980, pumping period reduced.

5.2 Principal consumers of ground water during 1980/81

Table 9 provides information on ground water consumption by irrigators for the 1980/81 agricultural season (refer to Enclosure 1 for borehole locations and Appendix 1 for borehole data). The total quantity of ground water abstracted by the eight main irrigators during 1980/81 was approximately $1,1 \times 10^6 \text{ m}^3$.

Where owners noticed declining yields from boreholes during 1980 this has been noted in Table 9. However as most farmers do not regularly measure borehole discharges the "set" pumping rate has been used to calculate ground water abstractions. Seasonal consumption fluctuates depending on the frequency and intensity of rainfall events and the availability and reliability of irrigation water from Wentzel Dam.

The pumpage listed for irrigator Van Rensburg is the minimum average amount annually required from ground water. If water is unavailable from the Wentzel Dam, or the privately owned dam on the Harts River, (river point 4) additional ground water abstraction is required to meet irrigation requirements, estimated at some $0,4 \times 10^6 \text{ m}^3/\text{a}$. Other former surface water users' on Rietput 60 HO and Palachoema 64 HO now obtain their water requirements directly from ground water. In previous years ground water for irrigation has been pumped on Hamburg 82 HO and London 112 HO, boreholes Hg 1 and Ln 1 respectively. However this practise has now ceased for economic reasons. In the near future pumpage will probably commence on Palachoema 64 HO from boreholes Pa 1 and Pa 2.

5.3 Piezometry

A piezometric contour map, refer to Enclosure 2, was compiled with 10 m contour intervals from ground water level measurements made during the period September - December 1981. 139 measurements were made, of which 40 had to be made while windpumps were operational. The collar elevations of most boreholes were estimated from the relevant topocadastral maps, although a few boreholes were levelled (see Table 5 and Appendix 1). A geohydrological profile along section line A-B is shown in Fig. 17.

Despite the low density of measurements, namely one measurement per 6 km^2 , and the absence of data from such critical areas as the Harts River valley and the Townland Stream valley, certain general observations can be made of the ground water conditions prevailing in late 1981, viz:

- ground water flow within the Harts River basin is from the major ground water divides, with a gradient of 1×10^{-3} , towards the central portion of the valley, from where ground water movement is in a south-westerly direction under a shallower hydraulic gradient of 1×10^{-4} ;
- a ground water divide coincides with the Vaal-Harts surface water divide. South of this divide ground water movement is in a southerly direction with a shallow gradient in the order of 2×10^{-4} ;

- the poorly defined surface water divide between the Dry Harts and Harts River catchments is not coincident with the ground water divide. North of the ground water divide ground water movement is in a north-westerly direction;
- piezometric contours indicate that effluent conditions prevailed in the channel of the Harts River, i.e. the Harts River was receiving discharge from the water table. Discontinuous stretches of standing water in the channel confirms this;
- dykes and other lineations do not appear to influence the direction of ground water flow.

5.4 Periodic water level measurements

The ground water levels in two boreholes are at present being monitored namely:

C3N046 (Schweizer-Reneke Townlands 62 HO) - monthly automatic water level recorder established in July 1980, an old municipal borehole Ss 4, and

C3N047 (Maraetchesfontein 54 HO) - H.W.K. meter read monthly, established in December 1981 on boreholes G34309.

As both observation boreholes have only been operational for a short period of time, no hydrographs have been compiled. Ground level measurements taken at the two recorder stations C2N046 and C2N047 are tabulated in Appendix 1.

Due to poor installation and servicing the data available from C3N046 prior to October 1981 should be viewed with circumspect. However the dewatering of the municipal well field is clearly shown by water level measurements in C3N046 as is the rapid recovery of the aquifer after municipal ground water pumpage ceased. Recent hydrographs show that there is a clear correlation between water level rises in C3N046 and surface water flow in the nearby Harts River (200 m distant with an elevation difference of 5 m).

The 0,77 m rise in water level recorded by C3N047 between October 1981 and February 1982 must presumably have been caused by recent rains in the area.

The reappearance of seepage points along the bed of the Hamburg stream, some 200 m west of G34313 and the flowing surface springs on Palachoema (S1) indicate that at the end of 1981 ground water levels were high and had completely recovered from the effects of subnormal rainfall over the period 1978-1980. This statement is further supported by the full recovery of those boreholes which experienced declining discharges during the latter half of 1980.

5.5 Drilling programme

Information on the exploration boreholes drilled is presented in Table 5, summarised borehole logs in Appendix 2, Figs 1 to 14, and complete borehole logs in G.H.P. envelope No. 635. Rock chip samples were collected during drilling and are available for inspection in Pretoria Head Office.

All boreholes were drilled using a Rock Giant air percussion rig supplied by the Vryburg Boring Inspectorate. Over the period 1981-09-21 to 1981-12-08 twenty-one boreholes were drilled in eight target areas. The combined meterage drilled was 1 017 m. For 57 working days this is equivalent to an average drilling rate of 18 m/d. 270,6 m of mainly 165 mm diameter steel casing was installed in the completed boreholes.

Of the 21 boreholes drilled five had yields in excess of 5 l/s while seven boreholes can be classed as potential production boreholes.

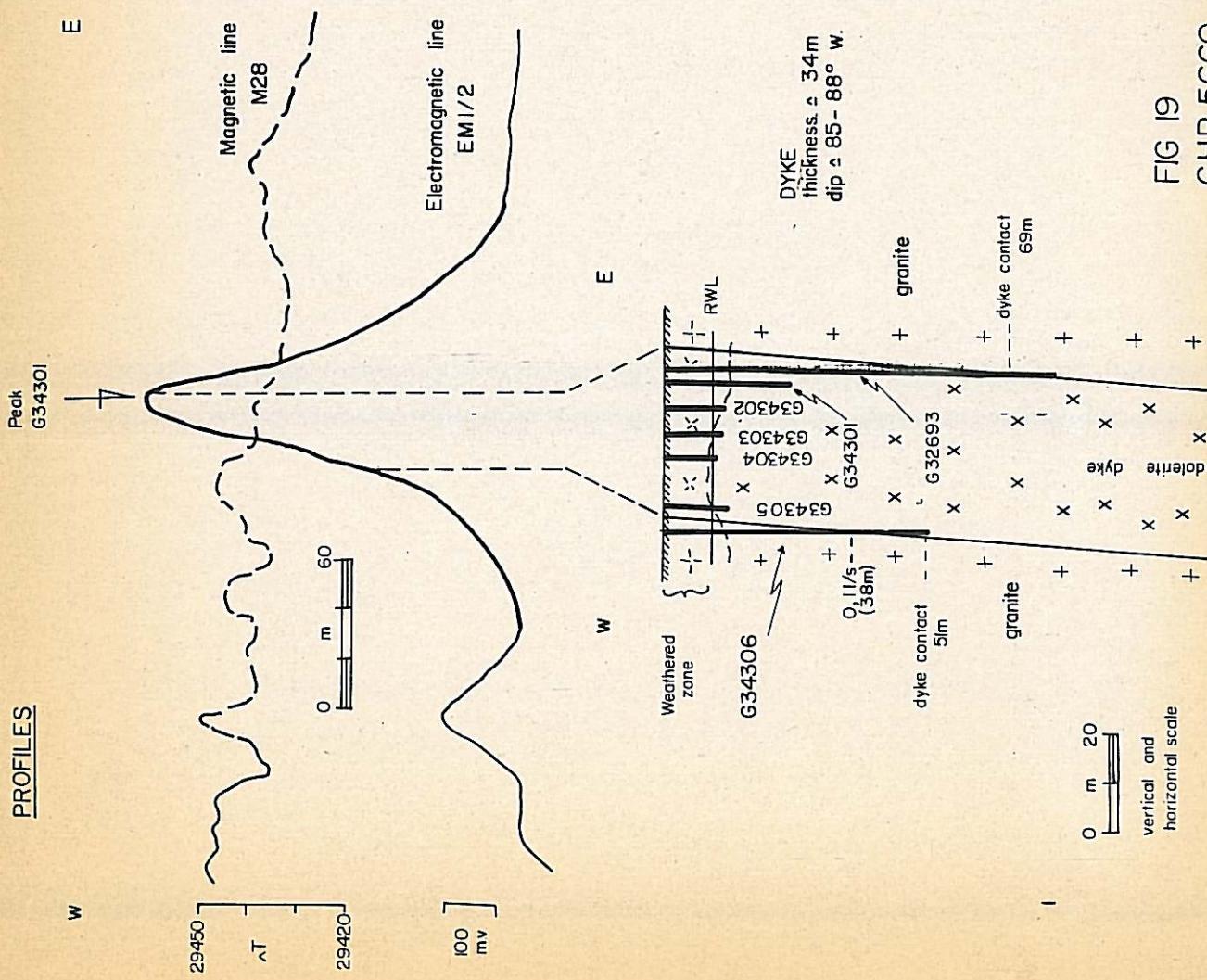
5.5.1 Drilling difficulties and recommendations

Progress was fairly rapid as drilling conditions were generally good. The weathered rock, particularly the granite, is rapidly drilled by air rig. In all cases steel casing had to be inserted during drilling to the weathered rock-fresh rock interface. Where necessary this casing was perforated on site using an oxy-acetylene

cutter. Borehole G34314 had to be abandoned as near-surface broken lava was not adequately cased off. The test pumping of borehole G34310 indicated that if heavy pumping of G34310 is planned the borehole should be reamed and fitted with a gravel pack (see section 8.2.1 for details) to prevent sand pumping and consequent damage to pumping equipment.

Progress through Karoo mudstone was slow, by air rig, as mudcake continually blocked the air vents in the drilling bit. This formation required complete casing off.

Drilling rates through the various hard rocks encountered were fairly uniform. However the Bothaville Formation presented a special problem as the exceptionally hard quartzite units were found to significantly grind down and thus reduce the diameter of the drilling bit. Consequently the diameter of boreholes penetrating the Bothaville Formation progressively decreased with depth e.g. from 165 to 154 mm approximately in G34315. Units of friable sandstone occurring within the Bothaville Formation may require complete casing off or grouting. When drilled these units appear as sand and local drillers have reported that this sand can cause the partial collapse of boreholes and damage to pumping equipment.



SITE PLAN

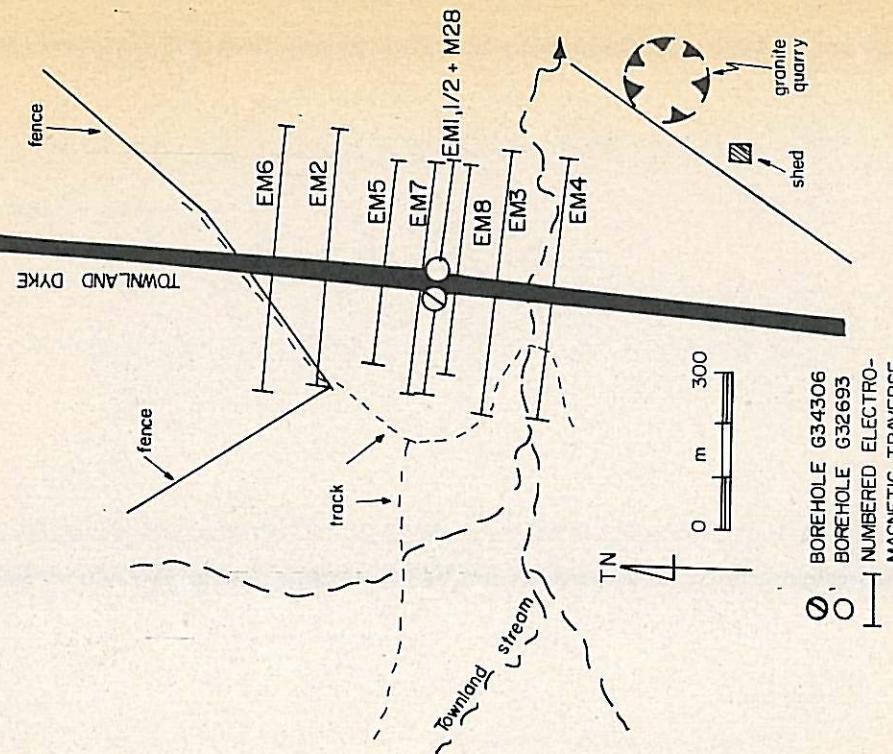
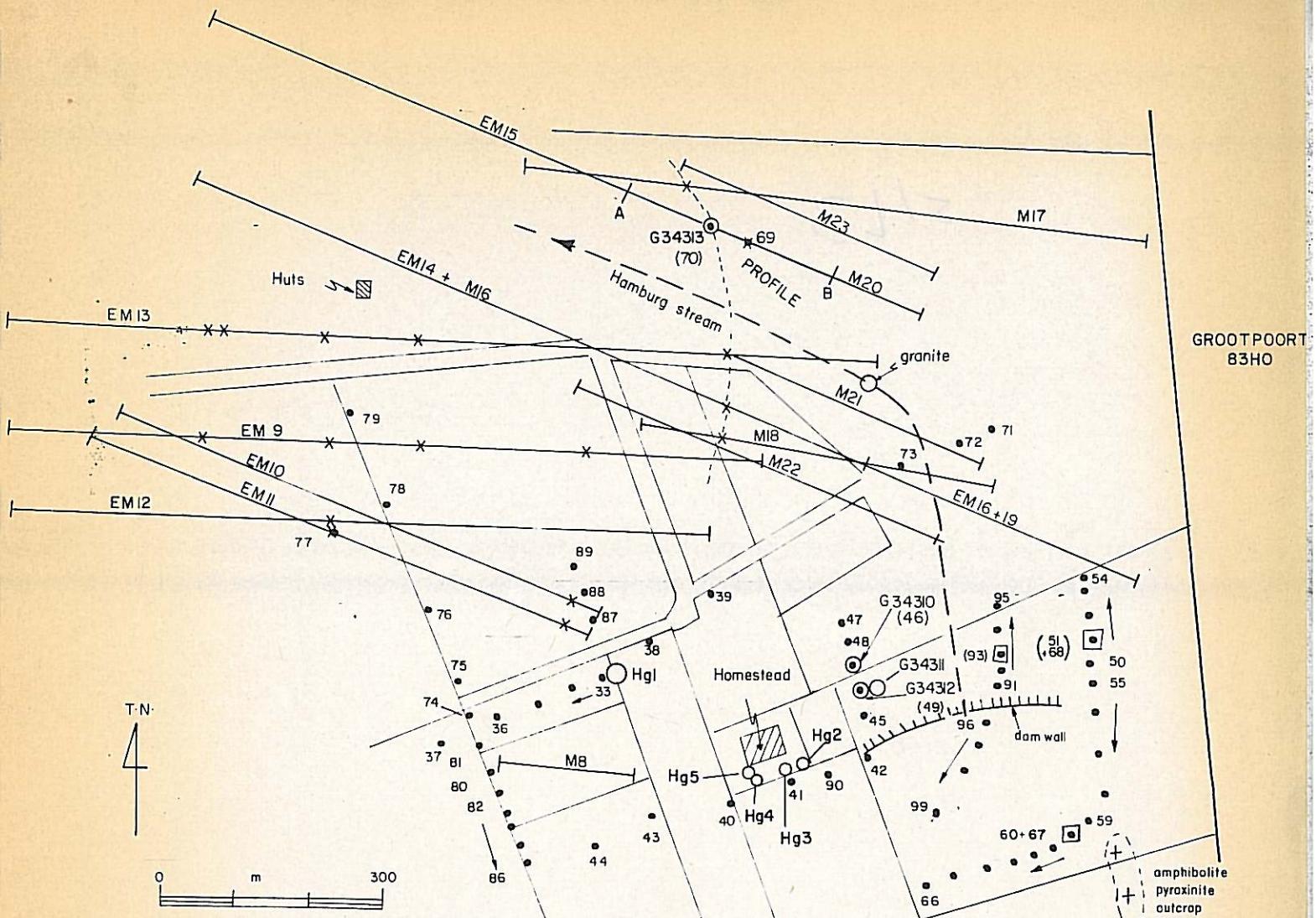


FIG 19
G.H.P. 5660

LOCATION OF BOREHOLE AND GEOPHYSICS
ACROSS TOWNLAND DYKE, SCHWEIZER-RENEKE
TOWNLANDS 62 HO.

FIG 20 LOCATION OF BOREHOLES AND GEOPHYSICS
GHP 5666 ON PART OF HAMBURG 82HO



Hg1 O Numbered borehole

Recommended borehole sites

Numbered electromagnetic (EM) and magnetic (M) traverses

• Numbered Wenner electrical resistivity sounding

fence

X prominent em and m feature
 x--x magnetic feature

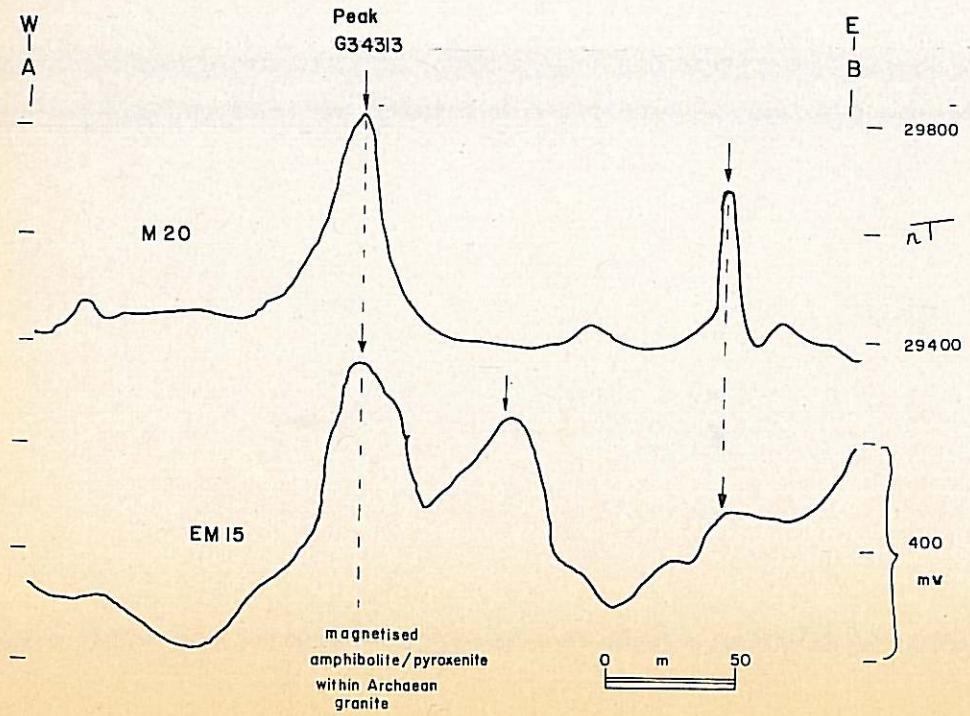
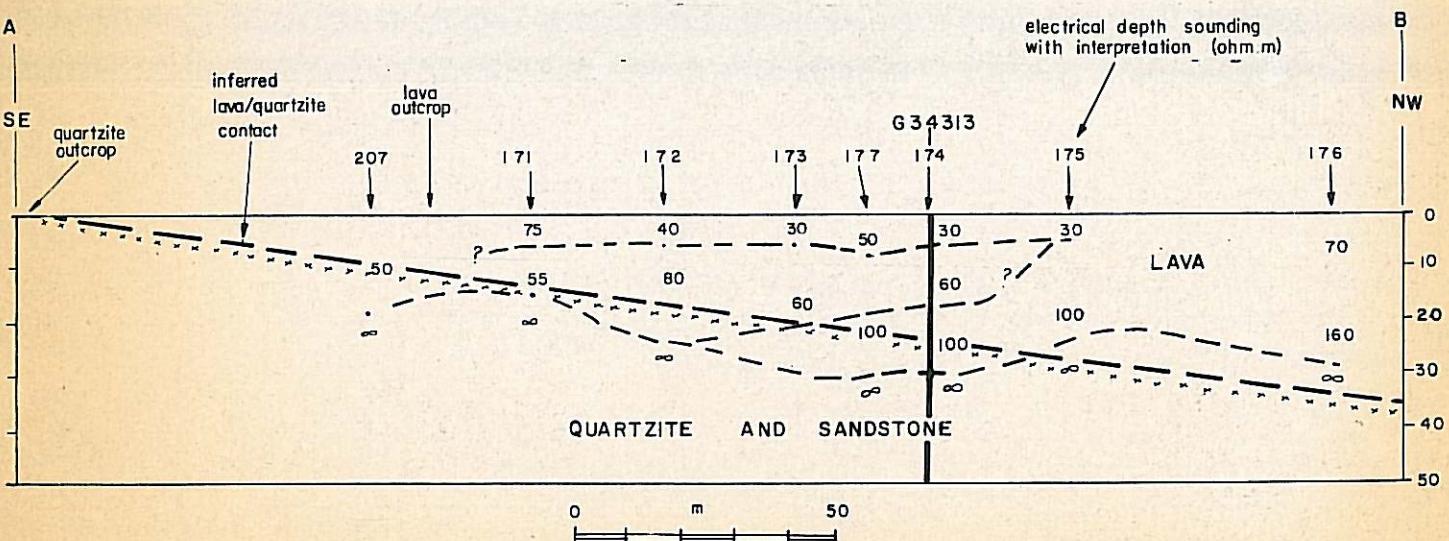
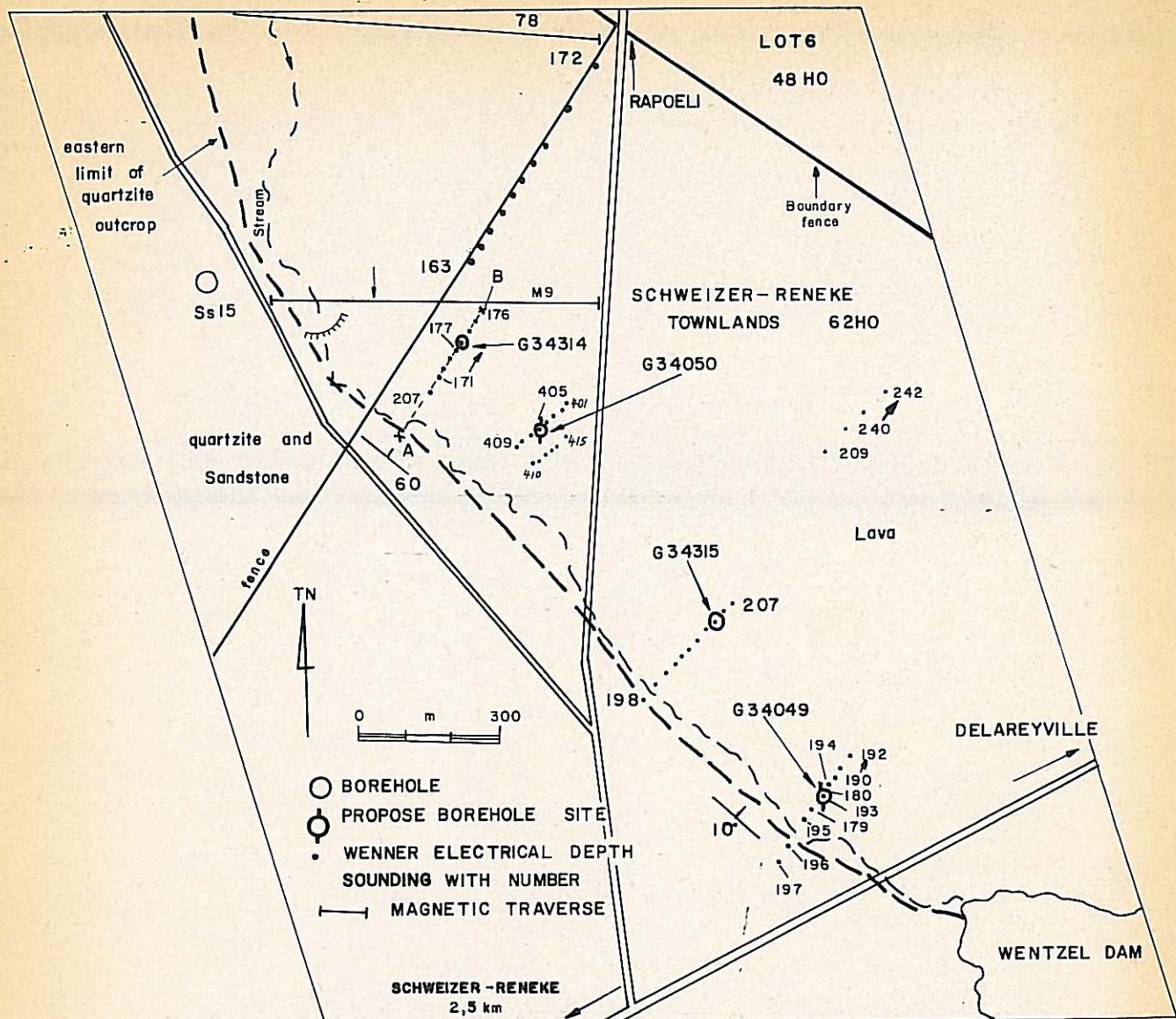


FIG. 21. LOCATION OF BOREHOLES AND GEOPHYSICS
GHP 5667 ON LAVA/QUARTZITE CONTACT.
SCHWEIZER - RENEKE TOWNLANDS 62HO



GEOPHYSICAL PROFILE A-B

FIG. 22 LOCATION OF BOREHOLES AND GEOPHYSICS ON
G.H.P. 5668 PART OF MARAETCHESFONTEIN 54HO.

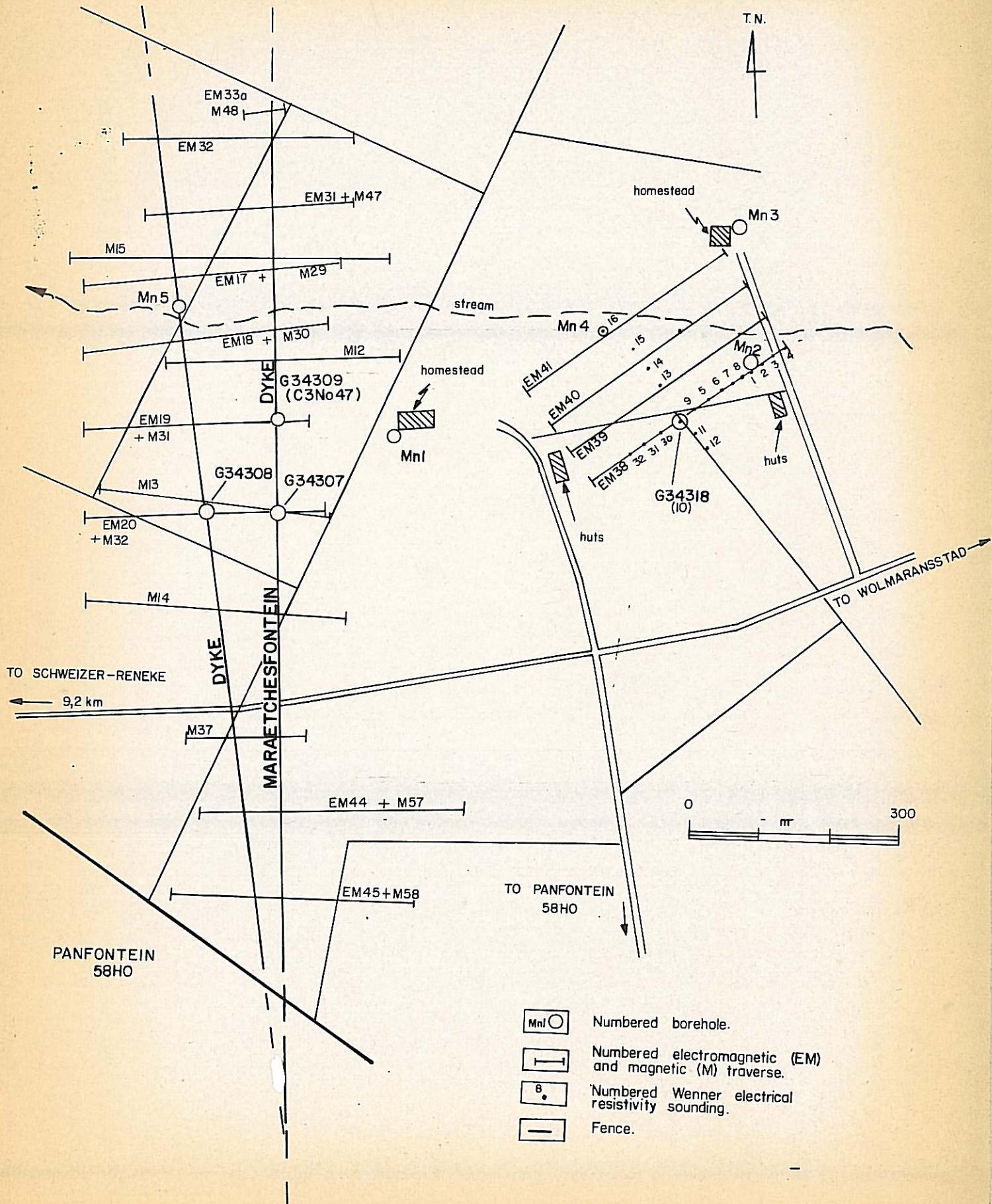
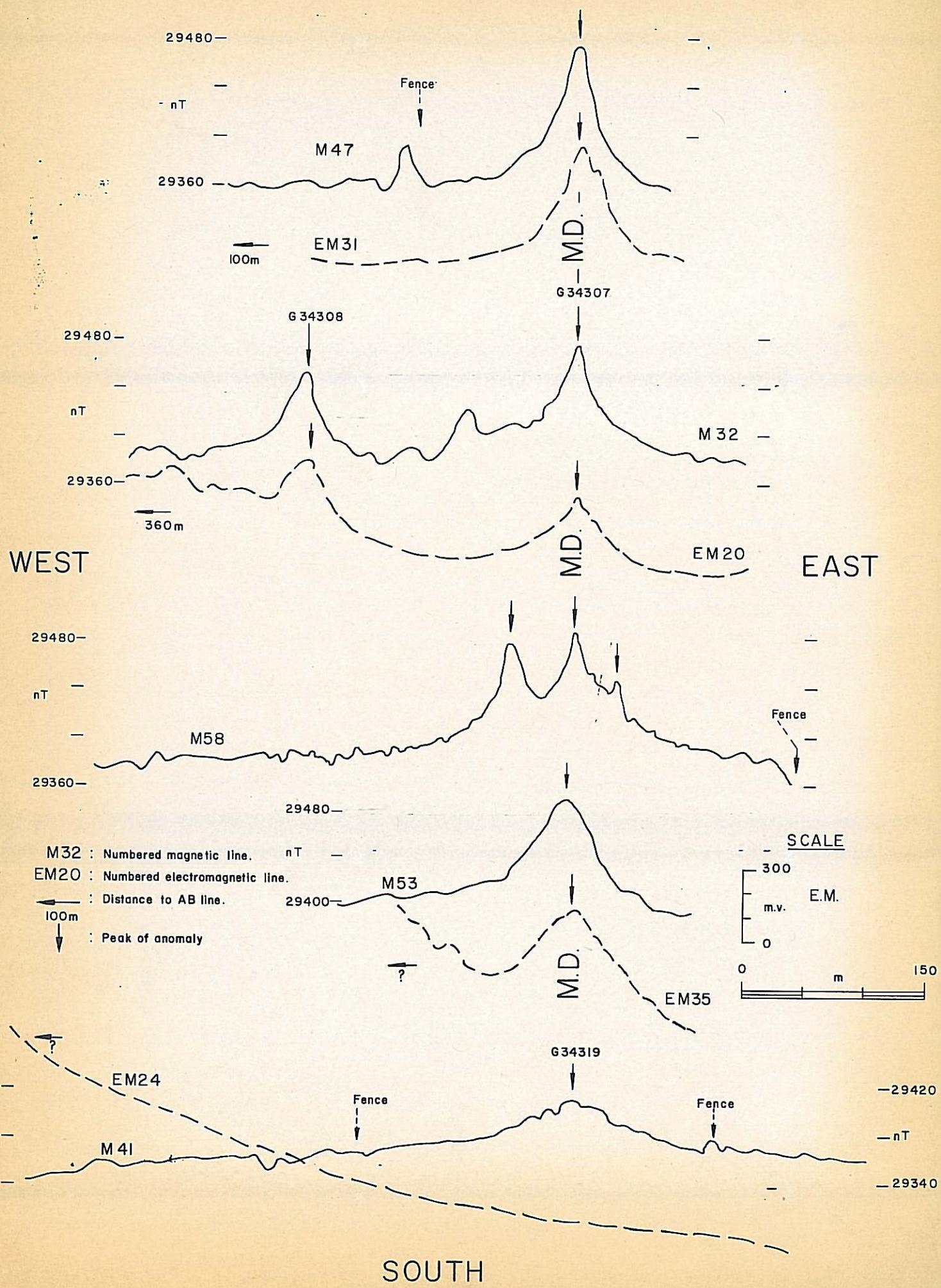


FIG. 23
G.H.P. 5669

MAGNETIC AND ELECTROMAGNETIC PROFILES
ACROSS MARAETCHESFONTEIN DYKE(MD) ON
MARAETCHESFONTEIN 54HO AND PANFONTEIN 58HO



6. GEOPHYSICS

6.1 General

Surface geophysical techniques were employed to investigate the favourable areas for ground water occurrence which had been previously identified as target areas by the indirect methods mentioned in section 1.5. The results obtained were used to select sites for exploration drilling. The following geophysical work was undertaken in the course of the investigation:

- 59 ground-magnetic traverses
- 47 electromagnetic traverses, and
- 452 electrical resistivity depth soundings.

In addition 20 boreholes were electrically logged.

Enclosure 1 and Figs. 19 to 23 show the location of the geophysical work, while field measurements, curves and profiles are available for inspection in G.H.P. envelope No. 635.

6.2 Methods employed

Ground-magnetics - The Chemtron Proton Magnetometer (models G2 and G3) was used. These instruments measure the total magnetic field. Traverse lines were short, maximum length 1 300 m, and as they were rapidly completed no corrections were necessary for diurnal drift. Distances between stations were paced, one pace being approximately 1 metre. Station intervals were 10 paces which was reduced to 2,5 paces over features of interest. The magnetic method was useful in detecting lineaments caused by magnetic dykes and magnetised amphibolite-pyroxenite bands. It could also be used to detect the very approximate surface position of the lava-quartzite contact.

Electromagnetics - The Chemtron models G30 and G31 were used, current being supplied by a separate alternating current source. Traverse lengths of up to 800 m were made when field conditions permitted. Station intervals were similar to those employed for the magnetic traverses. The electromagnetic technique accurately located

conductive lineaments, both magnetic and non-magnetic. It also indicated broad areas of weathering both within the granite and the lava.

Electrical resistivity depth soundings - The equipment available for making measurements of apparent electrical resistivity was the Gane-Enslin a/c field instrument. As the most appropriate information required was the relative depth to the underlying solid bedrock the Wenner electrode configuration was employed. Maximum spacings did not exceed 72 m and were commonly less than 42 m. The sounding curves produced were interpreted using the empirical method. The electrical resistivity method was successfully used to identify zones of deep weathering within the granite and the presence of Karoo mudstones. Limited success was achieved in identifying the lava-quartzite contact.

Borehole electrical resistivity logging - Resistivity readings were made using the Gane-Enslin instrument. A lateral electrode arrangement was employed with an M-N distance of 1 m. Boreholes were logged both down-hole and up-hole. The prime purpose of the logging was to check borehole logs, as deduced, from drilling chips, with geophysical information. Where necessary corrections were made to lithological logs. Information was also provided on the resistivity ranges of the lithological formations present, see Table 10.

TABLE 10: TYPICAL ELECTRICAL RESISTIVITY RANGES OF GEOLOGICAL FORMATIONS

Formation	Ohm.m
Karoo mudstone	20 to 40
lava, partly weathered	400 to 1 200
lava, fresh	1 000 to 9 000
quartzite	500 to 8 000
granite, weathered	100 to 300
granite, partly fractured	250 to 600
granite, fresh	3 000 to 12 000
amphibolite, fresh	10 000 +
dolerite, fresh	5 000 +
diabase, fresh	2 000 +

6.3 Results

6.3.1 Archaean granite gneiss

Townland stream valley - The electromagnetic technique was used to investigate a north-south, non-magnetic lineament on Schweizer-Reneke Townlands 62 HO. Profiles across this feature showed a prominent conductive zone flanked by minor subsidiary, positive anomalies (refer to Fig. 19). Exploration drilling (see Figs. 1 and 2) demonstrated that the prominent anomaly was caused by a dolerite dyke some 34 m thick, dipping at 85 to 88° to the west. This dyke was named the Townland Dyke.

Electrical resistivity depth soundings were carried out both upstream and downstream of the Townland Dyke. The maximum thickness of weathering detected was only 10 m, thus much of the weathered zone is currently above the present ground water table.

Ground-magnetic traverses were used to investigate a north-east to south-west trending low amplitude, positive aeromagnetic anomaly some 1 000 m east of the Townland Dyke. Traverses M10 and M11 failed to identify this feature on the ground however.

Hamburg 82 HO - Electromagnetic and magnetic work was initially concentrated on locating the north-south, low amplitude, positive, aeromagnetic feature that was initially thought to be associated with borehole Hg 1. South of Hg 1 a broad positive anomaly of some 80 nT amplitude was located by magnetic traverse M5. While north-west of Hg 1 a prominent narrow magnetic anomaly was located (refer to Fig. 20) the magnitude of which varied from 300 nT (M17) to over 1 200 nT (M18). Magnetic profiles indicated a steep dip to the west. A second smaller anomaly of some 250 nT was detected some 140 m to the east on Line M20. Magnetic measurements over an amphibolite outcrop indicated that this lithology was strongly magnetised. Exploration drilling (Fig. 8) on the peak of the large anomaly recorded by traverse M20 confirmed that this feature was caused by subcropping magnetised amphibolite-pyroxenite. Electromagnetic traverses detected a conductive zone coincident with the magnetic feature. Other broad electromagnetic anomalies are probably caused by areas of deep weathering within the granite.

Electrical resistivity soundings were used to locate the areas of deepest weathering. This was shown to be present in the valley of the Hamburg stream where depths of up to 30 m were interpreted from resistivity sounding curves. The successful borehole G34310, Fig. 6 was drilled on sounding No. 46 which indicated weathered material to 25 m with a resistivity value of some 115 ohm. m. Depicted on Fig. 20 are three recommended drilling sites.

Stream confluence, Mimosa 61 HO - The electrical resistivity sounding curves indicate that the granite is deeply weathered along the stream valleys. The depth of weathering confirmed by drilling (see Fig. 14) is up to 30 m with a corresponding saturated thickness of some 18 m. The resistivity range of the weathered rock is similar to that of Hamburg ie. 100 to 180 ohm.m.

Common boundary of Grootpoort 83 HO and Mimosa 61 HO - The possible northerly extension of a near vertical, north-south, dolerite dyke was investigated. Magnetic traverses indicated that the feature, if present, was non-magnetic. A positive electromagnetic anomaly recorded by profile EM 21 suggested that the dyke does continue northwards. However no further geophysical work was carried out across this dyke as access to the property was not permitted by the owner.

6.3.2 Contact between Bothaville and Allanridge Andesite Formations

Due to the low magnetic contrast and the low formation dip magnetic traverses were unsuccessful in locating the contact between lava and quartzite. The constant electrode spacing traverses on Rietput 60 HO were also unsuccessful in locating the contact. Therefore numerous electrical resistivity traverses lines were carried out across the approximate position of the formation contact and perpendicular to the expected strike. Sounding intervals were 25 to 30 m with AB lines aligned parallel to the strike.

The results obtained were somewhat variable. Where surface cover was thin, less than 2 m, sounding curves successfully located the formation contact by distinguishing solid quartzite from weathered

lava overlying solid quartzite. The dip of the contact could not be assessed from the empirical interpretation of the two- and three-layer sounding curves.

Four boreholes (see Figs. 9 and 12) were drilled through the lava-quartzite contact where depth soundings indicated favourable weathering within the overlying lava. The results of this drilling, in collaboration with observed rock exposures, enabled the dip of the contact to be estimated. As the exploration drilling was moderately successful two proposed borehole sites have been selected on the townlands (see Fig. 21, sites No. G34049 and G34050).

6.3.3 Allanridge Andesite Formation, including areas covered by Karoo Sequence

Geophysics provided information on the following:

- depth of weathering within the lava
- depth to Karoo - lava contact, and
- lineations

The major lineation investigated was the Maraetchesfontein Dyke. This dyke forms a low amplitude, north-south, positive, aeromagnetic anomaly which extends for over 110 km with only minor offsets by later north-east to south-west faulting. Geological and geophysical evidence suggests that this dyke is pre-Karoo. Ground-magnetics indicate that anomalies of 60 to 300 nT are associated with this feature on Maraethesfontein 54 HO. The anomalies are well defined and suggest that the dyke dips very steeply (80 to 85°) to the west (see Fig. 22 and 23 for typical profiles). South of the farm Maraetchesfontein 54 HO the peak anomaly of some 60 nT becomes broader and less well defined as the magnetised dyke is buried under a progressively deeper cover of younger Karoo sediments.

Ground-magnetics on Maraethesfontein 54 HO located a second, near vertical dyke trending 170 to 350° (see Fig. 22) with an anomaly of up to 80 nT. South of the common border between Panfontein 58 HO and Maraetchesfontein 54 HO the two identified dykes converge/diverge (see Fig. 22 and 23).

Electromagnetic profiles north of EM 35 detected conductive zones associated with the dykes while south of EM 35 no anomalies were identified under the presumably greater thickness of Karoo mudstones. The electromagnetic anomalies coincide within 5 m with the peaks of the magnetic anomalies. Three boreholes were drilled on these peaks into the two dykes (see Figs. 22 and 23). Chips recovered from borehole G34308, Fig. 4, suggested that the anomalies were caused by diabase dykes. Borehole G34319, Fig. 13, was drilled on the broad magnetic feature recorded on Panfontein 58 HO and penetrated 33 m of Karoo strata before intersecting lava.

Electrical resistivity traverses in the vicinity of borehole Mn 2, indicated the presence of substantial thicknesses (up to 25 m) of conductive rock (20 to 25 ohm.m) overlying infinitely resistive rock. Exploration borehole G34318 demonstrated that the conductive layers were Karoo mudstone and weathered lava. Subsequently electrical resistivity soundings were used to identify areas where a cover of Karoo rocks is present.

6.3.4 Superficial deposits

Electrical resistivity soundings on Belvedere 47 HO and Schoondal 22 HO indicated that only a thin cover of superficial deposits and weathered material, less than 12 m, overly solid bedrock.

7. HYDROGEOCHEMISTRY

7.1 General

As only a limited number of chemical analyses of ground water were available prior to this investigation an extensive ground water sampling programme was initiated with the following objectives:

- to assess the general ground water quality and potability
- to identify areas with unpotable ground water, and
- to detect any changes that might occur in ground water quality during the pumping of potential production boreholes.

Seventy-one major element chemical analyses were performed on samples collected from 62 sources. Sixty-nine of the samples were collected from existing boreholes while the other two were sampled from the Harts River.

The distribution of the sampling sites are shown in Enclosure 2 while the results of chemical analyses are tabulated in Appendix 3. To facilitate comparisons of water chemistry, sample analyses reported in mg/l were converted to milli-equivalents per litre (meq/l) and plotted on the accompanying Durov diagram, see Fig. 24 to 27.

7.2 Water quality and potability

Various chemical characteristics of the sampled ground water are summarised in Table 11. Most ground water are of good quality and with the exception of excessive nitrate concentrations satisfy the drinking water criteria of the South African Bureau of Standards (S.A.B.S.).

The majority of ground water are of the mixed cation - bicarbonate type (M-HCO₃). In only three ground water samples was a dominant cation found. However within the mixed cation field there is a tendency for ground water from the granite, Fig. 24, and from the lava overlain by Karoo mudstones, Fig. 27, to be more sodic than water obtained from the lava, Fig. 26, without a covering of Karoo. Anion types are more variable as bicarbonate (HCO₃), chloride (Cl) and nitrates (NO₃) are all important constituents. The presence of high nitrate concentrations and the difference in origin of nitrate and chloride justify the subdivision of the chloride-nitrate-fluoride field. Within this field a NO₃-Cl ratio of less than 1 is taken to designate a dominant chloride anion field, while a NO₃-Cl ratio in excess of a dominant nitrate anion field. Table 11 indicates that all aquifers can contain high nitrate concentrations (see also section 7.2.3). Low chloride-bicarbonate ratios and low sulphate concentrations indicate that most ground waters are of fairly recent origin.

TABLE II: SOME CHEMICAL CHARACTERISTICS OF THE WATER-BEARING FORMATIONS

Formation	No. of boreholes sampled	Fig. No.	Water type cation-anion (where N = mixed)	T.D.S. (mg/l) field values min. med. max.	Total hardness (as CaCO_3)	Silica concentration mg/l	Nitrate range
Archaean granite gneiss	25	24	mainly N-HCO ₃ with minor N-M N-Cl and N-NO ₃	88 454 869	very variable soft to very hard	20 to 30	3 to 184 NO ₃ generally exceeds 45 but but not 102
Granite-Bothaville	1	25	M - HCO ₃	430	hard	17	52
Bothaville	3	25	Variable M-HCO ₃ M-M and M-NO ₃	100 to 190	soft to moderately hard	20 to 26	less than 12, except for JL 2
Allanridge Andesite- Bothaville	3	25	M-HCO ₃	330 to 380	hard	15 to 20	below 36
Allanridge Andesite	18	26	mainly N-HCO ₃ with rare N-M N-Cl Ca-HCO ₃ and Ca-NO ₃	298 600 1 355	mainly hard but ranges from moderately hard to very hard	17 to 432 13 to 21	NO ₃ generally exceeds 45 but not 102
Karoo Sequence covering Allanridge Andesite	10	27	mainly N-HCO ₃ and single Na-HCO ₃	556 787 1 358	hard to very hard	generally range from 13 to 21	25 to 79 50% of samples below 45
Surface water - Harts River	2	25	Na-HCO ₃ and Na-M	230 to 490	soft to moderately hard	below 3	below 3

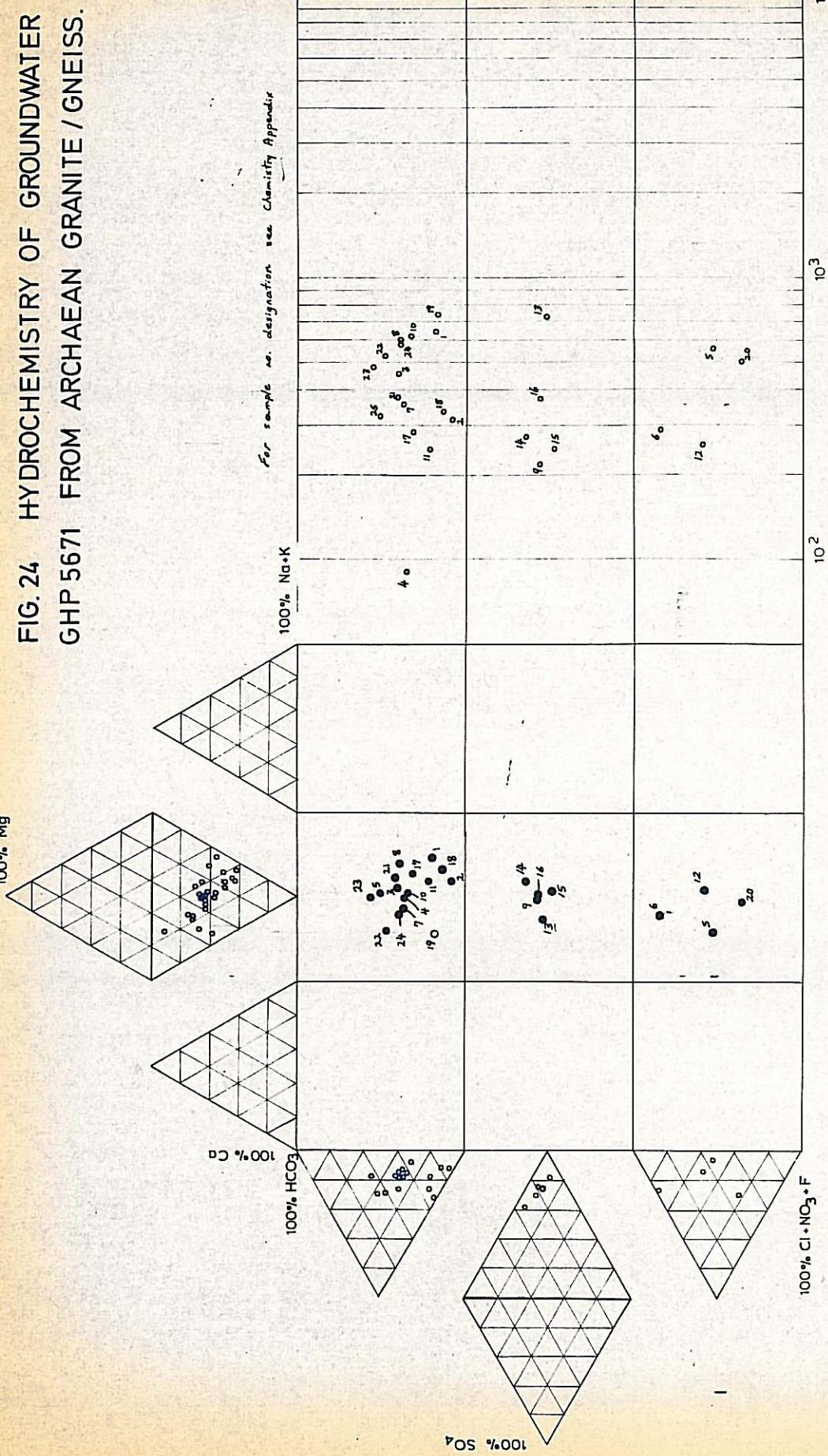
* Classification of hardness in mg/l CaCO_3
After Sawyer and Mc Cartt 1978

0 - 75 soft
75 - 150 moderately hard
150 - 300 hard
300 + very hard

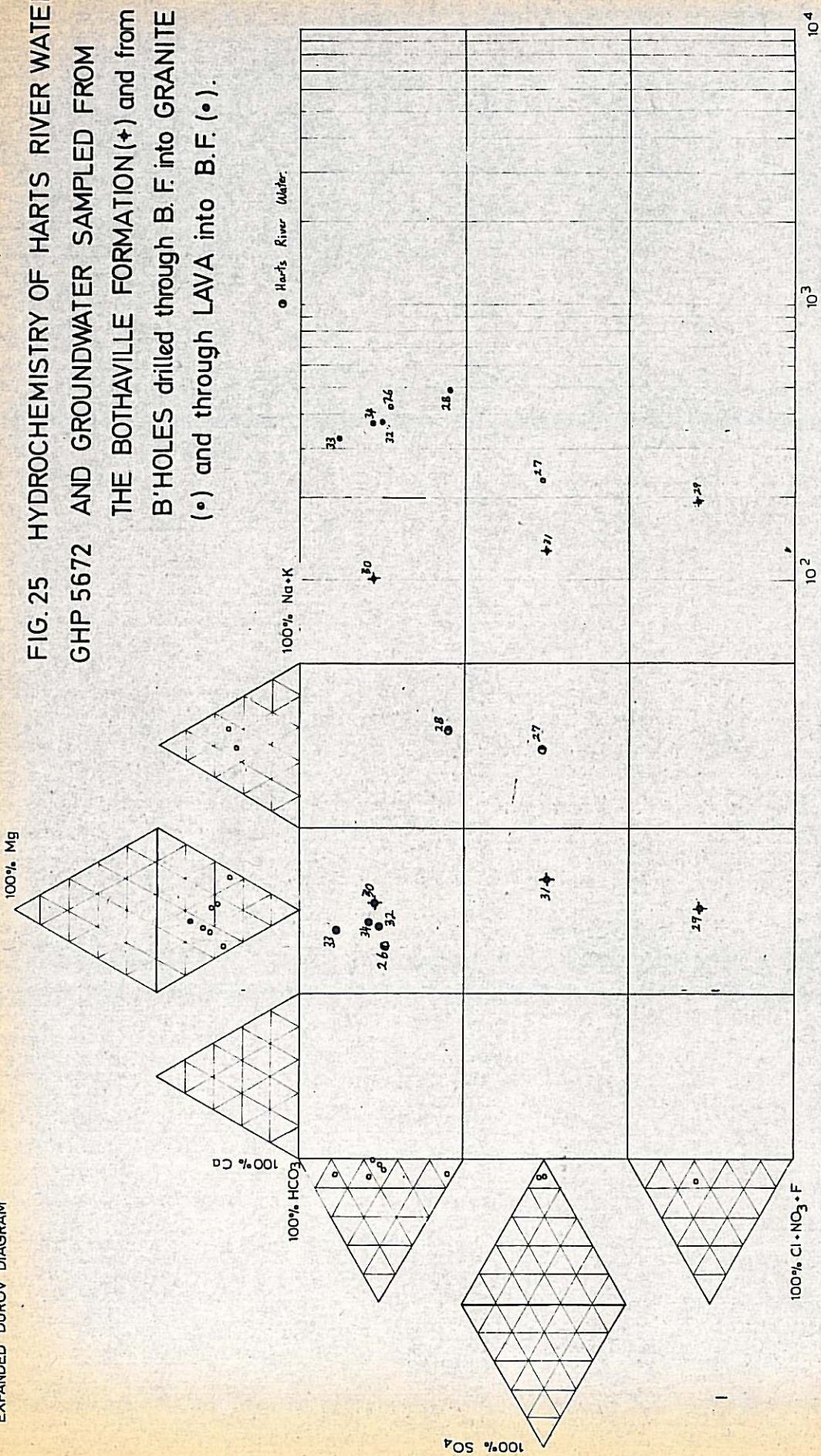
EXPANDED DUROV DIAGRAM

100% Mg

FIG. 24 HYDROCHEMISTRY OF GROUNDWATER
GHP 5671 FROM ARCHAean GRANITE / GNEISS.

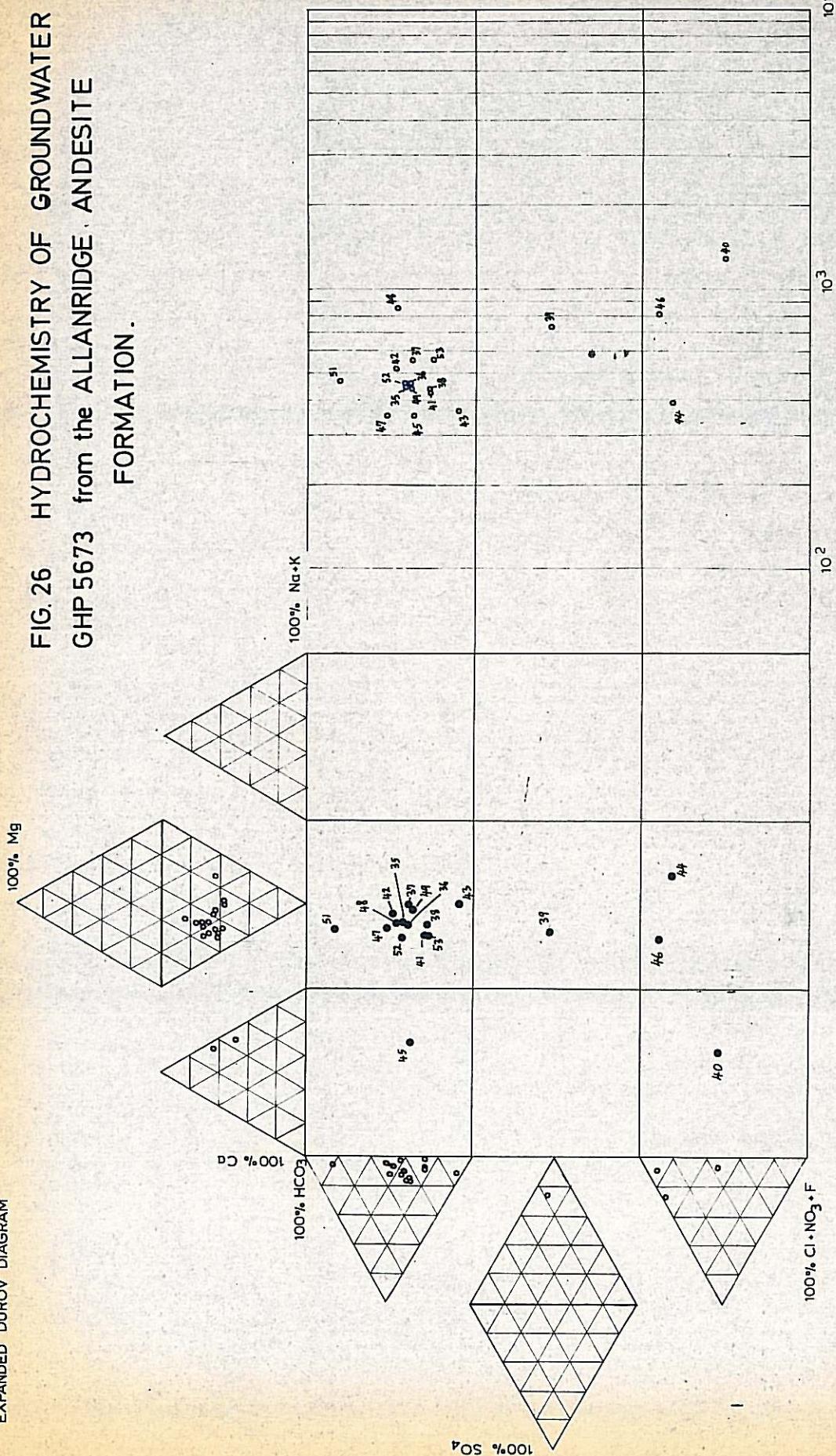


EXPANDED DUROV DIAGRAM



EXPANDED DUROV DIAGRAM

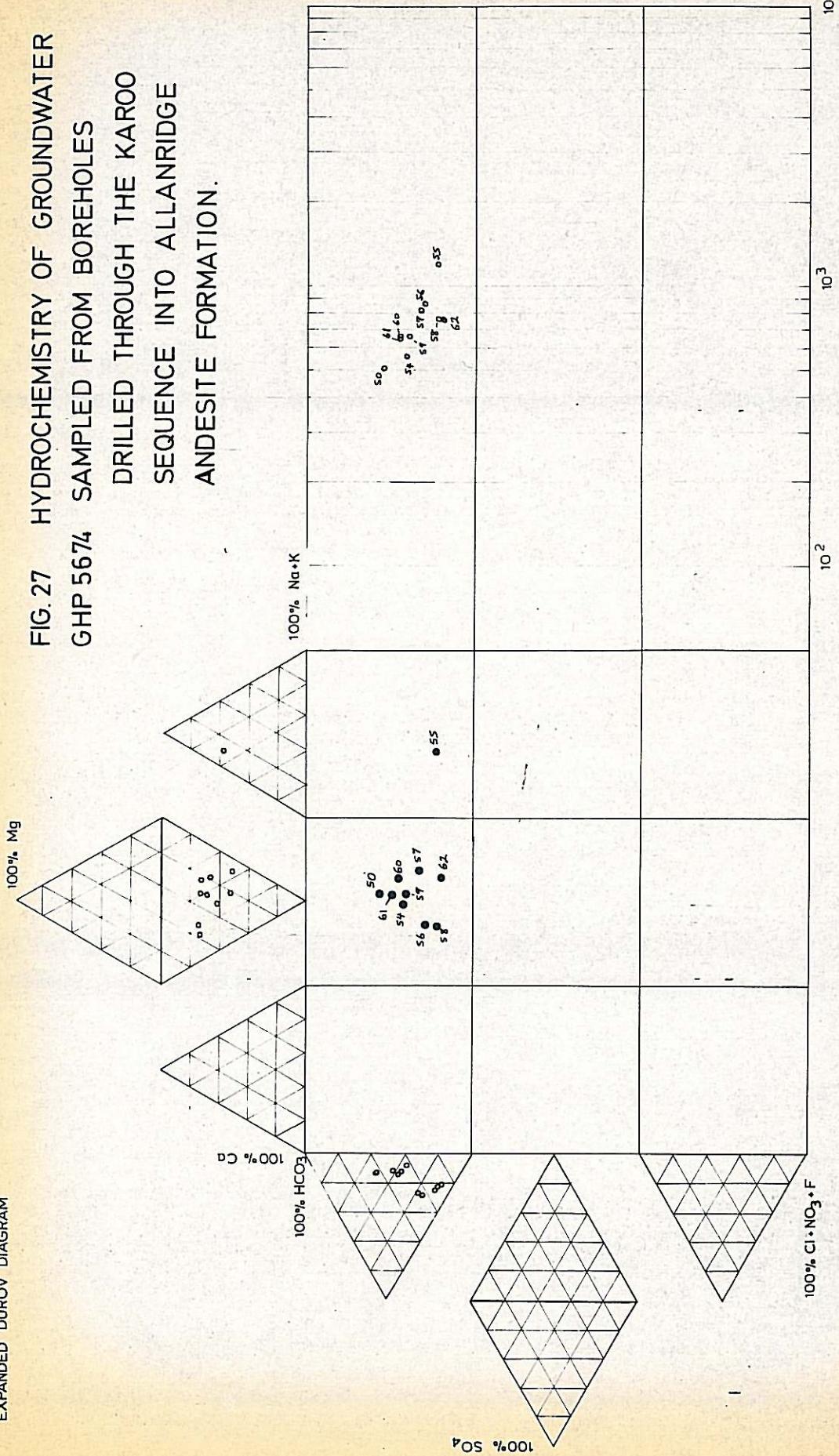
FIG. 26 HYDROCHEMISTRY OF GROUNDWATER
GHP 5673 from the ALLANRIDGE ANDESITE
FORMATION.



TDS. (from field E.C.) in ppm. at 25°C.

EXPANDED DUROV DIAGRAM

FIG. 27 HYDROCHEMISTRY OF GROUNDWATER
GHP 5674 SAMPLED FROM BOREHOLES
DRILLED THROUGH THE KAROO
SEQUENCE INTO ALLANRIDGE
ANDESITE FORMATION.



7.2.1 Total dissolved solids

In all cases T.D.S. values are below 2 000 mg/l, the maximum permissible limit set by the S.A.B.S. Most ground waters are classified as fresh with T.D.S. less than 1 000 mg/l, but some more mineralised brackish waters occur within the lava, with or without the presence of a Karoo cover. Waters from the arenaceous rocks are only very weakly mineralised.

The Harts River was sampled in November 1981 after light rainfall. Surface water upstream of Wentzel Dam had T.D.S. values of 700 to 800 mg/l, i.e. only slightly higher than local ground water. T.D.S. values downstream of Wentzel Dam were considerably lower i.e. 230 to 490 mg/l. These values are slightly lower than the T.D.S. content generally found in the surrounding granite and thus probably represents ground water diluted by recent surface water run-off.

7.2.2 Total hardness

The recommended upper limit of the S.A.B.S. is 200 mg/l, while the maximum permissible limit is 1 000 mg/l. Reference to Appendix 3 and Table 1 indicates that for many waters the total hardness exceeds 200 mg/l, however in all cases it is less than 1 000 mg/l. Waters which are weakly mineralised can be classified as soft to moderately hard while more heavily mineralised waters are hard to very hard.

7.2.3 Nitrate concentrations

Groundwater nitrate concentrations are high throughout most of this area. Previous recorded chemical analyses suggest that this is not a new development (de Villiers 1961).

Sixty-two percent of the sampled ground waters have nitrate levels in excess of 45 mg/l, the maximum limit for drinking water set by the S.A.B.S. If this standard is strictly adhered to much of the ground water currently being consumed in this area by the rural population would be classified as unpotable. However if a maximum concentration

of 102 mg/l (23 mg/l as N) reported by Kempster 1980 is considered acceptable for human consumption 95% of the waters sampled would be potable. To date no harmful effects (e.g. methemoglobinemia) caused by the consumption of high nitrate water have been reported from the Schweizer-Reneke district.

Nitrate, oxidised nitrogen and nitrite, can enter and pollute ground water by a number of mechanisms (Hern 1970), viz:

- leaching of nitrogeneous fertilisers applied to arable land,
- leaching of dung from livestock congregations,
- leaching from domestic sewage disposal systems and
- decomposition of nitrogen fixing leguminous plants.

All these mechanisms may be operating concurrently in the area studied. Once nitrate enters ground water it is highly mobile and travels with the local hydraulic gradient.

Actual nitrate levels vary considerably from less than 3 mg/l to an exceptional high of 432 mg/l reported for borehole Lt 91.

The areal distribution of nitrate concentrations are shown in Enclosure 2. This map indicates that nitrate levels are fairly uniform even over relatively large distances and down gradient of the ground water flow paths. Ground waters of low nitrate concentration are found on:

- the north-eastern portion of the Schweizer-Reneke Townlands 62 HO,
- Rietput 60 HO,
- east of the Maraetchesfontein Dyke in a north-south belt extending from Maraetchesfontein 54 HO to Nooitgedacht 86 HO, and
- along parts of the Harts River valley.

Low nitrate concentrations often appear to be related to a low level of farming activity where little or no fertiliser is applied to the land. This is particularly evident over the ground water recharge areas where quartzite and lava outcrop e.g. borehole Pe 1 with a NO_3 level of 9 mg/l. Exceptions to this generalisation are probably related to concentrations of livestock collecting around watering points e.g. J1 2, NO_3 level of 63 mg/l. Relatively low nitrate levels on Nooitgedacht 86 HO and part of Maraetchesfontein 54 HO may be due to the presence of a thick mudstone cover locally retarding the downward migration of nitrate ions to the ground water table.

In 1981 nitrate levels in the municipal wells at Schweizer-Reneke were low, borehole Ss 2 - 11,3 mg/l, this is presumably due to dilution effects caused by recharge from the low nitrate waters present in the Harts River.

During the initial period of well usage in 1980, nitrate concentrations were high (53 to 174 mg/l in Hattingh, 1980) but subsequently decreased with continuous pumpage. Part of the high nitrate may have been attributed to pollution from the sewage disposal method practised in the adjoining township.

Figure 27a shows nitrate concentration of ground water intercepted in documented boreholes and suggests that nitrate concentrations are high throughout the aquifers penetrated by the boreholes. However this cannot be confirmed as no progressive sampling of ground water has been undertaken during drilling.

7.3 Quality of ground water available from potential production boreholes

G34307 - Maraetchesfontein 54 HO (Analyses No. 49 and 67 to 71 in Appendix 3). This borehole was sampled when it was drilled and during subsequent pump testing when well head electrical conductivity (E.C.), temperature and pH were continuously monitored. Apart from the low values reflected by sample No. 71 values showed little variation. The T.D.S. content at 25°C was 470 mg/l (\pm 6%), well head temperature was 21°C with a pH of 7,2 (see Appendix 4 for further details). Similarly, apart from sample No. 71 with a lower T.H., the major ions showed little variation in concentration.

Ground water drained from the surrounding lavas towards G34307, in the Maraetchesfontein Dyke, are similar in composition to those waters sampled at G34307. However ground water to the east, under a Karoo mudstone cover, have substantially higher T.D.S. values at borehole Mn 2 as well as higher fluoride and lower nitrate concentrations.

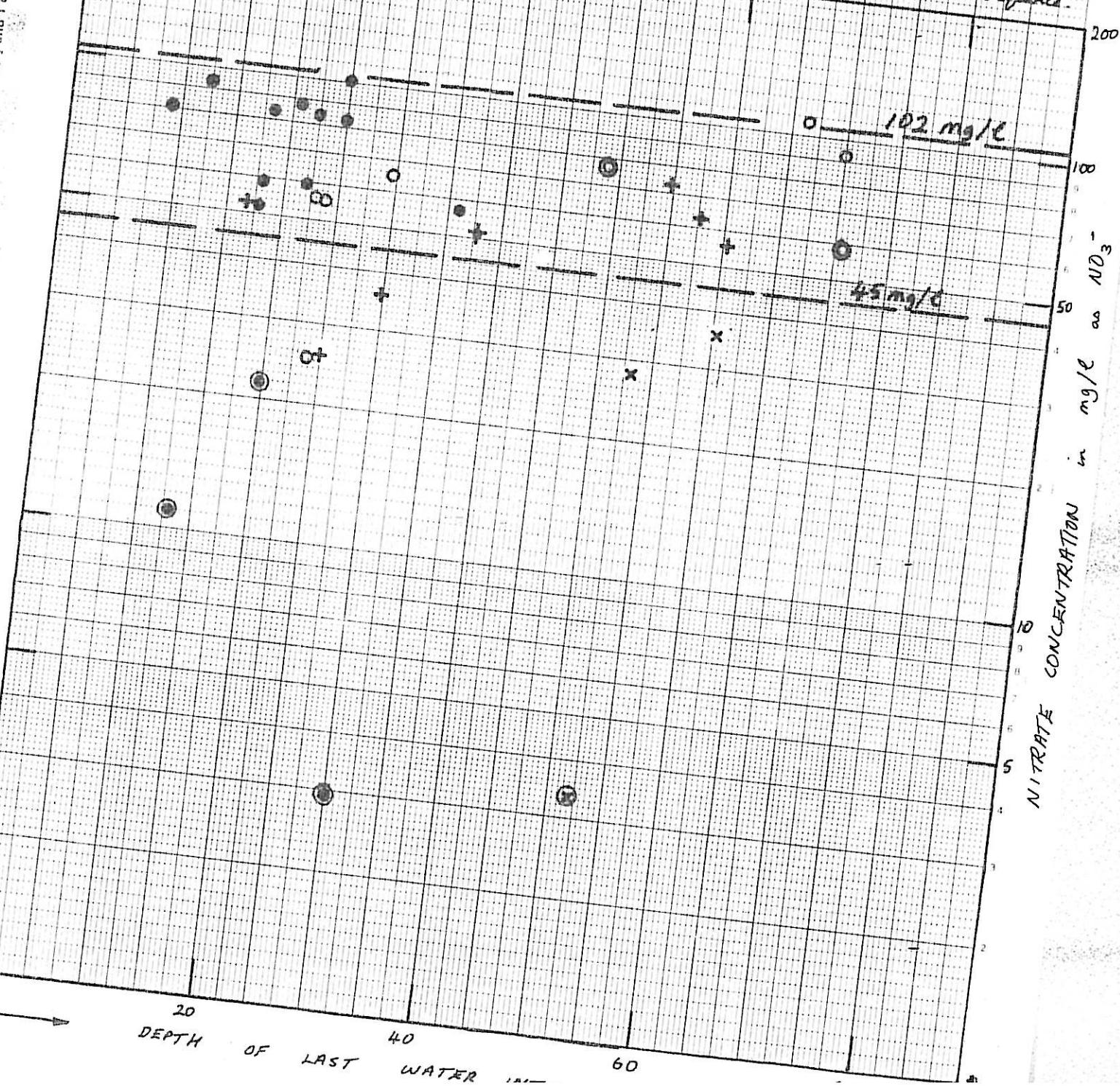
It can be expected that under long-term pumping ground water abstracted at site G34307 will not deteriorate substantially in quality and will continue to provide hard water with a low fluoride content and a nitrate concentration slightly in excess of the current S.A.B.S. upper limit of 45 mg/l.

FIG. 27a PLOT OF NITRATE CONCENTRATION AGAINST
GHP 5670 DEPTH OF LAST WATER INTERCEPT.

TOTAL NO. OF RESULTS USED = 32.

- Sampled from Archean Granite (ringed - Boreholes probably recharged by Hart River)
- ★ B'hole penetrates Bathurst - Granite contact
- ✖ B'holes penetrating Andesite - Bathurst contact (ringed - probably diluted by recharge from Hart River).
- Sampled from Allardige Andesite Formation (ringed - samples in or close to Marachefontein Dyke)
- + Sampled from Allardige Andesite Formation overlain by Kara sequence.

..... 1 and 1 cm



G34310 - Hamburg 82 HO (analyses No. 9, 63 and 64) - Ground water was sampled during drilling and later pumping tests, while E.C., temperature and pH were regularly monitored. Field pH varied from 6,9 to 7,15 with a T.D.S. content at 25°C of 220 mg/l (\pm 7%) and a water temperature of 19°C.

Chemical analyses indicated that during pumping the percentage concentration of sodium and nitrate slowly increased. Nearby water samples are more highly mineralised and have higher T.H. and nitrate ion concentrations. Therefore it is to be expected that with extended pumpage similar rises in T.D.S., TH and nitrate ion concentration, will be expected at G34310. However this will depend to some extent on the surface water recharge contribution made by the low T.D.S. and nitrate (?) waters periodically flowing in the Hamburg stream. Despite slight changes in water chemistry the pumped water should be moderately hard with only nitrate levels exceeding S.A.B.S. limits.

Pa 1 - Palachoma 64 HO (Analyses No. 15, 65 and 66) - Three samples for analysis were taken during a 24-hour pump test. E.C. and temperature measurements were made regularly. An average T.D.S. content of 236 mg/l (\pm 2%) was calculated, well-head temperature was 19°C. Major ion composition remained fairly constant and showed only a consistent increase in T.H. (from 102 to 111 mg/l). It is expected that this trend will continue under pumping in the short term. Nitrate concentrations are expected to remain fairly constant at approximately 60 mg/l. This water can be classed as moderately hard and apart from slightly high nitrate levels is potable.

G34315 - Schweizer-Reneke Townlands 62 HO (Analysis No. 33) - During trial pumping this borehole was regularly monitored and indicated a T.D.S. of 330 mg/l, pH of 6,8 and a well-head temperature of 21°C. The water is hard and eminently potable with all chemical parameter tested falling within the limits set by the S.A.B.S. Potential ground water recharge waters have higher nitrate concentrations thus it might be expected that nitrate levels in this borehole might increase under long-term pumping. However this will depend on a

number of factors including the relative ground water contribution made by the lava and the arenites and possible ground water recharge from the low nitrate waters in the Wentzel Dam.

Ss 20 (G32303) - Schweizer-Reneke Townlands 62 HO (Analysis No. 24) - This borehole is adjacent to the Wentzel Dam and has been pumped to supply water for the nearby caravan park. Major ion composition indicates that this water is potable with a low nitrate concentration. However when pumped a strong odour is emitted from the water thus bacteriological analysis is necessary before this water can be considered for domestic consumption. Borehole Ss 20 is partly recharged from water leaking through the flow of the dam, therefore periods of low surface recharge with the consequent concentration of chemical constituents in the dam water may render water pumped from this borehole untreatable, and therefore unpotable, at certain periods.

Pa 3 - Palachooma 64 HO (Analysis No. 16) - This borehole is heavily pumped to provide water for irrigation. Apart from high nitrate concentrations, approximately 90 mg/l, this moderately mineralised, moderately hard water is potable.

8. PUMPING TESTS

8.1 Introduction

As an integral part of this investigation test pumping was undertaken at four potential production boreholes namely G34307, G34310, G34315 and Pa 1. Constant rate tests were of two types; short duration tests (less than 1 000 mins.) and medium duration tests (1 000 to 10 000 mins.). In addition a 6-stage step-drawdown test was completed on borehole G34310 while a similar test on borehole G34315 had to be abandoned.

The objectives of the pumping tests were:

- to ascertain the production status of the tested boreholes, and when possible,
- to estimate transmissivity and storage parameters of the penetrated water-bearing formation,
- to calibrate the yielding capacity of the potential production well G34310,
- to monitor changes, if any, of ground water quality during pumping (refer to section 7.3), and
- to ascertain the influence of the geology on the movement of ground water.

All raw-pumping test data is presented in Appendix 4.

8.2 Archaean granite gneiss aquifer

8.2.1 Hamburg 82 HO

General - As borehole G34310 (Fig. 6) recorded a blow yield of some 10 l/s from weathered and fractured granite two observation boreholes, G34311 (Fig. 7) and G34312 (Fig. 7) were drilled 30 m to the south. Other boreholes utilised as piezometers during the tests were G34313 (Fig. 8), Hg 1 and Hg 2. Borehole locations are depicted in Fig. 20.

Prior to pump testing water levels were stable and thus no drawdown correction was necessary. The local ground water movement was north-west towards the Harts River with a hydraulic gradient of 6×10^{-3} . Ground water abstracted during pumping was fed via 90 m of 100 mm diameter outlet piping onto standing water in the nearby stream channel, from where it flowed downstream. Some return circulation of water may have taken place but this is considered to

be insignificant as the actual stream bed is composed of thick black clay. During the constant rate test thunderstorms occurred and some 20 mm of rainfall was recorded.

Step-drawdown test on G34310 - This test consisted of 6 pumping stages of 100 min. duration each under discharge rates ranging from 4,29 to 16,84 l/s. The pumping test data was plotted on single logarithmic paper (Fig. 28) and interpreted using the Bierschenk method (1964). Results are presented in Table 12 and Fig. 29. The formation loss coefficient, B, is estimated to be $0,66 \times 10^{-2} \text{ d/m}^2$ and the well loss coefficient, C, is calculated as $3,4 \times 10^{-6} \text{ d/m}^5$. However these figures are approximate only as poor pump performance caused fluctuating discharges. Additionally the water level recovery noted during step 6 was the result of the well screen being partly blocked by coarse granitic sand. The specific discharge-drawdown plot in Fig. 9 indicates that the above calculated B and C values are only valid when water levels are above 15 m. Fig. 28 indicates that pumpage at a rate slightly in excess of 13,3 l/s for a period of 100 mins. lowers the water level to the upper limit of the well screen (set between 15 to 18 m). For a discharge rate of 13,3 l/s (at $t = 100 \text{ min.}$) borehole efficiency is only 70%, the drawdown in the pumped well is 12,1 m of which 7,6 m is the result of formation losses and the remainder 4,5 m, is caused by well losses.

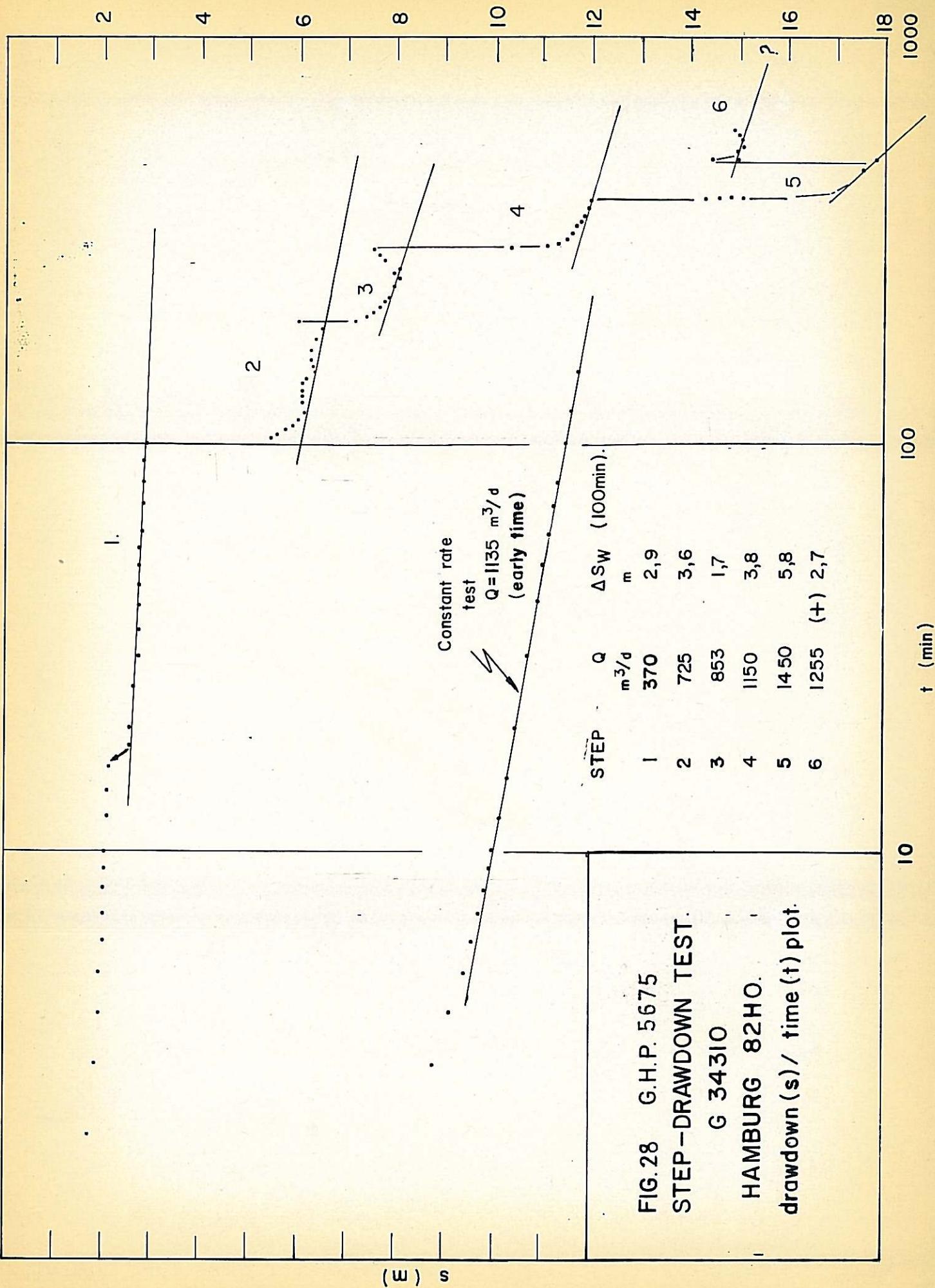
TABLE 12: WELL CALIBRATION G34310

STEP	Q 1/s	S _w (t=100 min) (m)	Cumulative S _w (m)	S _w /Q (X 10 ⁻²)	BQ	CQ ²	BQ+CQ ²	BQ BQ+CQ ²	X 100%	Specific capacity m ³ /d/m drawdown for t=100 min
1	4,29	2,90	2,90	0,784	2,44	0,47	2,91	84	127	
2	8,40	3,60	6,50	0,897	4,79	1,79	6,58	73	111	
3	9,87	1,70	8,20	0,961	5,63	2,47	8,10	70	104	
4	13,31	3,80	12,00	10,43	7,59	4,50	12,09	63	96	
5	16,84	5,80	17,80	12,23	(9,60)	(7,20)	(16,80)	(57)	82	
6	14,54 (-) 2,70	15,10	12,00	(8,28)	(5,36)	(13,64)	(61)	83		
Constant rate test 13,14	11,44			[10,98]	[7,49]	[4,38]	[11,87]	[63]	99	

Where B = 0,66 X 10⁻² days/m² and C = 3,4 X 10⁻⁶ days/m⁵

S_w/q = Specific drawdown, BQ = formation loss, CQ² = well loss and $\frac{BQ}{BQ+CQ^2} \times 100\% =$ well efficiency

[] Interpolated values () theoretical values



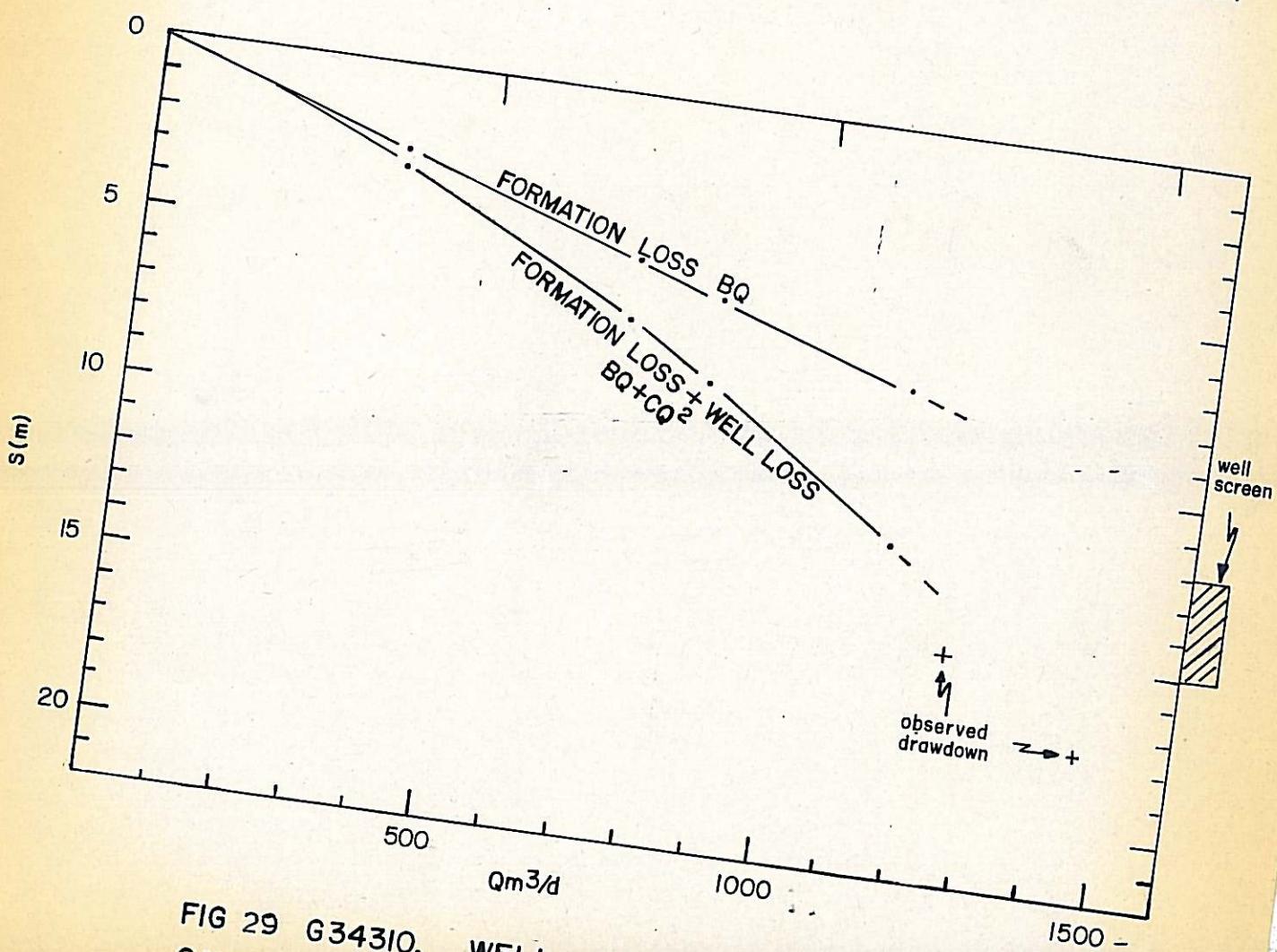
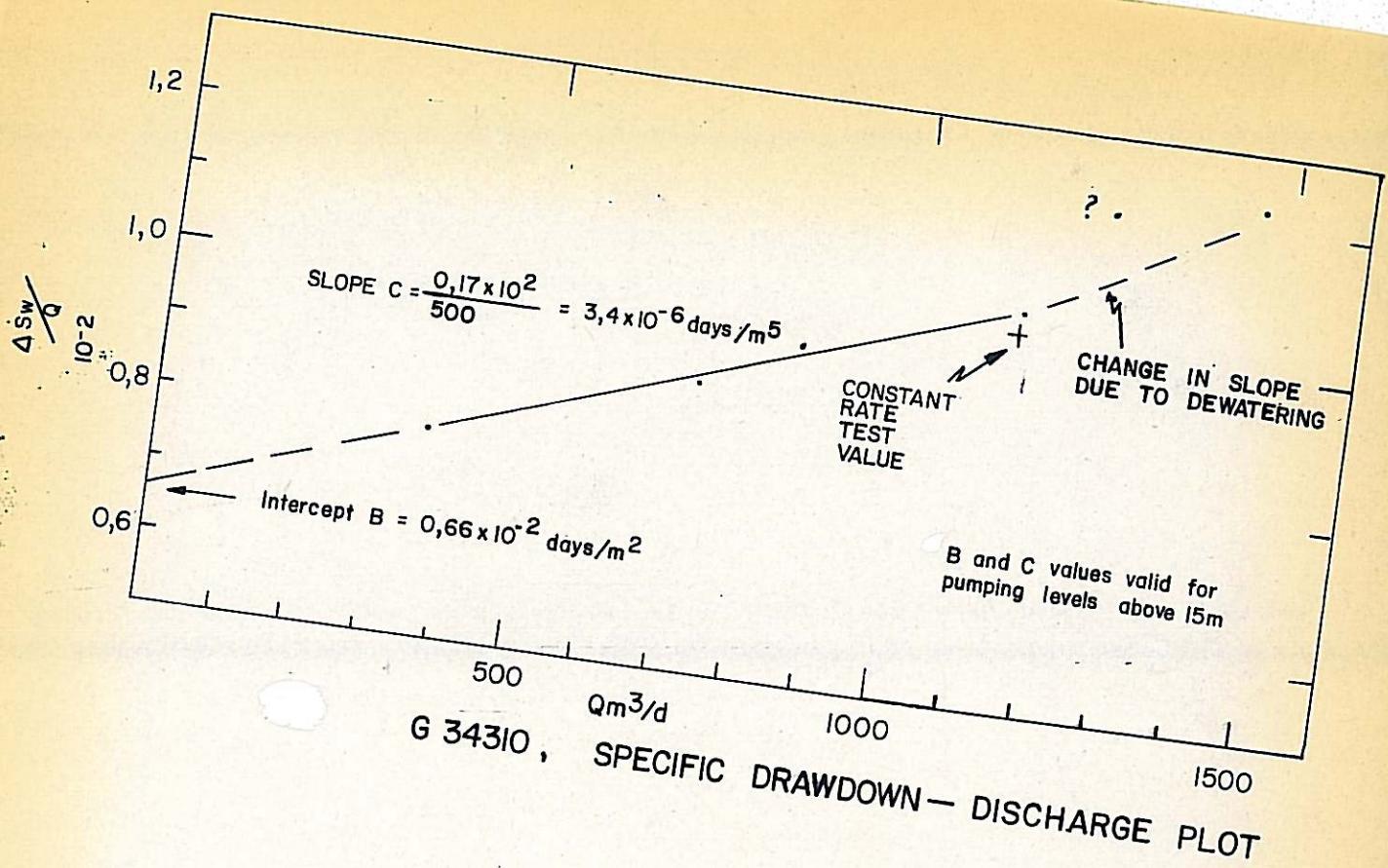


FIG 29 G 34310, WELL EFFICIENCY DIAGRAM
G.H.P. 5676

Constant yield test at borehole G34310 - The pumping period was 7 135 min. with an average discharge rate of approximately 12,65 l/s (13,14 declining to 12,2 l/s and subsequently to 10,91 l/s). Coarse granitic material partially blocking the pump screen caused this decline in pumping rate and recovery of water levels in G34310 after 5 000 min. After the cessation of pumping the recovery of the water levels was monitored for some 6 000 min.

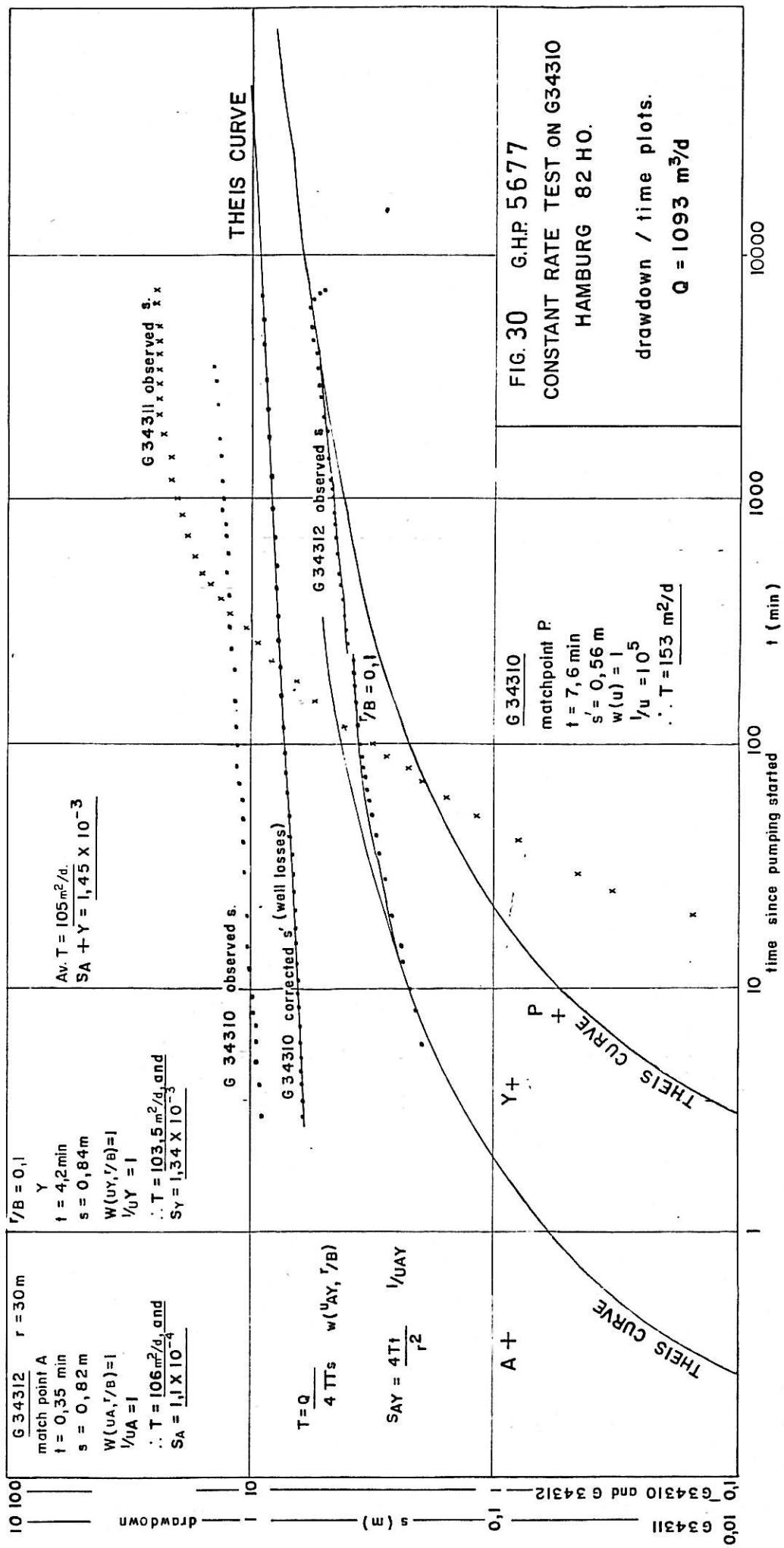
The drawdown (s) - time (t) curves have been plotted on double logarithmic paper and are presented in Fig. 30. Drawdown measurements indicate that the ground water flow to the borehole was at all times in an unsteady state. The appreciable drawdown observed in the covering layer of granitic sand (decomposed granite) by G34311 indicates that the aquifer type is semi-unconfined or unconfined.

The t-s curve for observation borehole G34312 was initially matched to the Theis curve for unconfined aquifers under unsteady conditions but produced a very low storage value (1×10^{-5}) indicative of confined conditions. Consequently the t-s (observed) curve was fitted to the Boulton type curves for unsteady state flow in semi-unconfined aquifers with delayed yield (refer to Fig. 30 for co-ordinates) and the following hydraulic properties were calculated:

average transmissivity (T)	$105 \text{ m}^2/\text{d}$
effective specific yield (S)	$1,45 \times 10^{-3}$

The t - s' (corrected for well losses) curve for the pumped borehole, G34310, was fitted to the Theis curve and a T value of $153 \text{ m}^2/\text{d}$ was calculated. This figure is in fair agreement with that produced by G34312 when it is recalled that more favourable hydraulic conditions prevailing initially at the pumped borehole produced a higher blow yield.

No quantitative analyses could be carried out on the drawdown measurements made at boreholes G34311, G34313, Hg 1 and Hg 2.



The recovery of the ground water levels after pumping ceased are shown in Fig. 31. Recovery was fairly rapid with water levels returning to within 0,5 m of original rest water levels some 5 800 min. after pumping stopped.

Results - Pumping test results indicate that aquifer conditions within the weathered and fractured granite aquifer are semi-unconfined with gravity drainage taking place from the overlying decomposed granite. Acceptable average hydraulic parameters are

$$T \quad 120 \text{ m}^2/\text{d}$$

$$S \quad 1,5 \times 10^{-3}.$$

The favourable hydraulic properties of the aquifer, the fast recovery of water levels and the relatively high specific capacities (see Table 12) indicate that conditions are suitable for the establishment of a number of high yielding production boreholes. Borehole G34310 requires redesigning if it is to be utilised effectively as a production borehole capable of yielding an estimated $1\ 000 \text{ m}^3/\text{d}$ (11,6 l/s). The structural changes suggested to the well include increasing the final well diameter to 203 mm and installing a well screen from 18 to 22 m with a 10% open area and a suitably graded gravel pack.

8.2.2 Palachoema 64 HO

General - Boreholes Pa 1 and Pa 2 were drilled by Vryburg Boring Inspectorate in early 1981, borehole logs and site positions are shown in Fig. 32. Electrical resistivity depth soundings suggest that the main ground water supply in the pumped borehole, Pa 1, was struck in weathered granite above 24 m. Aquifer conditions were thought to be confined or semi-confined with overlying clayey decomposed granite (see Fig. 32) acting as an impervious or semi-pervious layer.

Regional ground water flow is to the south towards the Harts River with a gradient of 6×10^{-3} . Water levels monitored prior to testing were stable (change 5 mm in 4 days) and no rain fell during the pumping period.

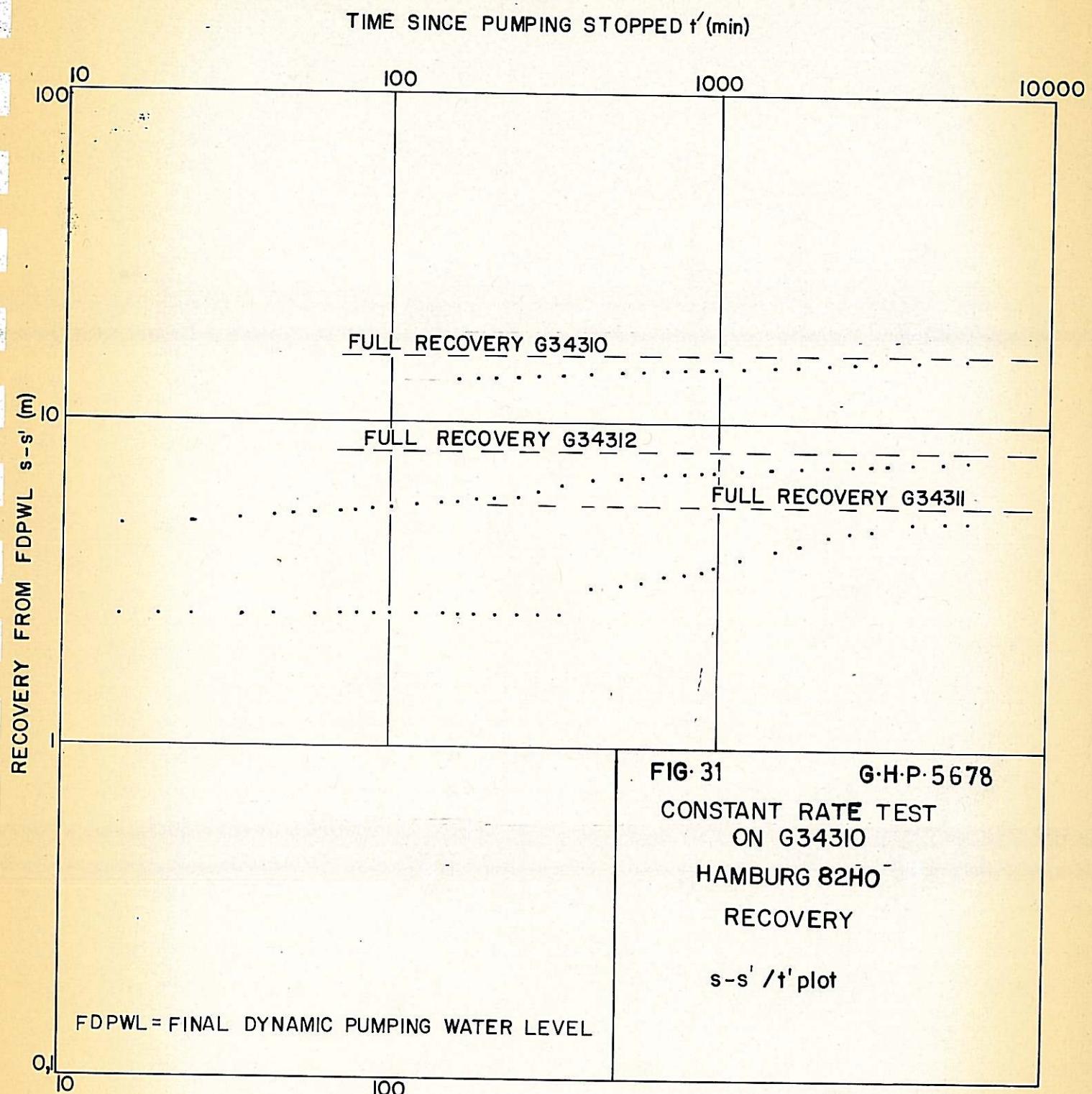


FIG. 32 G.H.P. 5679

CONSTANT RATE TEST

ON Pa 1

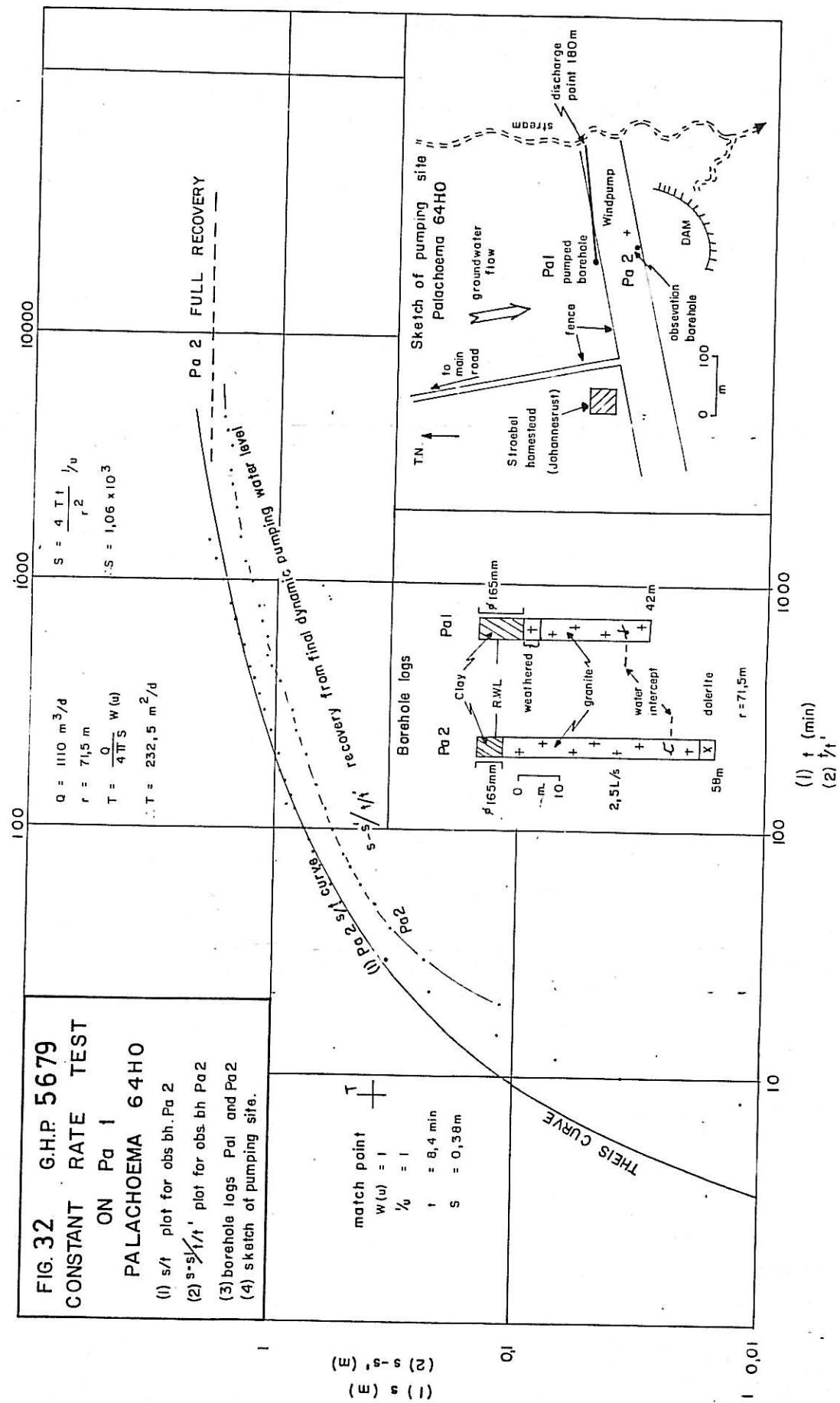
PALACHOEMA 64HO

(1) s/t plot for obs. bh. Pa 2

(2) s^2/t^2 plot for obs. bh. Pa 2

(3) borehole logs Pa 1 and Pa 2

(4) sketch of pumping site.



Constant yield test at Pa 1 - Borehole Pa 1 was pumped for 1 440 min. at an average discharge rate of 12,69 l/s (measured rate 12,87 to 12,64 l/s). As no water level measurements could be made in Pa 1 during pumping only drawdown data from observation borehole Pa 2 could be used for analysis. The t-s curve was matched to the Theis curve for unsteady state flow in confined aquifers (see Fig. 32 and Table 13) and the following hydraulic properties were estimated:

$$T = 230 \text{ m}^2/\text{d}$$

$$S = 1,06 \times 10^{-3}$$

The observed t-s curve showed deviation from the Theis curve after 800 min. The S-value calculated is in the range, albeit only marginally normally associated with unconfined aquifer conditions, therefore some doubt must be placed on the initially suspected aquifer conditions.

Boreholes Pa 1 and Pa 2 showed a fairly rapid recovery, (see Fig. 32) after 4 500 min. water levels had risen to within 0,3 m of pre-pumping levels.

Results - The weathered and fractured probably semi-unconfined aquifer has favourable hydraulic properties for the establishment of production wells.

Borehole Pa 1 was shown to be capable of producing yields in excess of $1\ 000 \text{ m}^3/\text{d}$ (11,6 l/s). However the actual pumping rate recommended would depend on the results of a step-drawdown test which was not undertaken because of inaccessibility into the borehole.

8.3 Borehole G34315 penetrating the lava-quartzite contact, Schweizer-Reneke Townlands 62 HO

General - Borehole G34315 recorded a blow yield of 4,3 l/s from 3 water interceptions in fractured lava and quartzite (refer to Fig. 10). This borehole was therefore pumped to ascertain the borehole yield. A step-drawdown test was attempted but failed as the

101 mm turbine pump available could not deliver water at the required low rate. Additionally water level measurements could only be made to some 17 m below the casing collar. No piezometers were available for drawdown measurements.

Before pumping water levels were stable. During both recovery phases light rain was recorded (approximately 5 mm). Groundwater movement is in a south-south-east direction with a relatively steep gradient of 1×10^{-2} . Abstracted ground water was discharged in a nearby stream (refer to Fig. 21 for location plan.)

Phase 1 - The pumping period was 180 min. with an average discharge of 5,78 l/s (6,02 l/s declining to 5,54 l/s). Drawdown measurements indicated that the pumping water level had dropped below the topmost water interception (17 m). For the first 50 min. the abstracted water was discoloured brown-red by suspended clay and silt. Recovery was fairly rapid after 205 min. water levels were within 0,27 m of the pre-pumping level. The Theis recovery method was applied and a transmissivity value of $70 \text{ m}^2/\text{d}$ (see Fig. 33) was estimated.

Phase 2 - After 205 min. of recovery the borehole was tested again for 150 min. and the discharge rate declined from 7,13 l/s to 5,29 l/s. The recovery was monitored for 840 min. by which time water levels had returned to within 0,18 m of the pre-phase 1 water level (see Fig. 33).

Results - The two short tests indicate that the highest-sustainable pumping rate of G34315 is probably of the order of 4 to 5 l/s. At this rate water levels are below the upper water strike in the lava.

Recovery from pumping is fairly rapid. A transmissivity value of $70 \text{ m}^2/\text{d}$ was calculated but the value of this is not apparent in this complex, fissure flow, multiple aquifer.

8.4 Borehole G34307 penetrating the dyke system on Maraetchesfontein 54 HO

General - Borehole G34307 (Fig. 3) produced a blow yield in excess of 10 l/s from fractured formation associated with the Maraetchesfontein

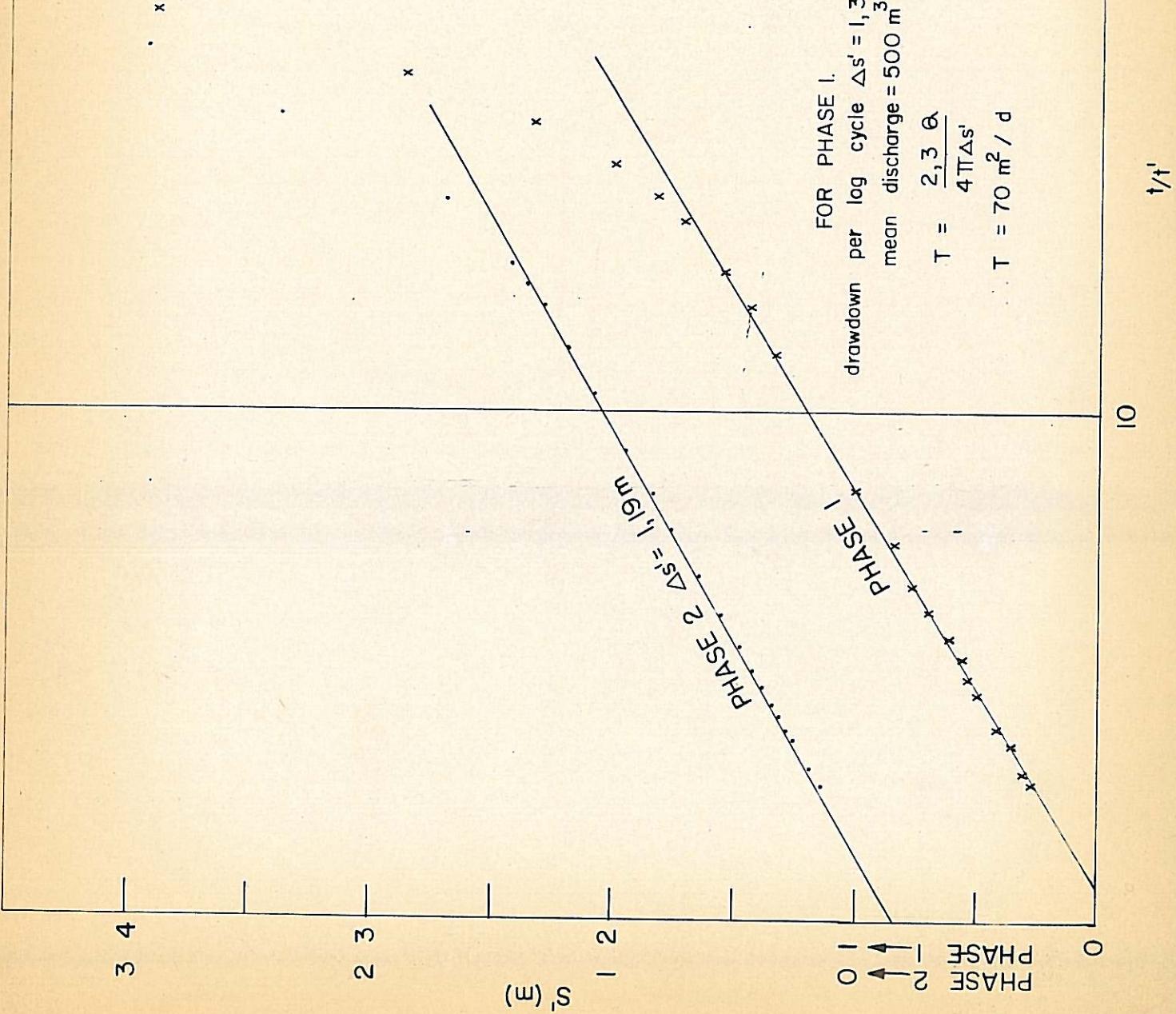


FIG. 33 G.H.P. 5680
THEIS RECOVERY ANALYSIS
G34315
SCHWEIZER-RENEKE TOWN-
LANDS 62HO
 $s' \sqrt{t/t'}$

Dyke. The additional boreholes G34308 (Fig. 9), G34309 (Fig. 10), G34318 (Fig. 13) and G34319 (Fig. 13) were later drilled (see Figs. 22 and 35) and were monitored during the 4-phase pump testing of G34307. (N.B. G34318 and G34319 were only measured during phase 4). Apart from phase 1 water level measurements could not be made in G34307.

Due to the expected conduit nature of the confined or semi-confined fractured aquifer it was expected that the pressure relief surfaces created by pumping would be ellipsoid with the maximum drawdown and hence transmissivity being along the line of the fracture system (see Fig. 35). The data obtained from pump testing was found to be unsuitable for quantitative analysis due to the lack of sufficient purpose-drilled observation boreholes and fluctuating yields from the pumped borehole.

The hydraulic gradient is 3×10^{-3} with ground water flow in a west-north-west direction. The dykes within the lava have apparently no effect on the piezometric surface contour pattern (Enclosure 2). Prior to pumping phases 1 and 4 water levels were stable while prior to phases 2 and 3 water levels were still recovering. Light rainfall 2,5 mm, was recorded during phase 4 while heavy rainfall, some 35 mm, was experienced during the recovery from pumping phase 4.

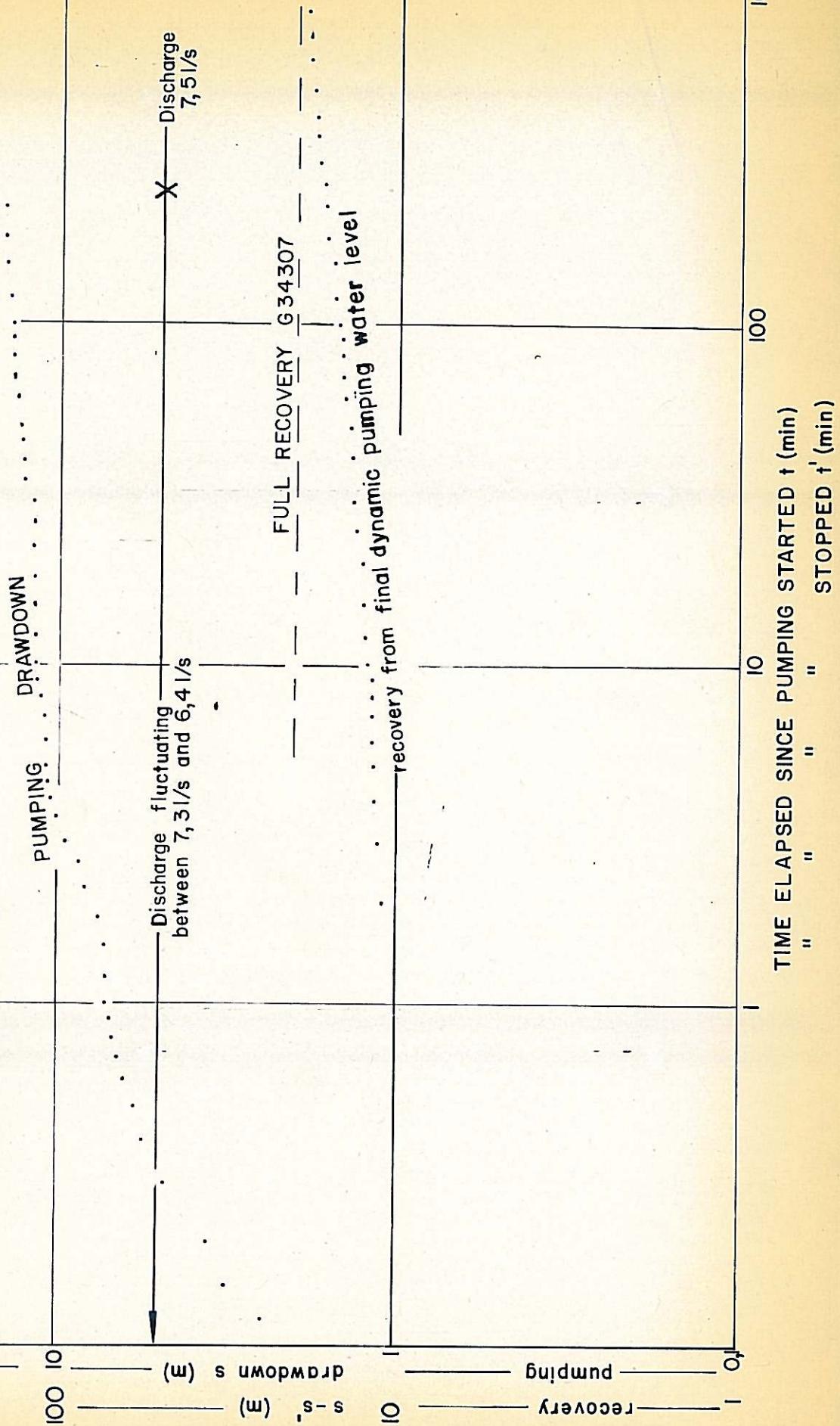
During phases 1 to 3 the 101 mm turbine pump intake was set at 62 m while during phase 4 the replacement pump was installed to 64 m. Four hundred metres of outlet piping conveyed the abstracted water down-gradient of G34308 (see Fig. 35).

Phase 1 - Borehole G34307 was pumped for 480 min. at 7 l/s, a pressure transducer coupled to an analog chart recorder was used to monitor drawdown and recovery of the ground water in the pumped borehole. The irregular time-drawdown plot presented in Fig. 34 is the product of poor pump performance. Observation borehole G34309 responded rapidly to the pumping of G34307 (see Fig. 34) and subsequently recovered fairly rapidly. The poor hydraulic connection between observation borehole G34308 and the pumped borehole is clearly shown by the continuous time-drawdown plot in Fig. 35.

FIG. 34 GHP. 5681
G 34307, PUMPED BOREHOLE
MARAETCHES FONTEIN 54HO

s/t plot, and
 $s-s'/t$ recovery plot

DATE 27.10.1981



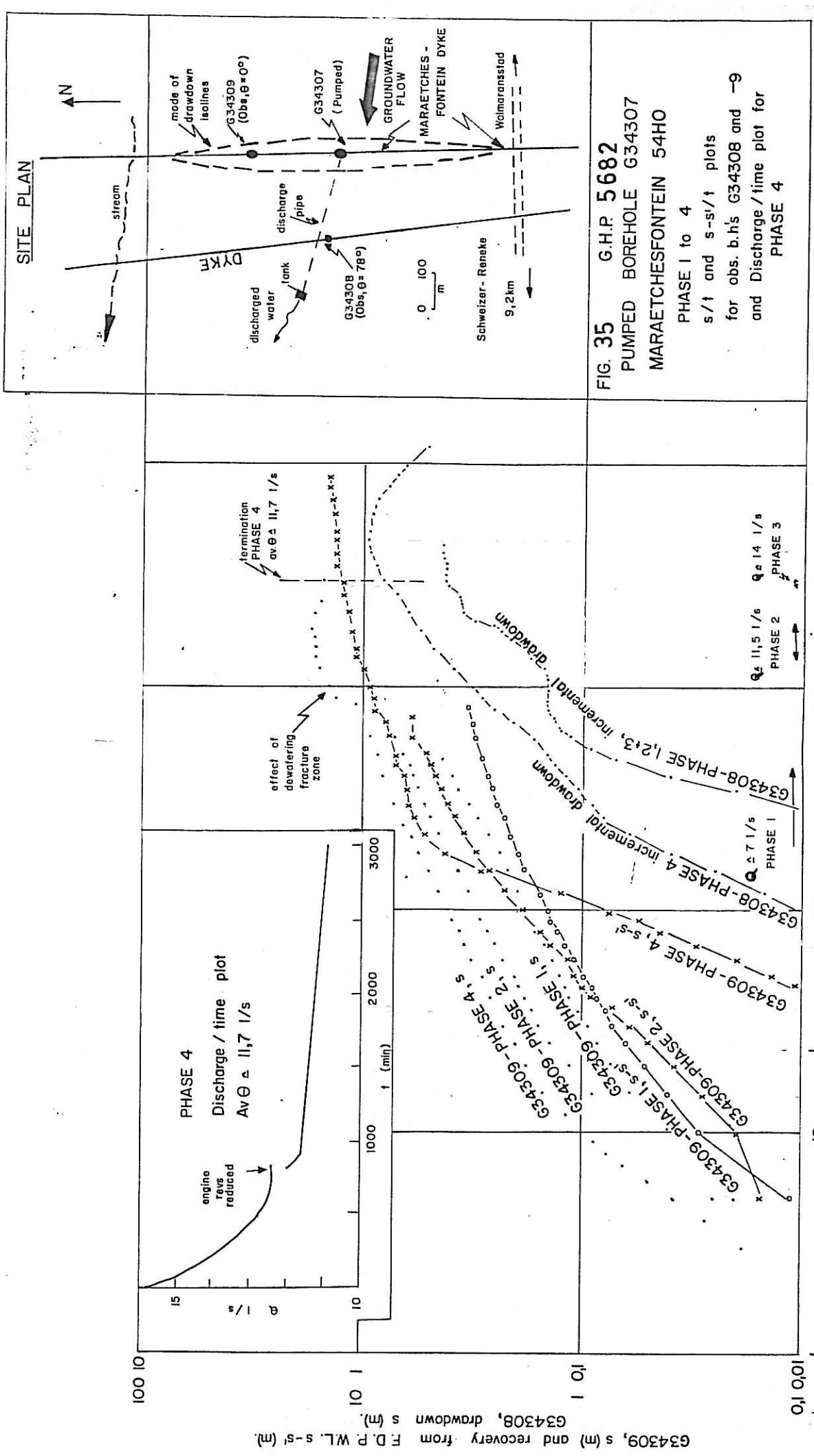


FIG. 35 **G.H.P. 5682**
PUMPED **BOREHOLE G34307**
MARAETCHESFONTEIN 54HO
PHASE 1 to 4
s/t and s'-s'/t plots
for obs. b.h.s G34308 and -9
and Discharge /time plot for
PHASE 4

FIG. 3
PUM
MAR

PHASE I to 4
 s' / t and $s-s'$ / t plots
 obs. b. hs G34308 and -9
 Discharge / time plot for
 PHASE 4

Phase 2 - Prior to pumping, water levels in the boreholes G34307, G34308 and G34309 were 1,07, 0,145 and 0,97 m respectively below pre-phase 1 rest levels. The pumping period was 600 min. with an average discharge rate of 11,5 l/s. Recovery in G34307 and 8 was fairly rapid.

Phase 3 - Before pumping commenced water levels were considerably lower than pre-phase 1. The initial discharge rate was 15,82 l/s which declined to 12,1 l/s after 110 mins. Consequently the pumping test was terminated. At this stage poor pump performance was blamed for the declining yield.

Phase 4 - The pumping period was 3 000 min. at an average discharge rate of 11,7 l/s. Declining yields were experienced (see Fig. 35) and a pump adjustment was necessary after 720 min. to prevent overheating. Time-drawdown curves and recovery data are presented in Fig. 35. The drawdown pattern experienced in G34309 after 800 min. pumping illustrates the effect of dewatering the shallow fracture zone and the consequent change to unconfined conditions. After some 120 min. of pumping observation borehole G34308 responded. No response to pumping was detected in the distant piezometers G34318 and G34319. Water-level recovery in G34307 and G34309 was fair. For G34307 the water level was some 0,63 m below pre-pumping level after 7 600 min. of recovery.

Results - Observations made during the tests indicate:

- the maximum drawdown available in G34307 is 63 m (for a rest water level of 8 m), which reduces to 56 m if the pump intake is set at 64 m,
- a discharge rate greater than $850 \text{ m}^3/\text{d}$ (approximately 10 l/s) is unlikely to be maintained for extended pumping periods. This is because the initial high discharges are provided by the partial dewatering of the fracture zone which is then replenished, at a lower rate, from interconnected fractures within the lava,

- apparently high well losses cause excessive drawdown in the pumped borehole,
- recovery is fairly rapid with over 90% of the drawdown recovered in a time period equivalent to that of the pumped period,
- the water bearing formation shows marked anisotropy with zone of preferential transmissivity associated with fractures along the Maraetchesfontein Dyke,
- although it is considered that the aquifer is capable of supplying substantial quantities of ground water for use by Schweizer-Reneke, further tests and development of additional observation boreholes as outlined below are recommended to enable hydrological parameters to be calculated and thus optimum pumping rates and spacing between production boreholes to be determined namely,
- pumping of G34307 for 5 000 min. at a constant rate of 7 to 8 l/s after the installation of at least three additional observation boreholes namely,
- borehole No. 1 - 3 m west of G34309 to penetrate the fractured aquifer at a depth greater than 30 m,
- borehole No. 2, 40 m east of G34307, and perpendicular analyses to the dyke,
- borehole No. 3, some 1 000 m north of G34307, to penetrate the dyke at a depth greater than 30 m.

Pumping test interpretation may be possible using the type curves prepared by Gilding (1980) for fractured aquifers.

9. EXPLOITABLE GROUND WATER POTENTIAL

9.1 General

In order to ascertain whether the municipal water supply of Schweizer-Reneke could be augmented or supplemented from ground water sources when surface water supplies are inadequate (see section 2.1) an attempt has been made to quantify the exploitable ground water potential in the most favourable areas investigated in this report, making use of the limited geohydrological information obtained in the field. The estimates are at best considered to be only first estimates, and can be expected to be significantly in error. However an attempt has been made to err on the conservative side when making these estimates.

According to the variable yield analyses undertaken of Wentzel Dam it is evident that under the most adverse conditions of the third scenario ground water would be required to supplement surface water for a continuous period of up to twelve months during periods of surface water deficiency in order to meet the municipal water demand of Schweizer-Reneke. As already mentioned as supplementary supplies of ground water would be required during periods of subnormal rainfall, the recharge to ground water during these periods would be minimal. Any abstraction of ground water for extended periods would therefore depend largely on favourable ground water storage potential combined with good transmissive properties of the underlying geological formations.

In order not to seriously over-exploit the available ground water resources in any one area it would be advantageous to recover ground water from more than one location at a time, whenever ground water is required as a supplementary source of municipal water.

It will be possible to revise the first estimates of the exploitable ground water potential in the event of abstracting large quantities of ground water by continuously monitoring the long-term fluctuations of the ground water levels and by metering the quantities of ground water being abstracted.

In making these first estimates of the exploitable ground water potential in the vicinity of Schweizer-Reneke cognizance was taken of such factors as the potential recharge to the ground water systems, the storage potential of the systems for ground water and the potential yields of boreholes.

The following areas are considered favourable for the recovery of ground water on a scale sufficient for municipal use and will each be discussed in this chapter:

Schweizer-Reneke Townlands 62HO

Hamburg 82HO

Mimosa 61HO

Palachoema 64HO

Maraetchesfontein 54HO

9.2 Municipal owned land on Schweizer-Reneke Townlands 62HO

Pumping history during 1980 indicates that the established borehole field adjacent to the Harts River to the southwest of the town, can provide only $0,2 \times 10^6 \text{ m}^3$ of ground water annually without additional replenishment either from the river or from rainfall. Sufficient boreholes have already been established in the area and no opportunity exists for further development.

Due to unfavourable geohydrological conditions in the remainder of the townlands underlain by granite little scope exists for further development of ground water with the possible exception of immediately adjacent to Wentzel Dam where some leakage through the base of the dam could be expected. Borehole Ss 20 could provide about $0,1 \times 10^6 \text{ m}^3$ annually when pumped at 6 l/s for 12 hours per day.

The most promising water-bearing formations for development of ground water are expected to be the lava and quartzite formations located to the north of Schweizer-Reneke (see section 8.3). Boreholes penetrating both formations at suitable depths may be able to take

advantage of considerable ground water storage within the fractured formations. The success of the two exploration boreholes G34314 and G34315 indicates that a number of relatively low-yielding borehole could be developed in this area. Fig. 21 indicates two further possible sites viz G34049 and G34050. The recharge to ground water from rainfall in this area (about 25 km^2) could be of the order of $0,5 \times 10^6 \text{ m}^3$ annually. Four boreholes yielding 3 l/s each pumped for 12 hours a day could yield about $0,2 \times 10^6 \text{ m}^3$ annually when the aquifer is initially fully recharged. Estimates of ground water storage, however, is not possible in these fractured formations. At present there is no abstraction of ground water in this area.

The total exploitable ground water potential in the Schweizer-Reneke Townlands is therefore estimated to be of the order of $(0,2 + 0,1 + 0,2) \times 10^6 \text{ m}^3$ or, $0,5 \times 10^6 \text{ m}^3$ annually.

9.3 Hamburg 82HO

Ground water contouring indicates that the geological formations around borehole G34310 are recharged by ground water moving in a north-westerly direction. Ground water abstraction in the Hamburg stream catchment during 1981 was low, probably of the order of $0,03 \times 10^6 \text{ m}^3$. In the past the weathered and fractured granite aquifer has been exploited via borehole Hg 1 for irrigation purposes for several months a year. The maximum quantity of ground water that has been recovered from this borehole was probably less than $0,15 \times 10^6 \text{ m}^3$ annually.

Assuming a 5% infiltration from rainfall over the catchment of about 60 km^2 the average annual recharge to ground water is estimated to be about $1,5 \times 10^6 \text{ m}^3$. This is considerably more than the potential ground water storage in the catchment estimated at about $0,6 \times 10^6 \text{ m}^3$. It is considered that at least $0,6 \times 10^6 \text{ m}^3$ of ground water could be recovered from this catchment by judicious siting of production boreholes, which is equivalent to a continuous pumping rate of about 19 l/s . It would be reasonable to expect that boreholes G34310 (8 l/s), G34313 (2 l/s), Hg 1 (6 l/s) and a further

4 boreholes could provide at least $0,6 \times 10^6 \text{ m}^3$ annually when pumped for 12 hours per day. The additional boreholes would be spaced along a line parallel to the ground water contours as indicated in Fig. 20. Further sites could also be selected east and west of the Hamburg stream valley using geophysical methods to identify areas of maximum depth of weathering in the granite. This would increase the exploitable ground water potential beyond $0,6 \times 10^6 \text{ m}^3$ annually. The effect of heavy abstraction of ground water in this area would have the effect of lowering the ground water levels and so enhancing the opportunity of recharge from rainfall. Recharge could be increased further by constructing a dam in the Hamburg stream valley north of borehole G34310.

9.4 Palachoema 64H0

Boreholes Pa 4, 5 and 6 are currently pumped all the year round for irrigation purposes and are capable of a maximum combined delivery of about 18 l/s (refer to section 5.2). The weathered granite aquifer exploited has been proved to be reliable and productive. The maximum annual abstraction from these three boreholes is estimated to be about $0,3 \times 10^6 \text{ m}^3$ equivalent to a constant pumping rate of about 10 l/s. The fact that yields declined slightly in 1980 indicates that this is the maximum yield that can be relied upon. However some scope for further exploitation of ground water may exist however toward the western boundary of the property. The ground water is recharged mainly from direct rainfall recharge and flows in a south catchment south-easterly direction.

Borehole Pa 1 is located in a relatively unexploited sub-catchment to the west of boreholes Pa 4, 5 and 6. The ground water storage commanded by this borehole is estimated to be of the order of $0,15 \times 10^6 \text{ m}^3$ and the annual recharge is estimated to be about $0,4 \times 10^6 \text{ m}^3$. On catchment size considerations, the ground water potential is estimated to be also about $0,3 \times 10^6 \text{ m}^3$.

The total exploitable ground water potential therefore on the property Palachoema is estimated to be approximately $0,60 \times 10^6 \text{ m}^3$, equivalent to a continuous pumping rate of 19 l/s.

Additional production boreholes could be selected on the results of electrical resistivity soundings.

Ground water recovered from this area is conveniently situated for use in the Black residential townships located to the west of Schweizer-Reneke. An additional reservoir would probably be required.

9.5 Mimosa 61HO

It is considered that boreholes developed in the deeply weathered granite along the surface drainage courses in the vicinity of borehole G34320 could be capable of supplying an additional $0,2 \times 10^6 \text{ m}^3$ annually, equivalent to a continuous pumping rate of about 6 l/s.

Recharge and storage conditions in this area are not as favourable as at Hamburg and at Palachooma because of steep hydraulic gradients and comparatively deep ground water levels. Several possibly low-yielding boreholes similar to borehole G34320 could be developed yielding a maximum of 3 l/s each. At present virtually no ground water is being abstracted in the vicinity of borehole G34320. This area is conveniently near to Schweizer-Reneke.

9.6 Maraetchesfontein 54HO

The assessment of the exploitable ground water potential associated with the Maraetchesfontein dyke on the farms Maraetchesfontein and Panfontein is difficult. The fractured contact zone of the dyke, with evidently high transmissive properties but low ground water storage properties, is fed under pumping conditions from the adjoining lava formations with evidently low transmissive properties but relatively high ground water storage properties. The total extent of the influence of pumping borehole G34307 was not determined by the pumping test (see section 8.4) because of the linear configuration of the conduit system. It is estimated though that perhaps, as much as $0,5 \times 10^6 \text{ m}^3$ of ground water could be recovered annually from boreholes located along this feature.

Assuming a pumping rate of 8 l/s and 4 l/s respectively for boreholes G34307 and G34308 and a 12 hour per day pumping schedule a total of $0,2 \times 10^6 \text{ m}^3/\text{a}$ could immediately be recovered. Further boreholes to the north and south of borehole G34307 could be sited following further geophysical work and additional pumping tests.

Deep fracturing within the lava indicates the ground water storage within the lava may be more extensive than that in fractured and weathered granite and may be a more reliable source during periods of subnormal rainfall. During the period of investigation no boreholes were known to be recovering ground water from the vicinity of the dyke and abstraction from the lava was low.

This area however is a considerable distance from Schweizer-Reneke and development costs may probably be prohibitive.

10. SUMMARY

The following are the main findings of the investigation and which have been previously referred to in this report:

(i) Wentzel Dam will remain the primary source of municipal water for Schweizer-Reneke. By the year 2000 the projected municipal water demand will be $1,52 \times 10^6 \text{ m}^3$ per annum equivalent to a constant pumping rate of 50 l/s, with seasonal peak demands rising to 75 l/s. Supplementary supplies of ground water will be required for continuous periods of up to 12 months during periods of subnormal rainfall.

More surface water could be made available from Wentzel Dam if the irrigation rights of the remaining downstream irrigators are expropriated. Repairs to the irrigation canal could significantly reduce conveyance losses.

(ii) The existing borehole field recovering ground water south of Schweizer-Reneke adjacent to the Harts River has been estimated to have an exploitable ground water potential of

only $0,2 \times 10^6 \text{ m}^3$, without additional ground water recharge from the Harts River.

- (iii) The mean annual rainfall at Schweizer-Reneke from 1953 to 1981 was 509,5 mm/a. Subnormal rainfall over the period 1978-1980 was responsible for the water crisis developing in Schweizer-Reneke during 1980.
- (iv) Regional ground water levels have fully recovered from the effects of the subnormal rainfall.
- (v) Ground water flow within the Harts River basin is from the drainage divides to the environs of the Harts River and subsequently in a south-westerly direction with a shallow gradient of 1 in 10 000. South of the Vaal-Harts watershed ground water flow is to the south with a steeper gradient of 1 in 5 000.
- (vi) The numerous dykes present in the investigation area, inferred from a study of aerial photographs and aeromagnetic maps, do not impede the movement of the regional ground water flow.
- (vii) The principal geohydrologic units present are the granite-gneiss, quartzite-sandstone, lava, mudstones and various superficial deposits.
- (viii) Apart from the mudstone and superficial deposits all geohydrological units are fairly easily developed by the rural population to provide adequate supplies of ground water for domestic and stock watering purposes. Borehole yields though are generally low, less than 2 l/s, and only 41 boreholes of the surveyed boreholes (including five recently drilled exploration boreholes) have reported yields in excess of 5 l/s. Regional ground water abstraction is therefore low in aggregate.

(ix) Irrigation from ground water resources is limited. During the agricultural season 1980/81 only twelve boreholes were pumped at rates in excess of 5 l/s from eight properties to provide about $1,1 \times 10^6 \text{ m}^3$ for irrigation purposes. Yields from these boreholes were reasonably reliable but declined slightly in 1980 from boreholes receiving ground water from the granite.

(x) The most favourable geological formations for the recovery of ground water on a scale sufficient for a municipal supply are weathered zones within the granite and in the lavas (particularly near surface drainage courses), deep fractures within the lavas, and fractures associated with the quartzite-granite contact. It has been proved by exploration drilling that dyke-lava contact zones and lava-quartzite contacts are also capable of yielding good supplies of ground water.

(xi) Geophysical methods can be successfully applied to locate moderately yielding boreholes within weathered granite, dyke-lava contacts and possibly dyke-granite contacts. Geophysical methods can also be successfully employed to determine the presence of magnetized amphibolite bodies within the granite and to delineate those areas overlain by Karoo mudstones. Boreholes penetrating lava-quartzite contacts are best sited geologically.

(xii) Ground water in general is of good quality except for relatively high levels of nitrate. Ground water is only moderately mineralized. Mineralization is greater in Karoo formations and lavas than that from granite and quartzite.

Nitrate levels though generally high seldom exceed 102 mgm/l. Apart from ground water occurring in the granite of the Schweizer-Reneke Townlands, ground water remaining in other locations favourable for exploitation are only marginally high in nitrate concentrations. The

chemical composition of the ground waters occurring in the area is characteristically not dominated by any particular cation or anion type.

- (xiii) Pumping tests on borehole G34310, on Hamburg 82HO, indicate that this borehole is capable of supplying about 1 000 m³/day during a 24-hour pumping schedule. The weathered and fractured granite is semi-confined and is partially fed by gravity drainage from the overlying less permeable decomposed granite. Transmissivity values average about 120 m²/d.
- (xiv) Borehole Pa 1, Palachoema 34HO, has been shown to be capable of yielding 1 000 m³/d during a 24-hour pumping schedule. Conditions within the weathered granite are semi-confined or semi-unconfined. Transmissivity values average about 230 m²/d.
- (xv) Borehole G34325, Schweizer-Reneke 62HO, is probably capable of supplying 350 m³/d. The multiple fracture system is probably semi-confined. A transmissivity value of 70 m²/d was estimated from recovery data after a pump test.
- (xvi) Borehole G34307, Maraetchesfontein 54HO, could possibly be capable of yielding 850 m³/d. Ground water movement is by fissure flow along a preferential zone of transmissivity namely the Maeretchesfontein dyke. Pumping tests did not permit the calculation of any geohydrologic parameters.
- (xvii) Estimates of the exploitable ground water potential in favourable geological formations indicates that adequate supplies of ground water are available within a radius of 10 km of Schweizer-Reneke. Provisional estimates of these exploitable supplies are summarized below:

Schweizer-Reneke Townlands 62HO

- ° Borehole field Ss 2 to Ss 11, located 1 km from town centre
- = $0,2 \times 10^6 \text{ m}^3$

- ° Borehole Ss 20 adjacent to Wentzel Dam, located 2,8 km from town centre
- = $0,1 \times 10^6 \text{ m}^3$

- ° Area centred on borehole G34315, located about 4,0 km from town centre
- = $0,2 \times 10^6 \text{ m}^3$

Mimosa 61HO

Area centred on borehole G34320, located about 4,5 km from town centre

- = $0,2 \times 10^6 \text{ m}^3$

Hamburg 82HO

Area centred on borehole G34310, located about 7,5 km from town centre

- = $0,6 \times 10^6 \text{ m}^3$

Palachoema 64HO

- ° Area centred on borehole Pa 1, located about 7,7 km from town centre
- = $0,3 \times 10^6 \text{ m}^3$

- ° Area centred on boreholes Pa 4, 5 and 6, located about 6,4 km from town centre
- = $0,3 \times 10^6 \text{ m}^3$

Maraetchesfontein 54HO

Area centred on borehole G34307, located about 10 km from town centre

$$= 0,5 \times 10^6 \text{ m}^3$$

In total this gives an exploitable ground water potential of about $2,4 \times 10^6 \text{ m}^3$.

This amount is sufficient to meet the projected municipal water demand of Schweizer-Reneke up to about the year 2010. However this potential is not all located in one area but in many areas within a radius of about 10 km of the town centre.

In chapter 11 some views are expressed on what approach should be adopted to integrate the available ground water resources into the present municipal water supply.

11. RECOMMENDATIONS

The results obtained from this geohydrological investigation, and as recorded in this report, indicate that sufficient ground water to supplement existing municipal water supplies during extended periods of surface water shortage cannot be met solely from ground water resources occurring on the municipal townlands. However adequate supplies of good quality ground water (apart from moderately high concentrations of nitrate) are available within various favourable geological formations within a 10 km radius of Schweizer-Reneke. The development of alternative ground water resources located at greater distances from Schweizer-Reneke namely on the farms Kameelkuil 88HO and Springbok 191HO, which are known to occur but which were not investigated at this stage, are not considered feasible options as supplementary supplies of municipal water.

As indicated in this report although the exploitable ground water potential is considerable, conservatively estimated at about $2,5 \times 10^6 \text{ m}^3$, it is distributed over an extensive area around

Schweizer-Reneke in various geological environments. This fact does impose additional problems of a development and management nature. It is recommended that the ground water resources available to the municipality be integrated into the municipal water supply system in an ordered sequence based primarily on accessibility considerations. The various recommended phases of integration are enumerated below.

Develop fully the available ground water resources on the Schweizer-Reneke Townlands as previously reported on. This would make provision for about $0,5 \times 10^6 \text{ m}^3$ of municipal water annually sufficient to meet a large proportion of the town's present municipal water demand.

This ground water could also be used to supplement the normal municipal supply from surface water sources during periods of peak demands experienced in early summer. This peak demand is expected to reach a continuous pumping rate of 75 l/s by the year 2000.

At a later stage and depending on the actual increase in demand for municipal water in Schweizer-Reneke, and also on the actual demands on the existing surface water resources made by downstream irrigators a second phase of ground water development could be initiated. This would necessitate developing a number of production boreholes either on the farm Hamburg or on the farm Palachoema. Palachoema is located slightly nearer to the town centre and has the advantage of being at a higher elevation and is conveniently situated as a supply point for the Black residential township to the west of Schweizer-Reneke. Both Hamburg and Palachoema could supply an estimated $0,6 \times 10^6 \text{ m}^3$ of ground water annually. Together with the ground water supplies developed in phase one a total supplementary excess of $1,0 \times 10^6 \text{ m}^3$ annually could be developed sufficient to meet the municipal water demand of Schweizer-Reneke until beyond 1990.

It is accepted that the surface water resources in Wentzel Dam will continue to supply the bulk of the municipal water of Schweizer-Reneke and that the alternative sources of ground water will be made available only for short critical periods. In order to

improve on the estimates made of the exploitable ground water potential in certain areas it is recommended that ground water be recovered at this stage for short periods to relieve the seasonal demands for municipal water. To achieve this improvement additional field data would need to be monitored. Ground water of high nitrate levels could successfully be blended with surface water from Wentzel Dam to achieve acceptable potable water quality standards.

Once a decision has been reached on the various options available to the municipality additional production boreholes can be sited and brought into service at reasonably short notice.

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APPENDIX 1

1. Index of investigated farms surveyed.
2. Completed borehole survey forms.
3. Water levels in observation boreholes.

APPENDIX 1

INDEX OF INVESTIGATED FARMS

CADASTRAL FARM NAME	FARM NUMBER	BOREHOLE PREFIX
Bellevue	68 HO	Be 68/
Belvedere	47 HO	Be 47/
Blyvooruitzigt	51 HO	Bt
Damplaats	38 HO	Ds
De Park	87 HO	Dk
Doornplaats	55 HO	Dt
Elim	24 HO	Em
Geluk	56 HO	Gk 56/
Geluk	81 HO	Gk 81/
Groenpan	23 HO	Gn
Grootpoort	83 HO	Gt
Hamburg	82 HO	Hg
Holpan	59 HO	Hn 59/
Jacobsdal	25 HO	J1
Jala Jala	10 HO	Ja
Kameelkuil	88 HO	K1
Koppiesfontein	52 HO	Kn 52/
Krompan	85 HO	Kn 85/
London	112 HO	Ln
Lot	48 HO	Lt
Lot 25	9 HO	Lt 9/-
Lot 26	16 HO	Lt 16/
Lot 22	17 HO	Lt 17/
Lot 22	18 HO	Lt 18/
Lot 20	19 HO	Lt 19/
Lot 19	20 HO	Lt 20/
Lot 49	39 HO	Lt 39/
Lot 15	44 HO	Lt 44/
Lot 16	45 HO	Lt 45/
Lot 6	48 HO	Lt 48/
Lot 7	49 HO	Lt 49/
Lot 9	63 HO	Lt 63/-
Luciana	43 HO	La

CADASTRAL FARM NAME	FARM NUMBER	BOREHOLE PREFIX
Maraetchesfontein	54 HO	Mn
Mimosa	61 HO	Ma 61 /
Mooilaagte	91 HO	Me
Niekerksrust	79 HO	Nt 79 /
Nooitgedacht	86 HO	Nt 86 /
Palachoema	64 HO	Pa
Panfontein	58 HO	Pn
Pasop	67 HO	Pp
Pierville	41 HO	Pe
Rietput	60 HO	Rt
Schoondal	22 HO	S1
Schweizer-Reneke Townlands	62 HO	Ss
Seringboomkop	66 HO	Sp
Springbok	191 HO	Sk
Vaalpoort	84 HO	Vt
Weltrevreden	28 HO	Wn
Zandfontein	90 HO	Zn
Zorgvliet	80 HO	Zt

COMPLETED BOREHOLE SURVEY FORMS

KAD. PLAASNAAM EN NR.

GEBIED.

2. estimaat
van toepassing
van kaart en/of lugfotoverwyssing 1: 50000 streeks
1: 250000 streeks
1: 250000 streeks.

Periode van Opname SEPTE- DEC 1921.

B.G. NR.	Naam van Eienaar en adres	Coördinate N. of Plat N.	Totaal grootte van watergebruik	Gebiedskewe- ring en datum bewerkingstempo van watergebruik	Onttrekkingsteveelheid van grondwater	Gebiedskewe- ring en datum bewerkingstempo van grondwater	Gebiedskewe- ring en datum bewerkingstempo van grondwater	13		14		15		
								1	2	3	4	5	6	
Kn 59/1	J.H. FOURIE	59 HO (PREFIX HN SP/)	1											
Ha 59/2	al.ito.	d+s P 2s												
Farm : JAKA JAKA 10 HO (PREFIX Ja.)														
J.J. A.J.J. JORDAN														
Jr. 2 ditto			0,4											
Jr. 3 ditto			0											
Jr. 4 ditto			0											
Farm : JACOBSEAL 25 HO (PREFIX Je.)														
Jr. 1 S. BAGNAUT			0											
Jr. 2 ditto			1,3											
Jr. 3 136037/5 ditto			0											
Farm : KATIEZKUL 33 HO (PREFIX Kc.)														
KL 1 J.S. VILJOEN			0											
KL 2 A.J. VAN DER REEVE		d+s t	1375											
KL 3 ditto			0											
KL 4 ditto			0,4											
KL 5 ditto			0											
KL 6 absent			0											
Farm : KOOPLES FONTEIN 52 HO (PREFIX KN&S/)														
Kn 52/1 VAN TONDER		d w.p.												

* last pumped by Agrius machinery (c.c.)
for 20 days at rate 2,5 to 25 t/ha.
Pump out at 70 m.

pump running.

KAD. PLAASNAAM EN NR.

GEBIED.

Periode van Opname **SEPT- DEC 1951**
 Kaart en/of Lugfotoverwyssing 2725 RA, AB, AC + AD.
 e - estimatief
 b - met.
 Z

1	2	3	4	5	6	7	8	9	10	11	12	13	14	15
BO. NR.	Naam van Elment en adres	Geometrie nr	Totale grondoppervlakte van watergebruik	Geometrie leeuw	Watergebruik	Onttrekkingsoort	Watergebruik	Oppervlakte onder bestrooiing uit grondwater						
1475/1	Farm : KROMPAN	85 HO	(PREFIX K-85)											
1475/2	Farm : FOUBIE	5 W.P.												
1475/3	Farm : LUCIANA	43 HO	(PREFIX L-43)											
1475/4	Farm : Absent.	0												
1475/5	Farm : LONDON	112 HO	(PREFIX L-112)											
1475/6	D.E.H. PUTTER	5 T	131971	6										
1475/7	Farm : Absent.	0												
1475/8	Farm : AOT	48 HO	(PREFIX L-48)											
1475/9	K. DUVIER	5 P												
1475/10	Farm : A.S.J. JORDAN	25 S	9 HO	(PREFIX L-9)										
1475/11	Farm : A.S.J. JORDAN	26 S	16 HO	(PREFIX L-16)										
1475/12	P. GOOSSEN	0 W.P.	175											
1475/13	A.S.J. JORDAN	0	0,3											
1475/14	H.B. SMITH	0 W.P.	0,19											
1475/15	A. H. T.	0,5												
1475/16	R.G. ERASMUS	222 17 HO	(PREFIX L-17)											
1475/17	Farm : Absent	20 19 HO	(PREFIX L-19)											
1475/18	Farm : Absent	20 19 HO	(PREFIX L-19)											
1475/19	Farm : Absent	5 W.P.												
1475/20	F. STRANSS	0												
1475/21	P. GOOSSEN	0 W.P.	1,25											
1475/22	Farm : Absent	22 18 HO	(PREFIX L-18)											

Prop
munity.

WATER LEVELS IN OBSERVATION BOREHOLES

C3NO 46 (Ss 4)

Monthly automatic
recorder

C3NO47 (G34309)

H.W.K. meter

Date

Water level

Date

Water level

(m)

(m)

30/07/1980	16,3	19/10/1981	6,06
30/08/	17,9	24/10/	6,00
02/10/	17,0	03/12/	6,09
03/11/	19,3	15/01/1982	5,80
26/11/	19,8	13/02/	5,32
13/01/1981	13,0		
02/02/	12,1		
03/03/	11,6		
01/04/	7,2		
01/10/	6,54		
12/10/	6,44		
02/11/	6,81		
27/11/	7,07		
08/12/	6,94		
08/01/1982	7,3		
08/02/	7,5		

APPENDIX 2

EXPLORATION BOREHOLE LOGS

G32693 and G34306 to G34320

FIG. I

G 32693

CASING GEOHYDROLOGY DEPTH PROFILE. DESCRIPTION

FIG. 2

G 34306

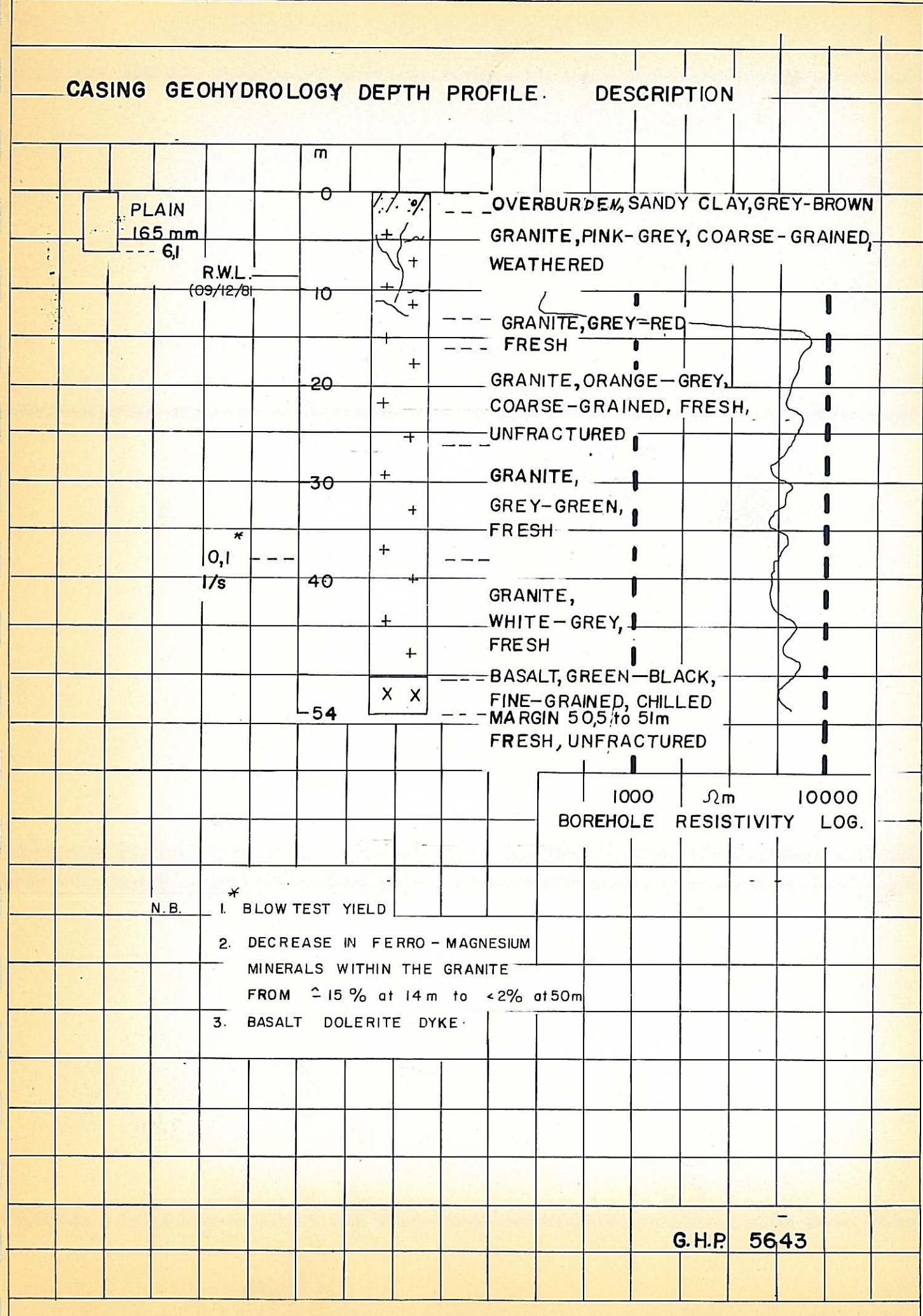


FIG. 3

G 34307

CASING GEOHYDROLOGY DEPTH PROFILE. DESCRIPTION

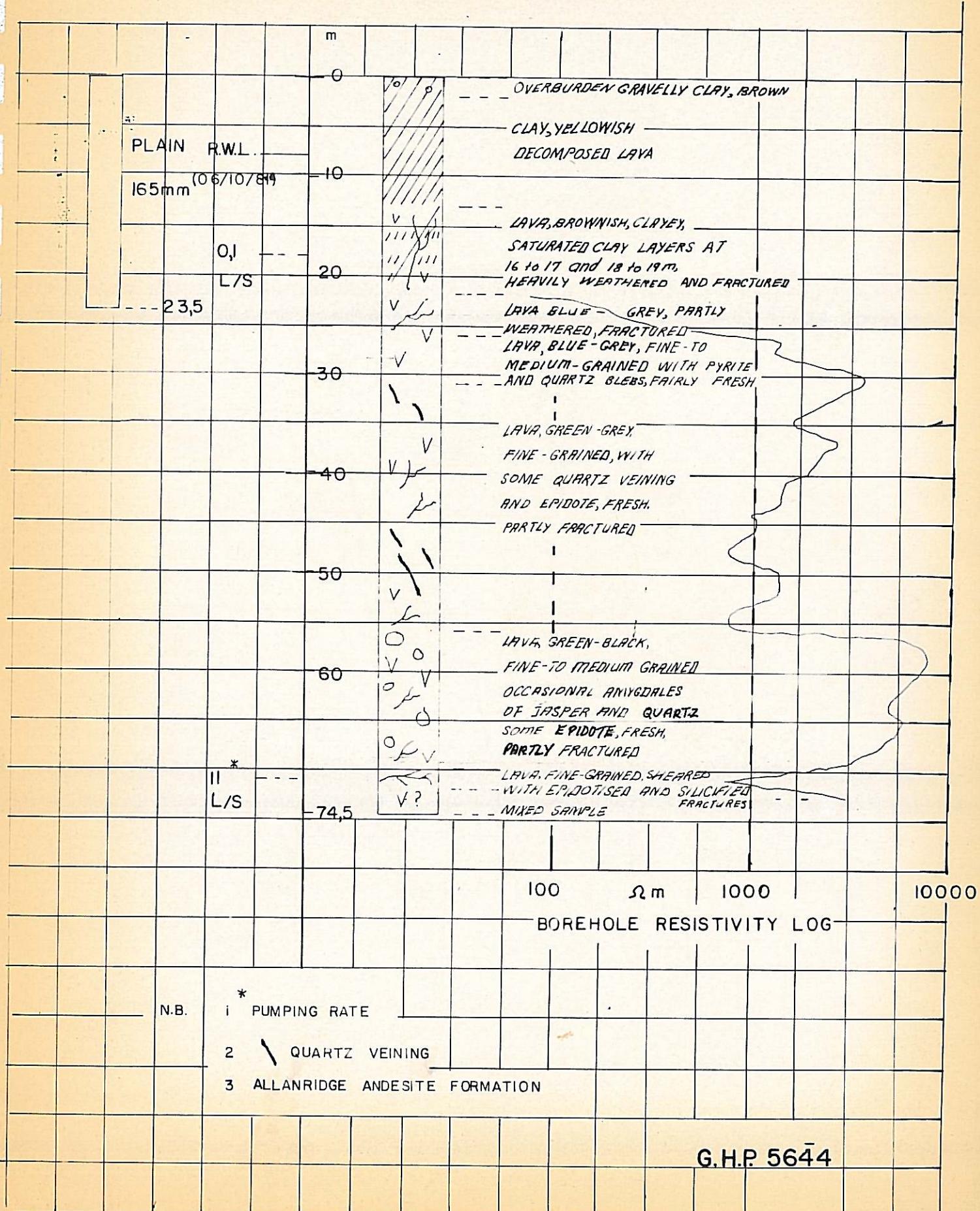


FIG.4

G 34308

CASING GEOHYDROLOGY DEPTH PROFILE. DESCRIPTION

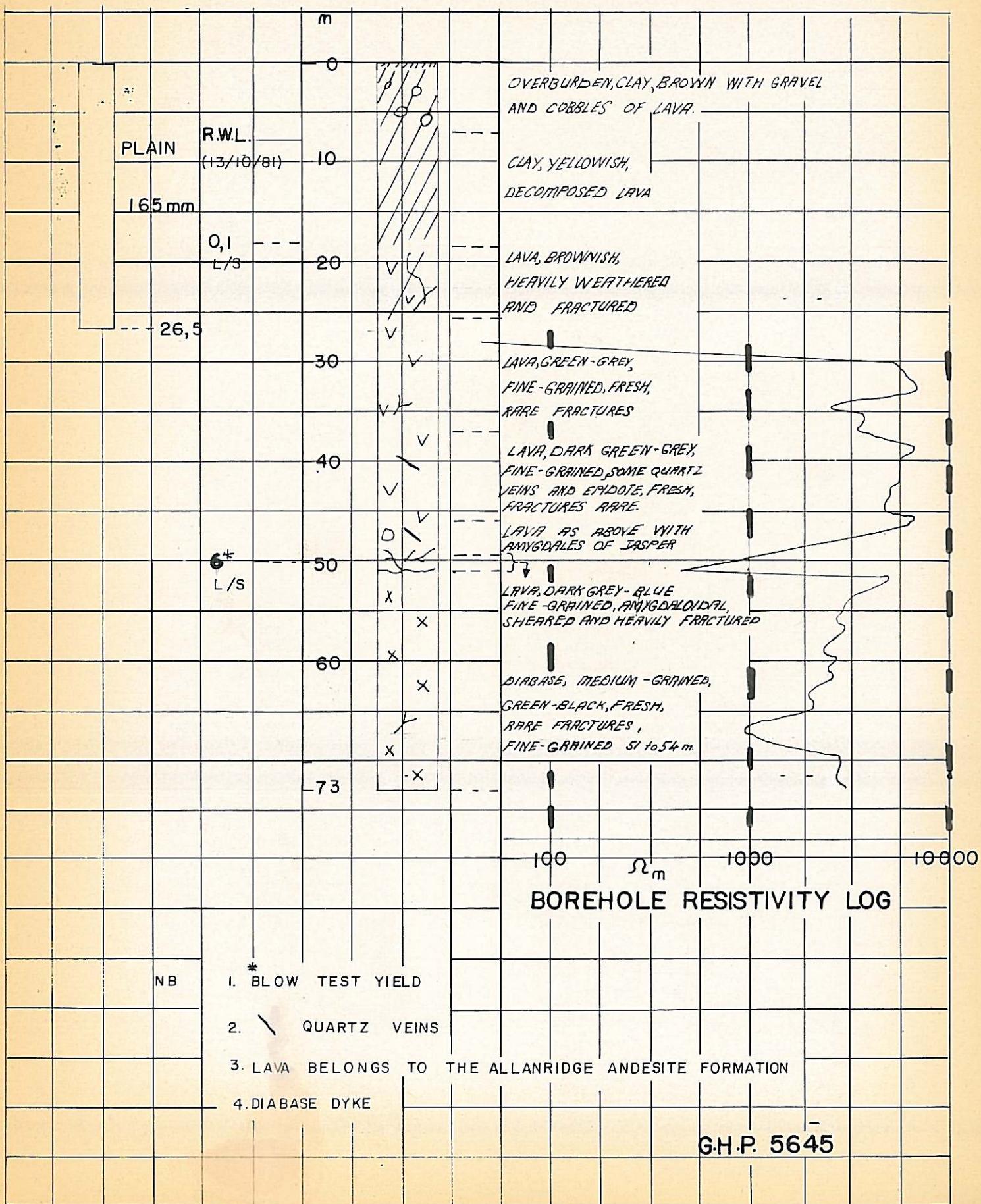


FIG. 5

G 34309

CASING GEOHYDROLOGY PROFILE. , DESCRIPTION

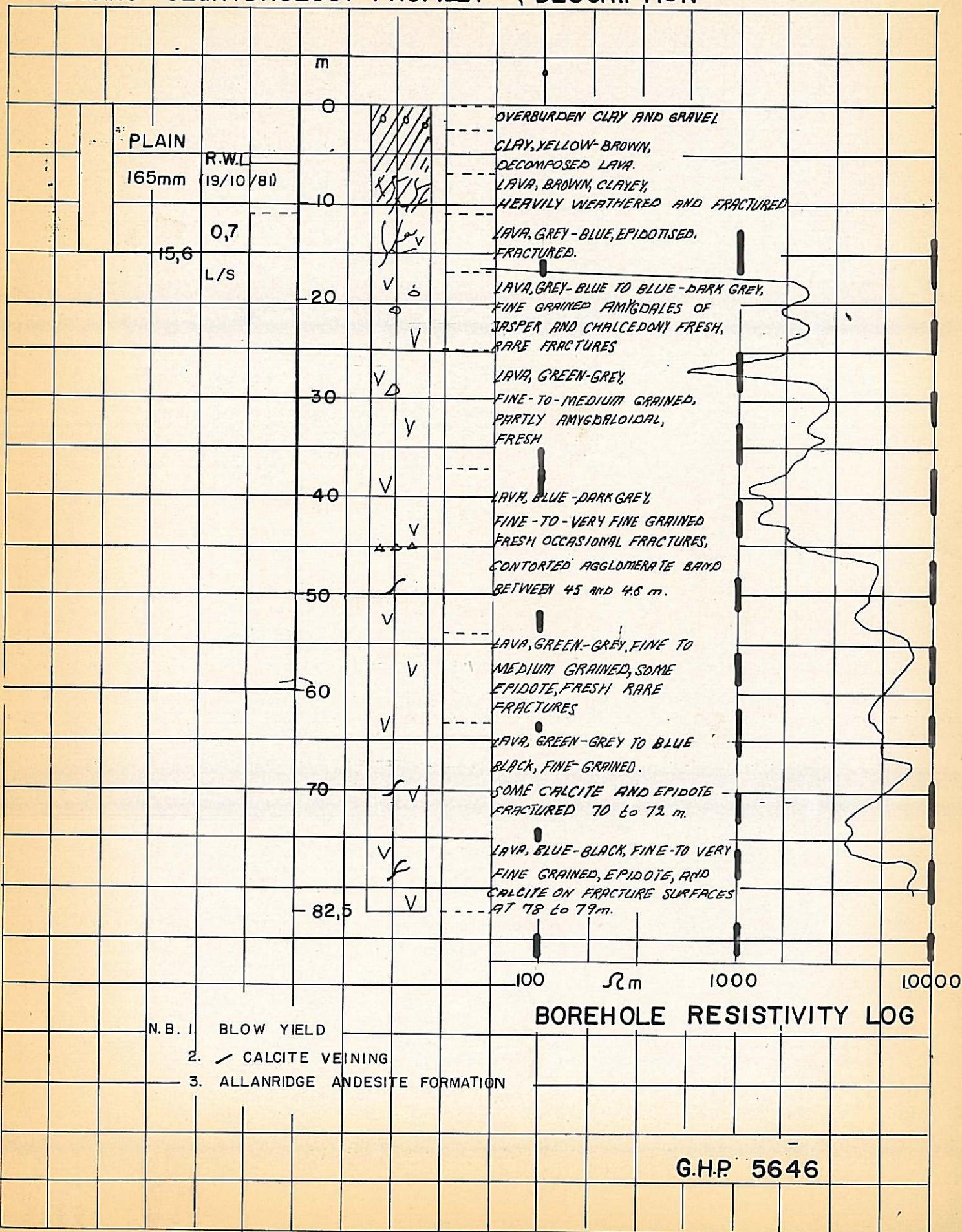
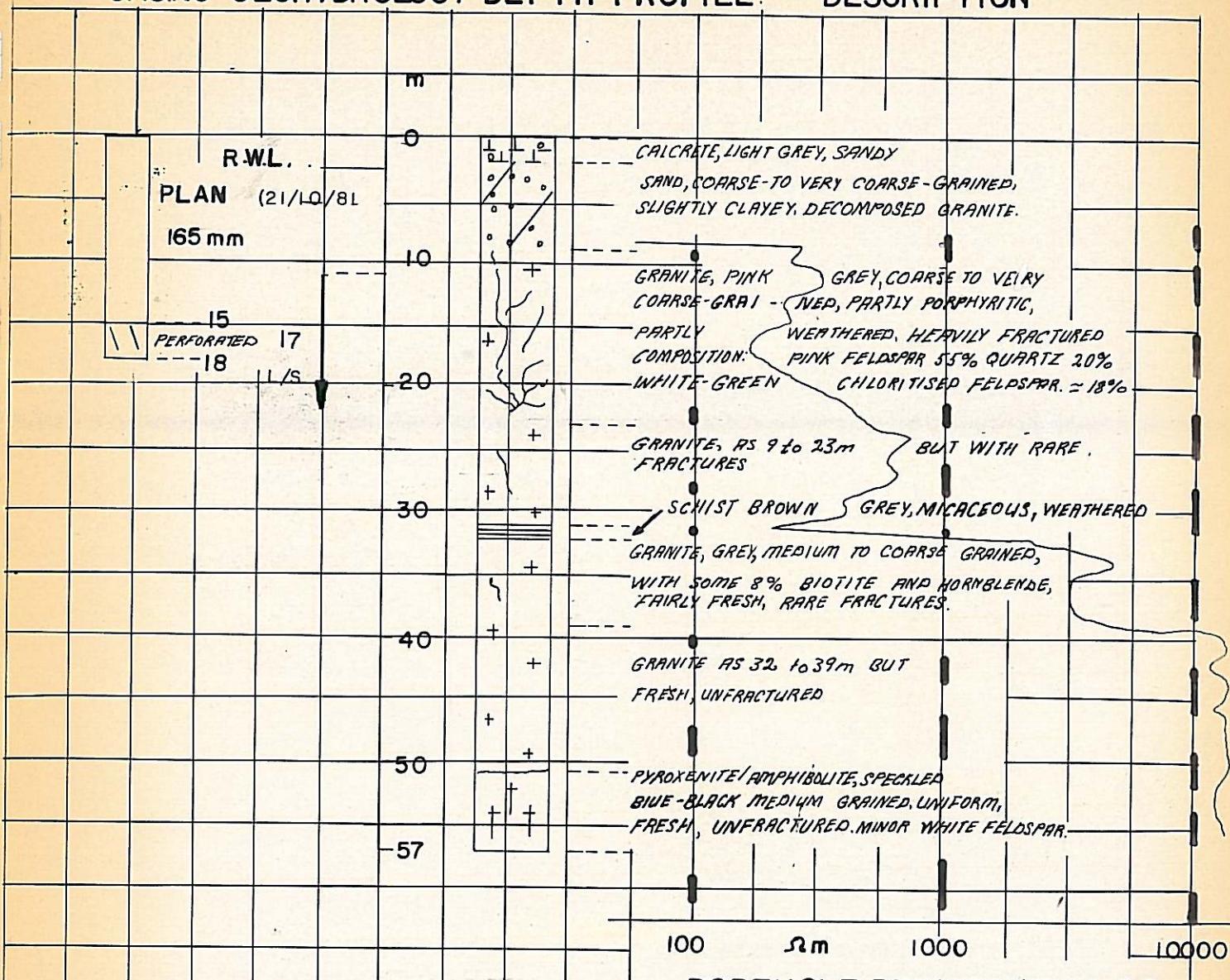


FIG.6

G 34310

CASING GEOHYDROLOGY DEPTH PROFILE DESCRIPTION

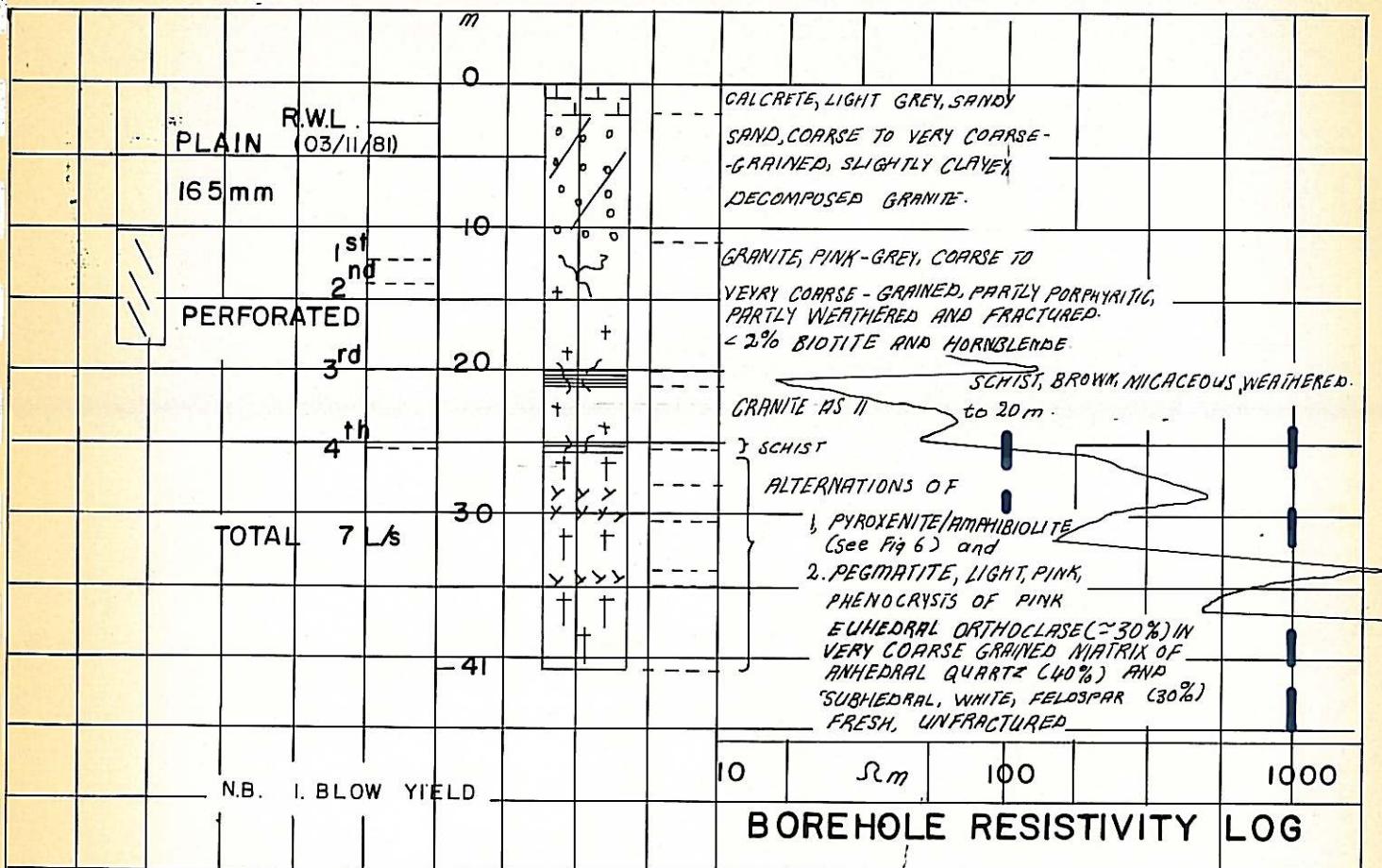


G.H.P. 5647

FIG. 7

G 34312

CASING GEOHYDROLOGY DEPTH PROFILE. DESCRIPTION



G 34311

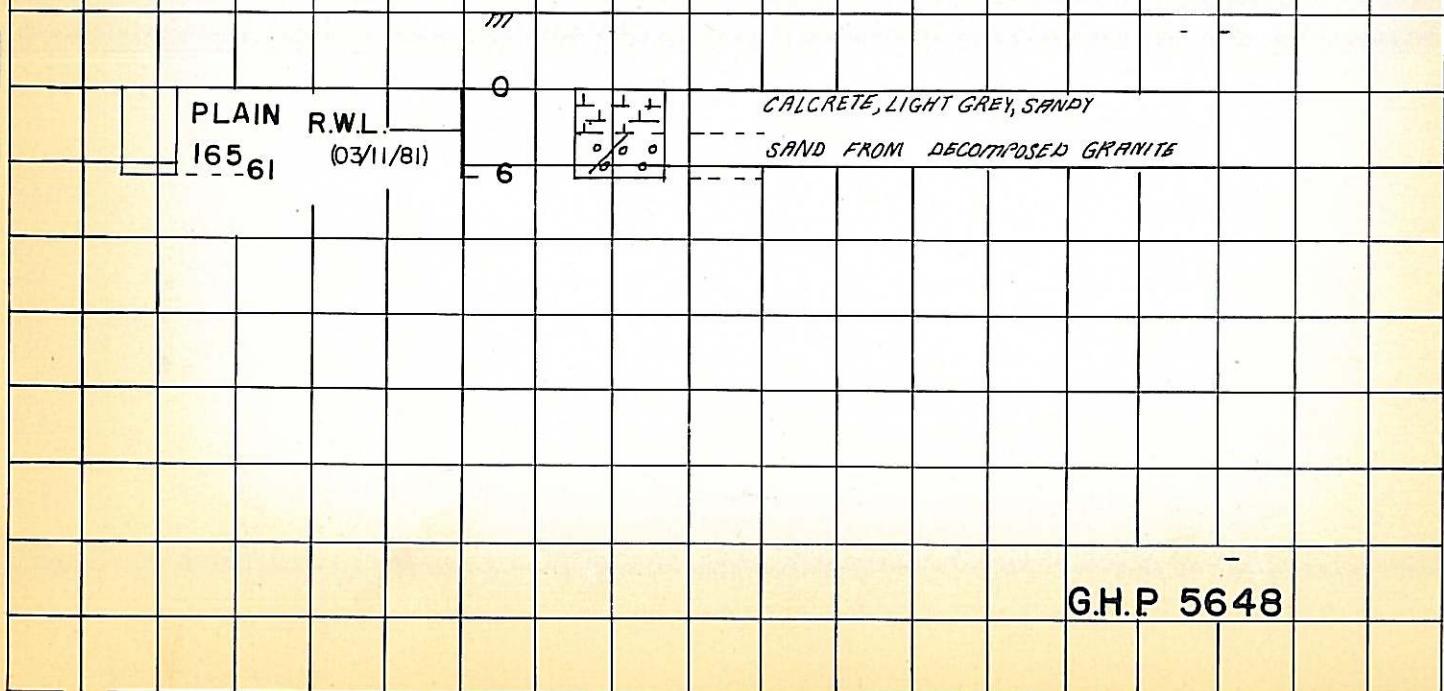


FIG. 8

G 34313

CASING GEOHYDROLOGY DEPTH PROFILE. DESCRIPTION

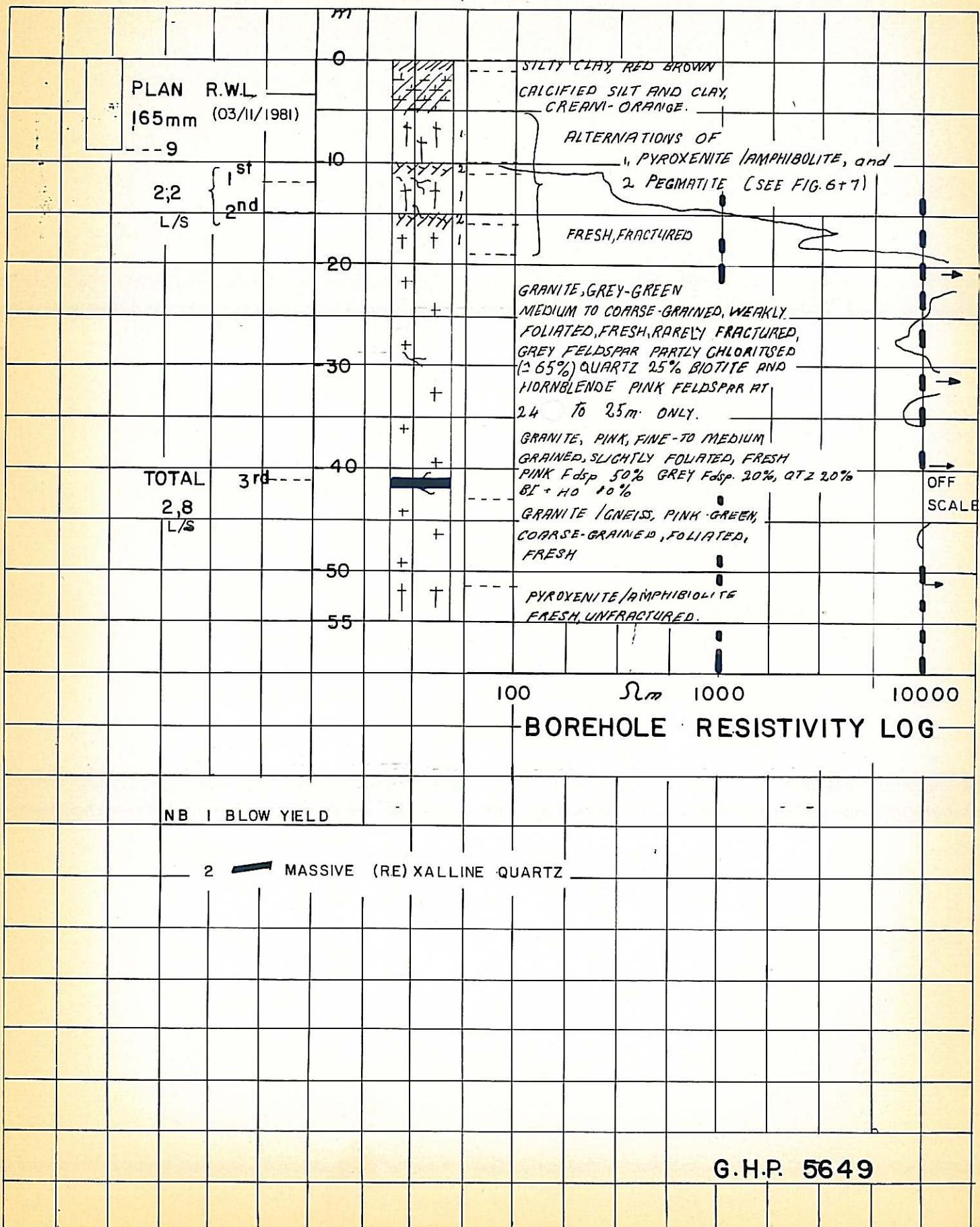


FIG. 9

G 34314

GEOHYDROLOGY DEPTH PROFILE. DESCRIPTION

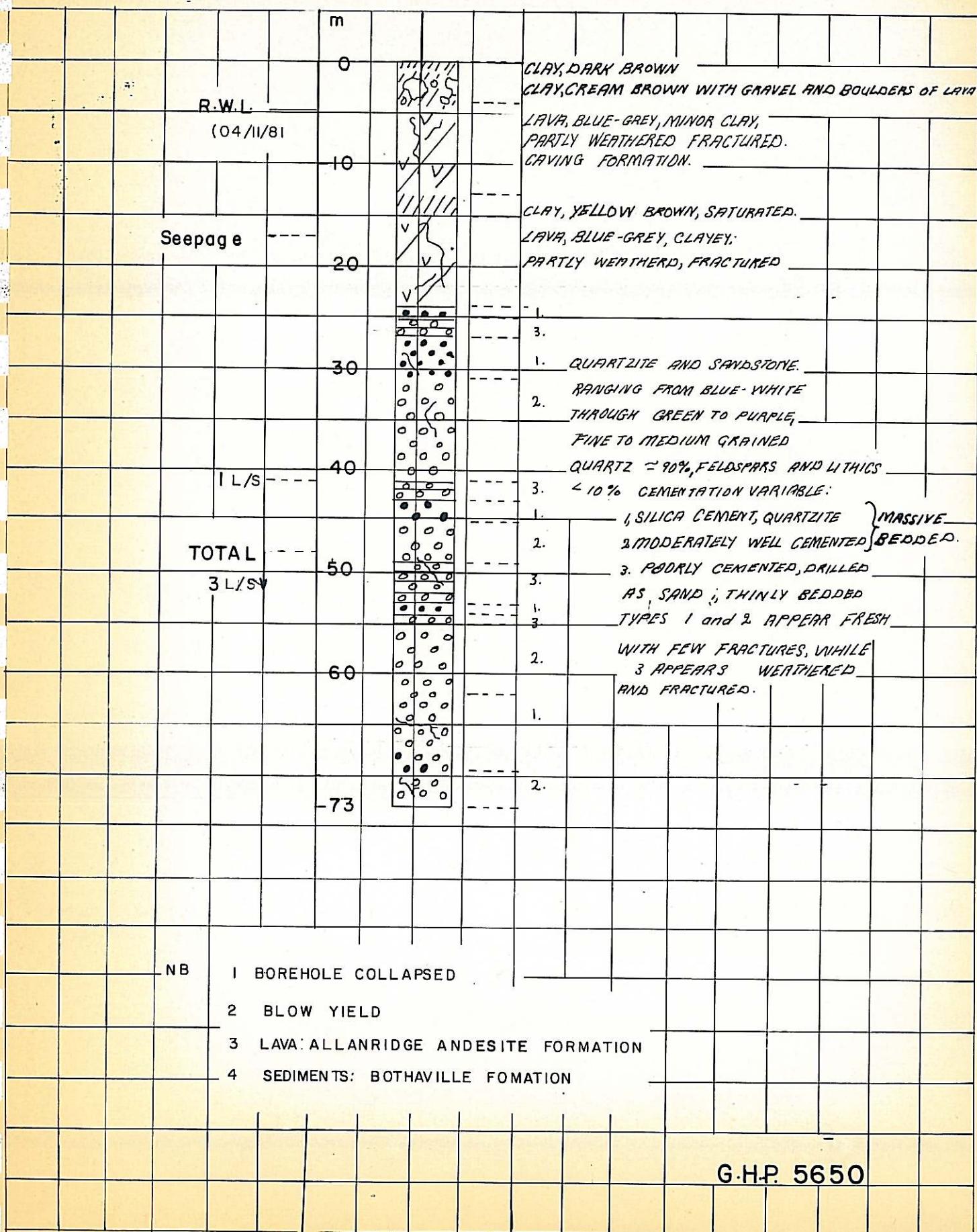
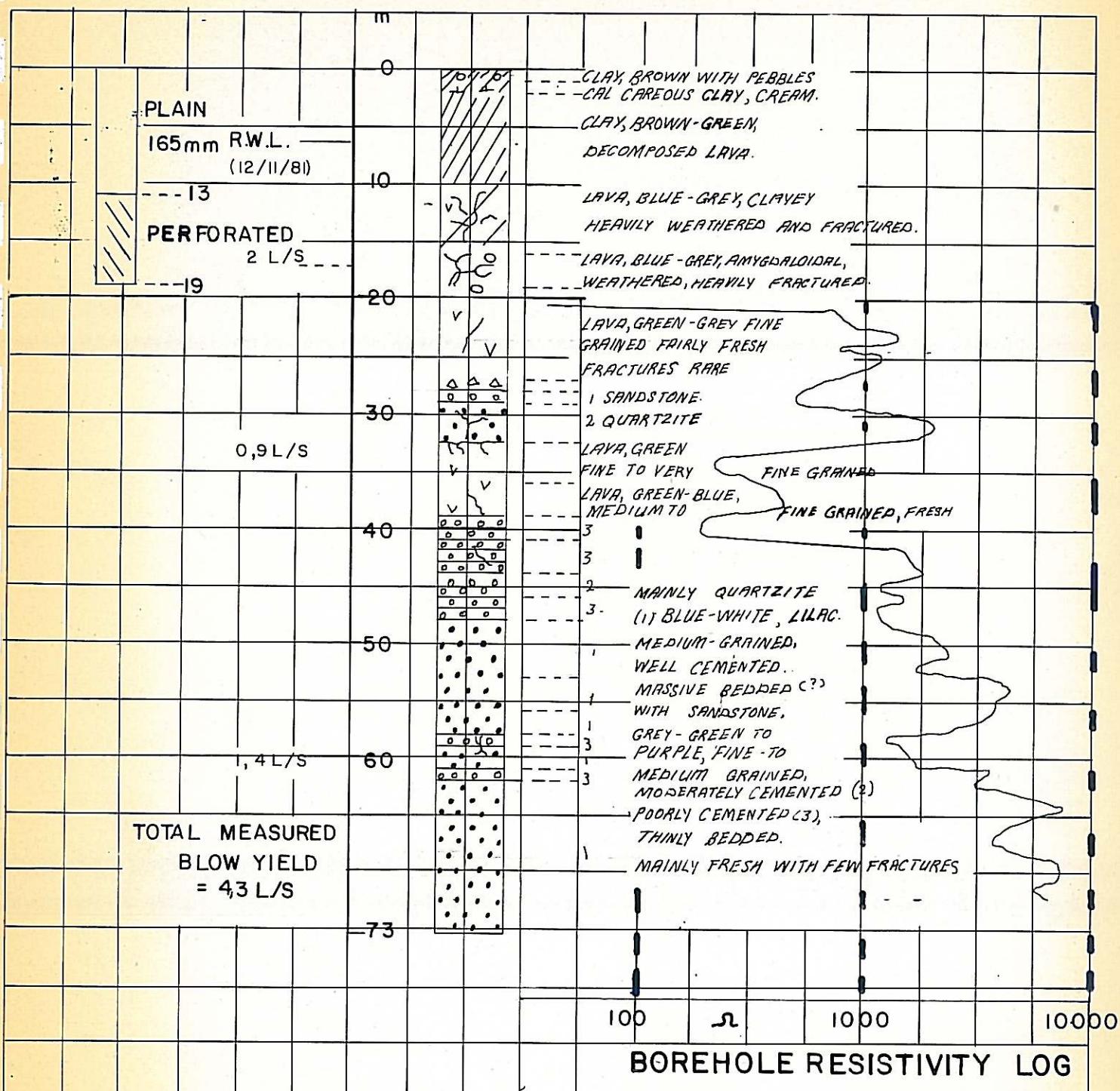


FIG.10

G 34315

CASING GEOHYDROLOGY DEPTH PROFILE . DESCRIPTION



NB 1 PUMPING RATE 5,5 L/S AFTER 6 HOURS

2 LAVA: ALLANRIDGE ANDESITE FORMATION

3. SEDIMENTS: BOTHAVILLE FORMATION

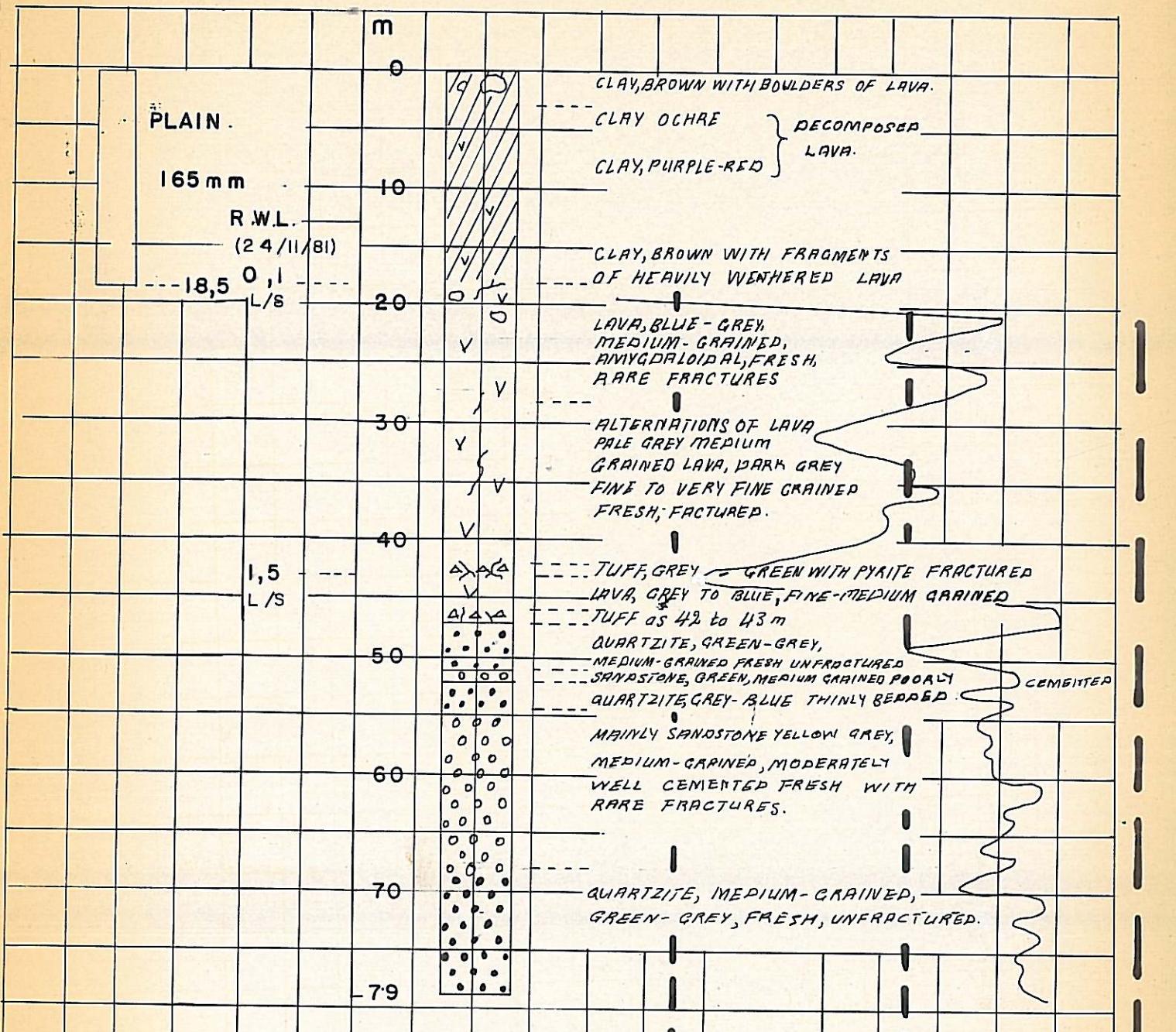
4 Δ Δ TUFF, GREY-GREEN

G.H.P. 5651

FIG. II

G 34316

CASING GEOHYDROLOGY DEPTH PROFILE. DESCRIPTION



N.B. 1. BLOW YIELD

2. LAVA: ALLANRIDGE ANDESITE FORMATION

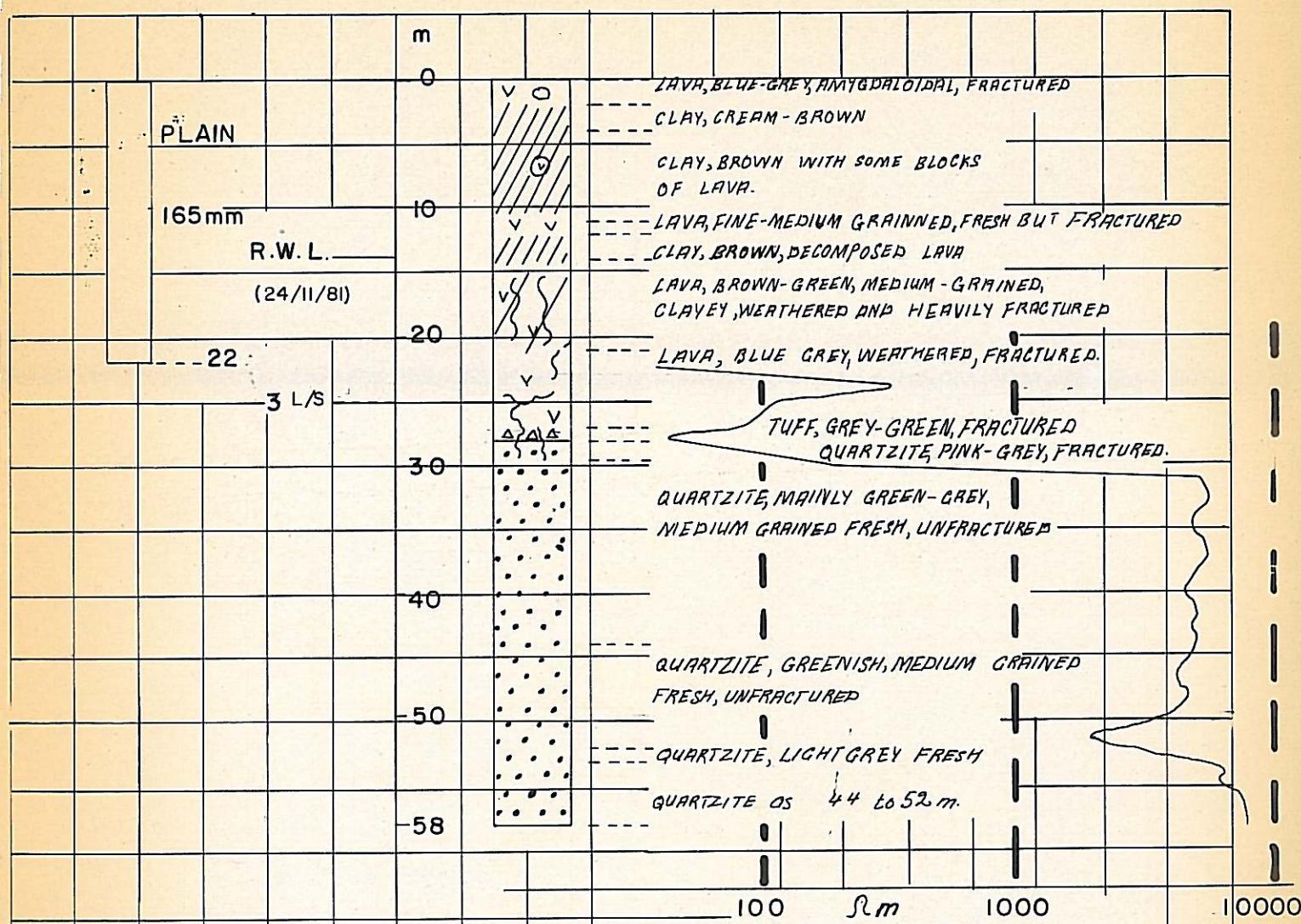
3. SEDIMENTS: BOTHAVILLE FORMATION.

G.H.P 5652

FIG. 12

G 34317

CASING GEOHYDROLOGY DEPTH PROFILE. DESCRIPTION



N.B. 1 BLOW YIELD

2. LAVA: ALLANRIDGE ANDESITE FORMATION

3. SEDIMENTS: BOTHAVILL FORMATION

G.H.P. 5653

FIG. 13

G 34318

CASING GEOHYDROLOGY DEPTH PROFILE DESCRIPTION

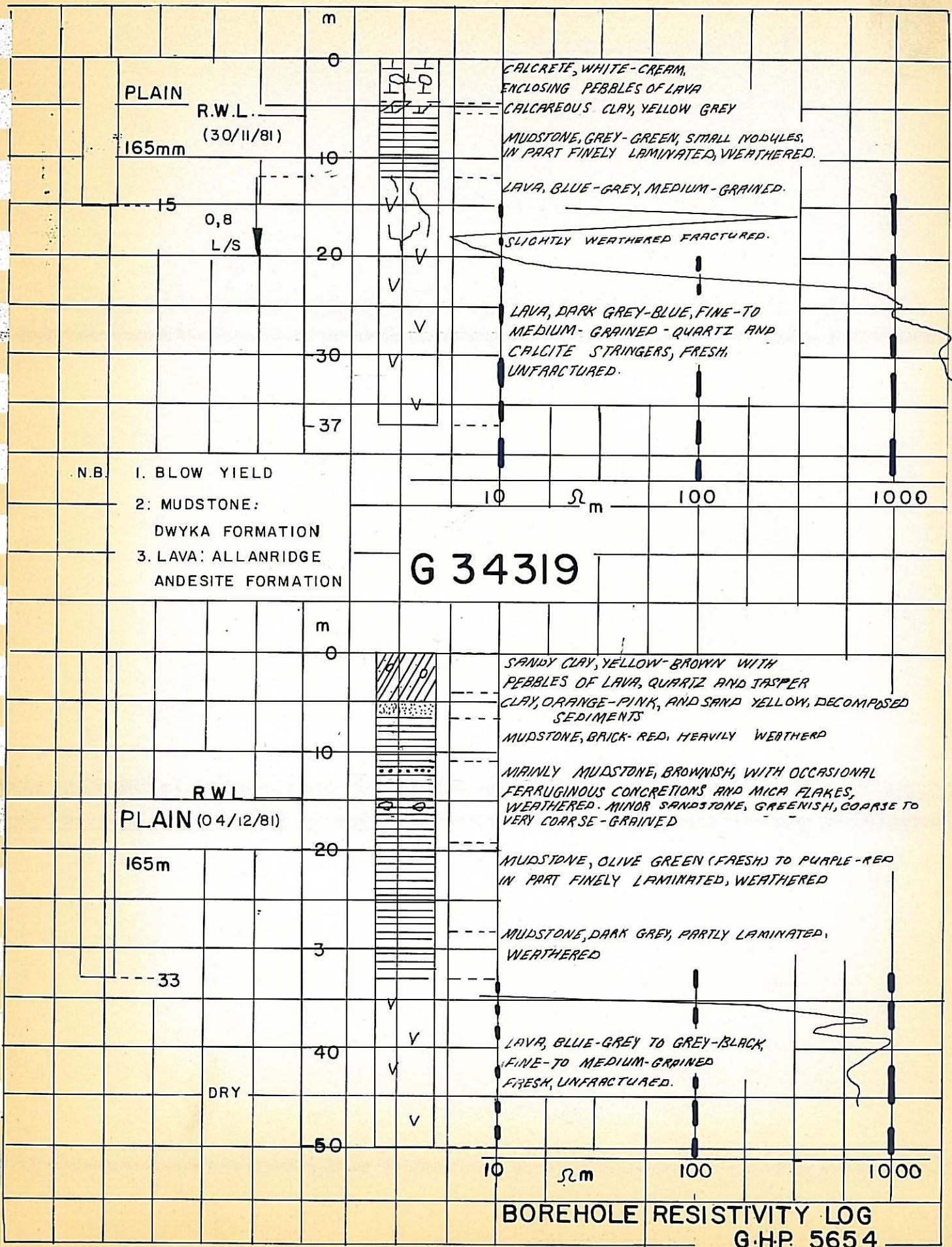
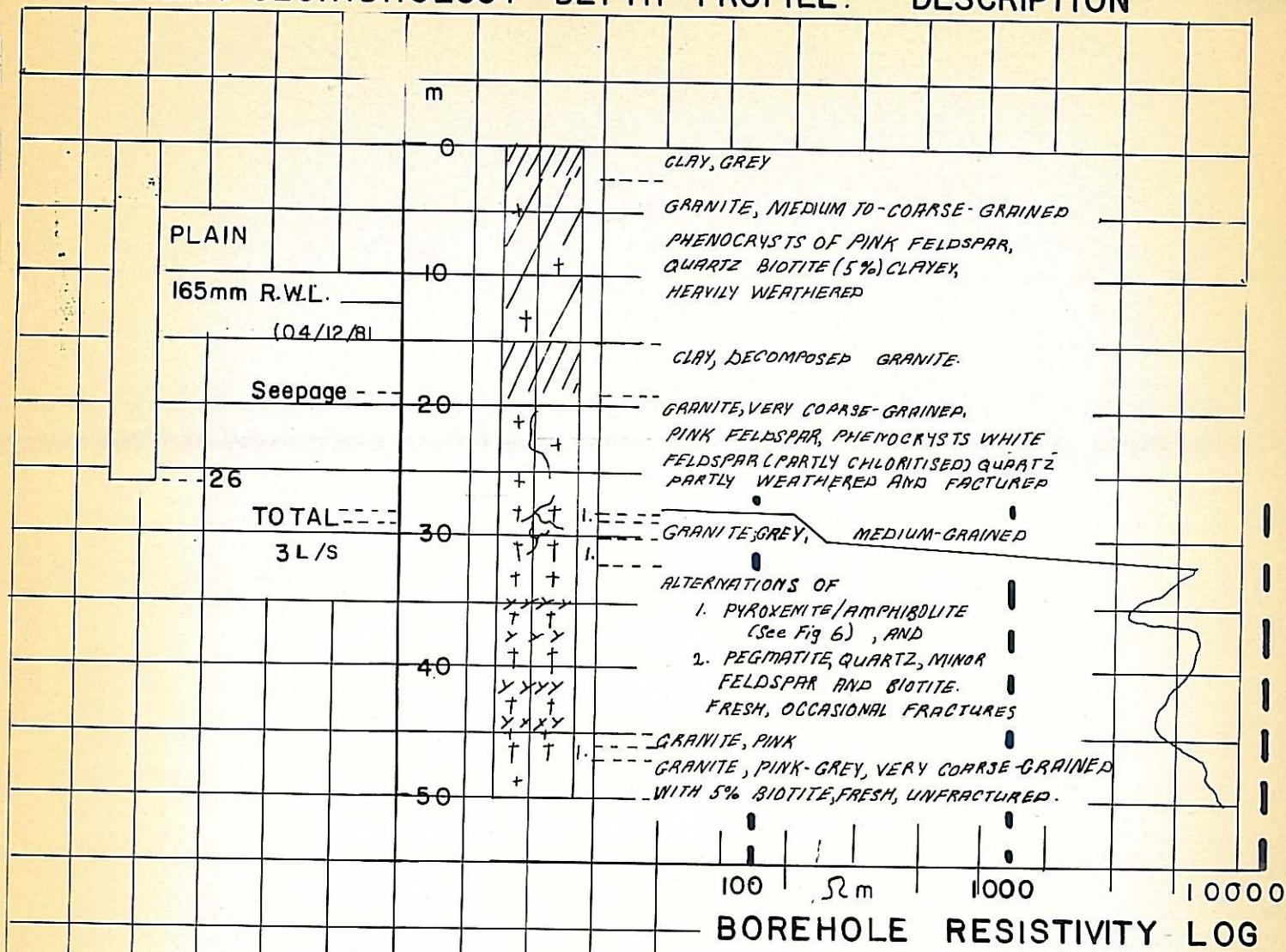


FIG. 14

G 34320

CASING GEOHYDROLOGY DEPTH PROFILE. DESCRIPTION



N.B. 1. BLOW YIELD

G.H.P. 5655

APPENDIX 3

WATER CHEMISTRY DATA

CATION AND ANION COMPOSITION IN MG. PER LITRE

F-B1 OGNVRV

SAMPLE	NA	K	CA	MG	T	CAT	S04	HCO3	CL	F	NO3	T ANS	TDS	SI	PH	E COND	DRWF (1981)	H/NO 4745
1 BE68/1	111.90	8.70	48.80	36.32	205.72	46.12	344.98	100.65	0.92	37.27	529.94	735.66	25.48	7.60	0.0	1911	5437	
2 DSI	43.33	1.30	29.80	11.47	85.89	11.68	138.60	35.68	0.05	52.49	238.50	324.40	25.04	7.00	0.0	2711	6014	
3 GKB1/3	57.73	1.87	43.32	30.43	133.35	28.74	281.40	26.48	0.27	58.51	395.40	528.75	22.34	7.60	0.0	1611	5018	
4 GT1	10.63	2.16	10.28	4.98	28.04	6.26	59.14	4.58	0.20	17.39	87.97	115.62	20.45	6.30	0.0	1611	6056	
5 GT2	30.54	3.01	60.86	44.33	138.73	8.17	123.70	159.50	0.20	65.77	357.40	496.14	15.65	7.80	0.0	1611	4499	
6 HG1	24.23	3.51	32.31	18.24	78.29	23.09	84.10	17.39	0.25	84.14	208.97	287.26	23.05	6.70	0.0	2911	4431	
7 HG6	39.04	1.01	43.81	21.22	105.08	14.98	221.98	26.81	0.51	43.73	308.01	413.09	23.70	7.20	0.0	1211	6048	
8 HG9	95.86	3.03	48.19	34.69	181.76	28.06	385.10	63.40	0.64	39.48	516.68	698.45	26.94	8.10	0.0	0112	5948	
9 G34310-3	24.89	2.66	22.32	13.54	63.41	20.57	81.61	5.72	0.46	59.05	167.41	230.82	23.16	6.70	0.0	3616	1810.4481	
10 LT44/1	64.98	2.23	54.90	30.13	152.23	28.96	324.50	37.86	0.40	83.08	474.80	627.04	23.00	8.20	0.0	0312	6111	
11 LT63/1	34.26	4.43	24.24	11.52	74.44	22.08	110.80	13.06	0.50	43.07	189.51	263.96	20.84	6.90	0.0	0111	4753	
12 LT63/2	31.56	2.23	23.86	11.52	69.16	10.79	48.02	21.85	0.21	93.13	174.00	243.17	27.79	6.40	0.0	1911	4761	
13 M61/1	69.19	2.74	97.25	53.93	223.10	70.52	297.72	134.87	0.68	101.67	605.46	828.57	24.28	7.40	0.0	3616	2686	
14 NT79/1	33.62	2.67	23.22	15.31	74.81	31.93	114.70	20.03	0.23	44.75	217.64	292.46	23.21	7.30	0.0	3616	5995	
15 PA1-3	29.48	2.53	24.24	12.39	68.63	13.67	88.17	19.56	0.23	57.36	178.99	247.63	28.27	6.90	0.0	0511	0650.3915	
16 PA3	40.41	3.12	39.59	22.42	105.53	29.47	147.83	22.70	0.35	87.91	288.26	393.80	26.36	7.10	0.0	1911	4698	
17 PA7	43.18	2.74	26.17	13.54	85.63	7.50	164.91	18.41	0.46	49.84	241.12	326.75	24.03	7.10	0.0	0111	4711	
18 PA8	53.74	2.39	30.39	12.39	98.91	14.01	147.79	17.26	0.68	81.31	261.05	359.96	28.88	7.10	0.0	1911	4763	
19 RT1	69.93	1.51	121.66	46.93	240.02	93.81	386.08	126.01	0.62	22.53	629.05	869.08	13.02	7.50	0.0	3616	2446	
20 SP2	50.19	2.98	50.34	26.06	129.57	46.64	43.40	68.07	0.18	193.39	341.68	471.25	28.76	6.50	0.0	0111	4724	
21 SP5	59.49	2.85	39.21	15.56	88.11	17.43	23.56	227.68	26.95	0.52	41.43	320.14	437.57	24.02	7.30	0.0	4711	
22 SS2	43.16	3.89	75.75	37.87	160.67	53.65	344.03	48.96	0.58	11.28	458.50	619.17	21.97	7.50	0.0	0110	2701	
23 SS16	63.37	1.76	56.77	30.37	152.27	33.35	336.68	32.29	0.76	25.89	428.97	581.24	32.08	7.30	0.0	0110	2717	
24 SS20	56.39	2.80	67.37	43.40	169.95	53.62	341.58	74.49	0.45	2.92	473.06	643.02	18.30	7.20	0.0	0110	2735	
25 ZT4	34.14	4.13	28.07	15.92	82.26	29.24	47.60	30.44	0.22	128.23	235.73	317.99	25.59	7.00	0.0	0112	6006	

45100 - SAMPLES FOR DETERMINATION TR108.

102400 - NEW CONC. FOR DETERMINATION TR108.

SAMPLE	NA	K	CA	MG	T	CAT	S04	HCO3	CL	F	NO3	T ANS	T/BAL	NO3/CL	F-B1 OGNVRV
BE68/1	4.867	0.222	2.435	2.986	10.512	0.960	5.654	2.838	0.048	0.601	10.102	-1.557	0.121	F-B1 OGNVRV	
DS1	1.884	0.033	1.487	0.943	4.348	0.243	2.271	1.006	0.002	0.846	4.370	0.126	0.04	F-B1 OGNVRV	
GKB1/3	2.511	0.047	2.161	2.502	7.223	0.598	4.612	0.746	0.014	0.943	6.915	-1.441	1.126	F-B1 OGNVRV	
GT1	0.462	0.055	0.512	0.409	1.440	0.130	0.969	0.129	0.010	0.280	1.519	0.613	2.17	F-B1 OGNVRV	
GT2	1.328	0.076	3.036	3.645	8.088	0.170	2.027	4.499	0.010	1.061	7.768	-1.407	0.24	F-B1 OGNVRV	
HG1	1.054	0.089	1.612	1.500	4.256	0.480	1.378	0.490	0.013	1.357	3.720	-3.265	2.77	F-B1 OGNVRV	
HG6	1.698	0.025	2.186	1.745	5.655	0.311	3.638	0.756	0.026	0.705	5.438	-1.136	0.93	F-B1 OGNVRV	
HG9	4.169	0.077	2.404	2.652	9.505	0.584	6.311	1.787	0.033	0.636	9.354	-0.598	0.36	F-B1 OGNVRV	
G34310-3	1.082	0.068	1.113	1.113	3.378	0.428	1.337	0.161	0.024	0.952	2.903	-3.129	5.91	F-B1 OGNVRV	
LT44/1	2.826	0.057	2.739	2.477	8.101	0.602	5.318	1.067	0.021	1.340	8.350	1.056	1.26	F-B1 OGNVRV	
LT63/1	1.490	0.113	1.209	0.947	3.760	0.459	1.816	0.368	0.026	0.694	3.365	-2.493	1.89	F-B1 OGNVRV	
LT63/2	1.372	0.057	1.190	0.947	3.567	0.224	0.787	0.616	0.011	1.502	3.147	-2.749	2.44	F-B1 OGNVRV	
M61/1	3.009	0.070	4.852	4.435	12.367	1.468	4.879	3.803	0.035	1.640	11.827	-1.865	0.473	F-B1 OGNVRV	
NT79/1	1.462	0.068	1.158	1.259	3.948	0.789	1.879	0.564	0.012	0.721	3.966	0.118	1.28	F-B1 OGNVRV	
PA1-3	1.282	0.064	1.209	1.018	3.575	0.284	1.445	0.551	0.012	0.925	3.218	-2.201	1.68	F-B1 OGNVRV	
PA3	1.757	0.079	1.975	1.843	5.657	0.613	2.422	0.640	0.018	1.418	5.113	-2.828	2.22	F-B1 OGNVRV	
PA7	1.878	0.070	1.305	1.113	4.367	0.156	2.702	0.519	0.024	0.803	4.206	-0.940	1.55	F-B1 OGNVRV	
PAB	2.337	0.061	1.516	1.018	4.934	0.291	2.422	0.486	0.035	1.311	4.548	-2.181	2.17	F-B1 OGNVRV	
RT1	3.041	0.038	6.070	3.859	13.010	1.953	6.327	3.553	0.032	0.363	12.230	-2.635	0.10	F-B1 OGNVRV	
SP2	2.183	0.076	2.511	2.143	6.914	0.971	0.711	1.919	0.009	2.958	6.569	-1.656	1.54	F-B1 OGNVRV	
SP5	2.587	0.072	1.956	1.305	5.923	1.416	0.490	0.759	0.027	0.668	5.677	-1.261	0.28	F-B1 OGNVRV	
SS2	1.877	0.099	3.779	3.114	8.871	1.116	5.638	1.380	0.030	1.182	8.348	-2.214	0.13	F-B1 OGNVRV	
SS16	2.756	0.045	2.832	2.497	8.132	0.694	5.518	0.910	0.039	0.417	7.580	-2.460	0.46	F-B1 OGNVRV	
SS20	2.452	0.071	3.361	3.569	9.455	1.116	5.598	2.100	0.023	0.047	8.886	-2.330	0.02	F-B1 OGNVRV	
ZT4	1.485	0.105	1.400	1.309	4.300	0.608	0.780	0.858	0.011	2.068	4.327	0.153	2.41	F-B1 OGNVRV	

CHEMICAL ANALYSES OF GROUNDWATER SAMPLES FROM BOREHOLES PENETRATING BOTHNIVILLE FORMATION (X) AND SURFACE WATER SAMPLES FROM HAGET'S RIVER (A).

No. OF PLOTTED ON FIG. 6.

FIG. 6. CATION AND ANION COMPOSITION IN MG PER LITRE

SAMPLE	NA	K	CA	MG	T CAT	SO4	HCO3	CL	F	NO3	T ANS	TDS	SI	PH	E COND	DIST	H/H ₂ O	(1960)	g/1000
GN1	26	37.67	1.24	65.46	21.52	125.88	11.49	263.11		21.75	0.36	51.52	348.23	474.12	16.62	7.60	0.0	0.0	4157
H3	27	42.66	4.50	9.56	8.58	65.30	8.88	95.00	48.98	0.20	2.61	155.67	220.97	2.88	7.00	0.0	0.0	6022	
H4	28	107.12	7.71	12.40	20.43	147.66	20.53	212.30	96.47	1.10	1.32	331.72	479.38	1.42	7.40	0.0	0.0	6030	
JL2	29	20.29	2.33	17.29	9.51	49.41	9.12	39.64	25.47	0.27	62.85	137.35	186.77	20.35	6.60	0.0	1.11	4787	
PE1	30	15.19	1.88	11.98	6.51	35.55	0.00	76.60	7.47	0.24	8.94	93.25	128.81	24.26	6.70	0.0	0.0	5187	
SS17	31	17.69	3.31	10.28	5.26	36.54	5.00	50.38	21.82	0.14	11.64	88.98	125.52	26.52	6.30	0.0	0.0	2710	
G34314	32	34.28	1.31	44.48	23.00	103.07	7.97	239.13	24.07	0.29	35.54	307.00	410.07	16.76	7.70	0.0	0.0	4779	
G34315	33	36.30	0.84	44.57	20.62	102.33	11.96	252.58	13.31	0.36	3.14	281.35	383.68	19.50	7.10	0.0	0.0	3943	
ZNI	34	35.96	2.32	38.26	26.92	103.46	16.70	250.20	21.26	0.22	27.71	316.09	419.55	15.51	8.10	0.0	0.0	6064	

45.00
SABs. max. limit

MILLI-EQUIVALENTS PER LITRE AND IONIC BALANCE

SAMPLE	NA	K	CA	MG	T CAT	SO4	HCO3	CL	F	NO3	T ANS	I/BAL	NO3/CL	F-B1 OGNVRV
GN1	1.638	0.031	3.266	1.769	6.706	0.239	4.312	0.613	0.018	0.831	6.014	-3.462	1.36	F-B1 OGNVRV
H3	1.858	0.115	0.477	0.705	3.153	0.184	1.557	1.381	0.010	0.042	3.175	0.143	0.03	F-B1 OGNVRV
H4	4.659	0.197	0.618	1.680	7.155	0.427	3.479	2.729	0.057	0.021	6.706	-2.133	0.01	F-B1 OGNVRV
JL2	0.882	0.059	0.862	0.782	2.587	0.189	0.649	0.718	0.014	1.013	2.585	-0.007	1.14	F-B1 OGNVRV
PE1	0.660	0.048	0.597	0.535	1.842	0.000	1.255	0.210	0.012	0.144	1.622	-1.663	0.69	F-B1 OGNVRV
SS17	0.769	0.084	0.512	0.432	1.799	0.104	0.825	0.615	0.007	0.187	1.740	-0.445	0.30	F-B1 OGNVRV
G34314	1.491	0.033	2.219	1.891	5.635	0.165	3.919	0.678	0.015	0.573	5.352	-1.494	0.85	F-B1 OGNVRV
G34315	1.579	0.021	2.224	1.695	5.520	0.249	4.139	0.375	0.018	0.050	4.833	-3.784	0.13	F-B1 OGNVRV
ZNI	1.564	0.059	1.909	2.213	5.746	0.347	4.100	0.599	0.011	0.446	5.506	-1.251	0.74	F-B1 OGNVRV

MILLI-EQUIVALENTS PER LITRE OF CATIONS AND ANIONS
EXPRESSED AS PERCENTAGES OF THE TOTAL CATIONS AND THE TOTAL ANIONS
AND VARIOUS RATIOS (NA=NA+K CL=CL+NO3+F)

SAMPLE	NA	CA	MG	SO4	HCO3	CL	CL/HCO3	MG/CA	NA/CA	NA/CL	NA/CAT	ALK/CAT	SAR	THARD	LAT.	LONG.	CO-ORDINATES	OGNVRV
GN1	24.90	48.70	26.38	3.97	71.69	24.32	0.14	0.54	0.50	2.67	0.24	0.75	1.03	252.00	03°10'	14°25'	OGNVRV	
H3	62.49	15.12	22.37	5.82	49.02	45.15	0.88	1.47	3.89	1.34	0.58	0.37	2.41	59.17	15°18'	13°29'	OGNVRV	
H4	67.87	8.64	23.47	6.37	51.88	41.74	0.78	2.71	7.53	1.71	0.65	0.32	4.34	115.03	15°47'	11°05'	OGNVRV	
JL2	36.41	33.34	30.23	7.34	25.12	67.53	1.10	1.02	1.22	0.34	0.63	0.97	82.30	05°58'	15°35'	OGNVRV		
PE1	38.48	32.45	29.06	0.00	77.35	22.64	0.16	0.89	1.10	3.13	0.35	0.61	0.87	56.70	05°52'	14°03'	OGNVRV	
SS17	47.46	28.50	24.03	5.98	47.44	46.57	0.74	0.84	1.50	1.25	0.42	0.52	1.11	47.31	07°55'	17°07'	OGNVRV	
G34314	27.05	39.38	33.56	3.10	73.22	23.67	0.17	0.85	0.67	2.19	0.26	0.72	1.04	205.71	08°43'	20°03'	OGNVRV	
G34315	28.99	40.28	30.71	5.15	85.64	9.20	0.09	0.76	0.70	4.20	0.28	0.71	1.12	196.14	07°20'	20°08'	OGNVRV	
ZNI	28.25	33.22	38.52	6.31	74.47	19.21	0.14	1.15	0.81	2.60	0.27	0.71	1.08	206.31	18°36'	20°45'	OGNVRV	

NO. ON
FIGS.
CHEMICAL ANALYSES OF GROUNDWATER SAMPLES
PLOTTED ON F/G.

CATION AND ANION COMPOSITION IN MG PER LITRE

SAMPLE	NA	K	CA	MG	T CAT	SO4	HCO3	CL	F	NO3	T ANS	TDS	SI	PH	E COND	PAR	H. no. Y
54 KL2	60.21	4.39	61.56	30.68	156.83	22.08	333.59	39.33	0.43	55.28	450.66	667.50	16.61	7.90	0.0	0.0	814.965
55 MN2	245.01	11.28	64.00	71.00	391.29	10.31	633.60	19.87	1.83	24.74	967.65	1358.94	30.72	7.80	0.0	25.09	814.965
56 MN12	68.53	4.95	105.43	59.81	238.72	65.52	371.16	180.92	0.39	51.52	660.51	908.23	12.24	7.70	0.0	36.91	814.2638
57 NT86/1	135.44	4.93	75.74	43.05	259.15	72.10	450.80	120.75	0.80	13.14	657.59	916.75	16.50	7.80	0.0	16.11	814.4907
58 PN1	55.31	3.41	99.54	54.77	213.02	58.22	374.88	119.56	0.69	36.29	589.64	802.67	14.12	7.60	0.0	34.91	814.2662
59 PN9	83.35	6.31	66.25	44.67	200.58	15.58	426.26	76.58	0.44	42.40	561.26	761.84	20.74	7.60	0.0	13.11	814.4892
60 PN10-1	48.44	5.09	73.89	41.74	169.16	23.90	357.47	35.09	0.14	79.01	495.61	664.77	13.32	7.90	0.0	14.10	814.3867
61 PN10-2	51.60	4.45	73.75	39.19	168.99	27.78	368.37	28.26	0.31	67.19	491.91	660.90	17.42	7.70	0.0	13.11	814.4884
62 SK1	80.53	5.70	48.19	36.66	171.08	43.43	286.80	74.03	0.82	61.83	466.91	637.99	13.13	8.20	0.0	0.0	814.5971
+ NT 843 (59).										1.50	45.00						

MILLI-EQUIVALENTS PER LITRE AND IONIC BALANCE

SAMPLE	NA	K	CA	MG	T CAT	SO4	HCO3	CL	F	NO3	T ANS	1/BAL	NO3/CL	F-B1	OGNVRV
KL2	2.619	0.112	3.071	2.523	8.326	0.459	5.466	1.109	0.022	0.891	7.949	-1.638	0.80		
MN2	10.657	0.288	3.193	5.039	19.979	2.275	10.384	5.588	0.096	0.392	18.744	-3.109	0.07	F-B1	OGNVRV
MN12	2.981	0.126	5.260	4.918	13.287	1.364	6.083	5.101	0.020	0.831	13.401	0.361	0.16	F-B1	OGNVRV
NT86/1	5.891	0.126	3.779	3.540	13.337	1.501	7.388	3.405	0.042	0.212	12.549	-2.619	0.06	F-B1	OGNVRV
PN1	2.405	0.087	4.967	4.504	11.964	1.212	6.144	3.371	0.016	0.585	11.349	-2.176	0.17	F-B1	OGNVRV
PN9	3.625	0.161	3.305	3.673	10.766	0.324	6.986	2.159	0.023	0.684	10.177	-2.229	0.32	F-B1	OGNVRV
PN10-1	2.244	0.130	3.687	3.432	9.357	0.497	5.058	0.989	0.007	1.274	8.627	-3.035	1.21	F-B1	OGNVRV
PN10-2	2.244	0.113	3.680	3.222	9.261	0.578	6.037	0.796	0.016	1.083	8.513	-3.138	1.36	F-B1	OGNVRV
SK1	3.503	0.145	2.404	3.014	9.068	0.904	4.700	2.087	0.043	0.997	8.733	-1.386	0.48	F-B1	OGNVRV

MILLI-EQUIVALENTS PER LITRE OF CATIONS AND ANIONS
EXPRESSED AS PERCENTAGES OF THE TOTAL CATIONS AND THE TOTAL ANIONS
AND VARIOUS RATIOS (NA=NA+K CL=CL+NO3+F)

SAMPLE	NA	CA	MG	SO4	HCO3	CL	CL/HCO3	MG/CA	NA/CA	NA/CL	NA/CAT	ALK/CAT	SAR	THARD	as G.C.O.	LAT.	Long.	OGNVRV
KL2	32.80	36.89	30.30	5.78	68.76	25.45	0.20	0.82	0.85	2.36	0.31	0.67	1.56	279.96	27°53'			
MN2	54.78	15.98	29.22	12.14	55.40	32.45	0.53	1.82	3.33	1.90	0.53	0.45	5.01	451.97	18°33'	27°46'	OGNVRV	
MN12	23.38	39.59	37.01	10.17	45.39	44.42	0.83	0.93	0.56	0.58	0.22	0.76	1.32	509.37	08°57'	10°55'	OGNVRV	
NT86/1	45.11	28.33	26.54	11.96	58.87	29.16	0.46	0.93	1.55	1.73	0.44	0.54	3.07	366.27	09°35'	11°27'	OGNVRV	
PN1	20.83	41.51	37.64	10.67	54.13	35.18	0.54	0.90	0.48	0.71	0.20	0.79	1.10	473.92	16°47'	26°42'	OGNVRV	
PN9	35.17	30.70	34.12	3.18	68.64	28.16	0.30	1.11	1.09	1.67	0.33	0.64	1.94	349.24	16°46'	17°19'	OGNVRV	
PN10-1	23.91	39.40	36.68	5.76	67.90	26.32	0.16	0.93	0.57	2.12	0.22	0.76	1.11	356.26	12°28'	25°56'	OGNVRV	
PN10-2	25.46	39.73	34.79	6.79	70.92	22.28	0.13	0.87	0.60	2.81	0.24	0.74	1.20	345.42	11°26'	23°52'	OGNVRV	
SK1	40.23	26.51	33.24	10.35	53.82	35.82	0.44	1.25	1.45	1.67	0.38	0.59	2.12	271.18	14°50'	30°49'	OGNVRV	

CHEMICAL ANALYSES OF GROUNDWATER SAMPLES

CATION AND ANION COMPOSITION IN MG PER LITRE

SAMPLE	NA	K	CA	MG	T CAT	SO ₄	HC0 ₃	CL	F	NO ₃	T ANS	TDS	SI	PH	E COND	DATE	TIME	H ₂ O ₂	ESR
63 G34310-1	22.11	2.71	26.56	15.29	66.66	20.50	94.07	9.23	0.33	56.17	180.30	246.97	26.40	6.90	0.0	20/10	1015	3870	—
64 G34310-2	23.57	2.31	22.70	12.96	61.54	22.14	77.86	7.34	0.45	57.14	164.93	226.47	23.11	6.70	0.0	11/11	0945	4473	—
PA1-1	30.16	2.51	22.31	11.24	66.22	10.66	80.15	17.17	0.32	61.08	169.38	235.60	28.29	6.70	0.0	04/11	0722	5117	—
66 PA1-2	28.80	2.27	23.08	11.53	65.68	14.53	82.85	19.02	0.24	56.43	173.07	238.75	28.54	6.70	0.0	04/11	1930	3127	—
67 G34307-1	4.00	—	55.56	29.69	136.08	33.06	257.81	32.94	0.55	44.58	388.94	525.92	15.77	7.80	0.0	01/10	1310	2714	—
68 G34307-2	4.627	2.97	58.39	28.47	136.10	28.24	260.00	34.18	0.32	59.31	382.05	518.15	15.96	7.60	0.0	02/10	1045	3888	—
69 G34307-3	47.71	3.23	58.79	28.47	138.20	25.81	264.03	34.06	0.36	58.03	382.29	520.49	15.83	7.80	0.0	27/10	1610	3896	—
70 G34307-4	46.99	3.24	57.59	28.78	137.00	26.72	262.83	33.95	0.36	60.11	383.97	520.97	15.59	7.70	0.0	29/10	0440	3901	—
71 G34307-5	48.93	3.69	30.91	27.88	111.41	25.95	211.40	33.76	0.36	59.36	330.83	442.24	17.43	8.00	0.0	03/11	1200	5161	—
— S.A.S.S. ^{Max. conc. for drinking water (Prop.)} —																			
— Max. conc. for drinking water (Prop.) —																			
— G34307-3 (9) Pa / 3 (15) —																			
— G34307-1 (49) —																			

MILLI-EQUIVALENTS PER LITRE AND IONIC BALANCE

SAMPLE	NA	K	CA	MG	T CAT	SO ₄	HC0 ₃	CL	F	NO ₃	T ANS	T/BAL	NO ₃ /CL	F-B1	OGNVRV
G34310-1	0.961	0.069	1.325	1.257	3.613	0.426	1.541	0.260	0.017	0.906	3.152	-2.670	3.48	F-B1	OGNVRV
G34310-2	1.025	0.059	1.132	1.065	3.282	0.460	1.276	0.206	0.023	0.921	2.889	-2.600	4.47	F-B1	OGNVRV
PA1-1	1.311	0.064	1.113	0.924	3.413	0.221	1.313	0.484	0.016	0.985	3.021	-2.555	2.04	F-B1	OGNVRV
PA1-2	1.252	0.058	1.151	0.948	3.410	0.302	1.357	0.536	0.012	0.910	3.119	-1.879	1.70	F-B1	OGNVRV
G34307-1	2.076	0.102	2.772	2.441	7.392	0.688	4.225	0.928	0.028	1.041	6.913	-2.243	1.12	F-B1	OGNVRV
G34307-2	2.012	0.075	2.913	2.341	7.343	0.587	4.261	0.963	0.016	0.956	6.786	-2.630	0.97	F-B1	OGNVRV
G34307-3	2.075	0.082	2.933	2.341	7.433	0.537	4.327	0.960	0.018	0.935	6.780	-3.084	0.98	F-B1	OGNVRV
G34307-4	2.044	0.082	2.893	2.366	7.387	0.556	4.307	0.957	0.018	0.969	6.810	-2.723	1.01	F-B1	OGNVRV
G34307-5	2.128	0.094	1.542	2.292	6.058	0.540	3.464	0.952	0.018	0.957	5.933	-0.627	1.02	F-B1	OGNVRV

MILLI-EQUIVALENTS PER LITRE OF CATIONS AND ANIONS
EXPRESSED AS PERCENTAGES OF THE TOTAL CATIONS AND THE TOTAL ANIONS
CL=CL+NO₃+F

SAMPLE	NA	CA	MG	SO ₄	HC0 ₃	CL	CL/HC0 ₃	MG/CA	NA/CA	NA/CL	NA/CAT	ALK/CAT SAR	THAI as CaCO ₃	F-B1	OGNVRV	
G34310-1	28.53	36.67	34.79	13.53	48.90	37.55	0.16	0.94	0.72	3.69	0.26	0.71	0.84	F-B1	OGNVRV	
G34310-2	33.03	34.50	32.46	15.95	44.16	39.88	0.16	0.94	0.90	4.95	0.31	0.66	0.97	F-B1	OGNVRV	
PA1-1	40.31	32.61	27.07	7.34	43.47	49.10	0.36	0.83	1.17	2.70	0.38	0.59	1.29	F-B1	OGNVRV	
PA1-2	38.43	33.76	27.80	9.69	43.52	46.77	0.39	0.82	1.08	2.33	0.36	0.61	1.22	F-B1	OGNVRV	
G34307-1	29.46	37.50	33.02	9.95	61.12	28.92	0.21	0.88	0.74	2.23	0.28	0.70	1.28	F-B1	OGNVRV	
G34307-2	28.44	39.67	31.88	8.66	62.78	28.54	0.22	0.80	0.69	2.08	0.27	0.71	1.24	F-B1	OGNVRV	
G34307-3	29.03	39.46	31.49	7.92	63.82	28.25	0.22	0.79	0.70	2.16	0.27	0.70	1.27	F-B1	OGNVRV	
G34307-4	28.79	39.17	32.03	8.16	63.25	28.57	0.22	0.81	0.70	2.13	0.27	0.71	1.26	F-B1	OGNVRV	
G34307-5	36.69	25.46	37.84	9.10	58.39	32.50	0.27	1.48	1.37	2.23	0.35	0.63	1.53	191.90	F-B1	OGNVRV

ANALOGUE ANDESTITE FORMATION

MILLI-EQUIVALENTS PER LITRE OF CATIONS AND ANIONS
EXPRESSED AS PERCENTAGES OF THE TOTAL CATIONS AND THE TOTAL ANIONS
(NA=NA+K CL=CL+NO₃+F)

SAMPLE	NA	CA	HG	SO ₄	HCO ₃	CL	CL/HC03	MG/CA	NA/CA	NA/CL	NA/CAT	ALK/CAT	SAR	THARD	LAT	LONG	F-B1	OGNVRV	
BT2	25.42	41.03	33.54	5.12	69.02	25.84	0.11	0.81	0.58	3.41	0.24	0.74	1.07	277.83	06°58'	21°55'	F-B1	OGNVRV	
BT7	27.30	39.58	33.10	4.50	68.36	27.12	0.11	0.83	0.67	3.55	0.26	0.72	1.26	291.21	05°55'	25°05'	F-B1	OGNVRV	
DK1	36.81	34.18	28.99	8.02	66.22	25.74	0.22	0.84	1.01	2.43	0.34	0.63	1.84	278.87	17'36"	21'45"	F-B1	OGNVRV	
EM1	24.50	37.50	37.99	2.06	35.67	61.45	1.32	1.01	0.64	0.54	0.24	0.75	0.99	243.30	05°13'	15'38"	F-B1	OGNVRV	
LT9/2/1	17.24	52.25	41.53	30.70	12.83	46.67	40.49	0.59	0.73	0.64	1.01	0.26	0.72	1.53	432.46	05°05"	27'15"	F-B1	OGNVRV
LT11/2	24.57	41.51	33.91	30.49	3.53	25.14	71.31	1.38	0.58	0.32	0.50	0.16	0.82	1.17	821.58	03°18"	20'52"	F-B1	OGNVRV
LT19/2	30.86	38.35	30.77	6.30	61.26	33.40	0.29	0.81	0.57	1.45	0.23	0.75	0.99	244.08	05°35"	24'06"	F-B1	OGNVRV	
LT20/1	36.18	35.31	28.49	5.49	53.25	41.24	0.22	0.80	0.79	4.88	0.30	0.69	1.50	289.33	06°32"	22'16"	F-B1	OGNVRV	
39/1	42.70	26.59	30.70	13.51	35.78	50.70	1.11	1.15	1.58	1.07	0.42	0.57	2.28	130.20	05°28"	05°22"	F-B1	OGNVRV	
49/3	21.65	51.91	26.42	6.42	66.01	27.56	0.13	0.50	0.40	2.49	0.21	0.78	0.81	241.24	09°52"	01°25"	F-B1	OGNVRV	
61/6	24.47	45.10	30.42	5.11	42.09	52.79	0.94	0.67	0.52	0.61	0.23	0.75	1.39	226.07	07°17"	21'40"	F-B1	OGNVRV	
MN10	24.16	37.53	38.30	3.41	74.20	22.38	0.13	1.02	0.62	2.53	0.23	0.75	0.94	493.16	14°22"	23'25"	F-B1	OGNVRV	
MN15	24.66	39.57	35.76	4.22	72.21	23.56	0.13	0.90	0.59	2.56	0.23	0.75	0.94	228.13	08°15"	24'33"	F-B1	OGNVRV	
G34307-6	31.07	38.10	30.81	7.49	65.57	26.93	0.20	0.80	0.78	2.31	0.29	0.75	1.14	335.14	05°35"	05°44"	F-B1	OGNVRV	
NT86/3	39.01	33.26	27.72	6.38	74.46	19.14	0.17	0.83	1.09	2.89	0.36	0.68	1.37	253.66	07°05"	10°04"	F-B1	OGNVRV	
RT11	28.25	42.30	29.43	2.13	90.07	7.79	0.04	0.69	0.66	7.69	0.27	0.71	1.81	232.10	15°25"	28°35"	F-B1	OGNVRV	
VT6	23.81	41.54	34.63	0.56	61.06	38.37	0.24	0.83	0.55	1.66	0.23	0.76	0.97	255.25	11°14"	22'22"	F-B1	OGNVRV	
ZN2	26.53	43.23	30.23	0.32	64.03	35.64	0.20	0.69	0.60	2.13	0.25	0.73	1.18	281.96	18°07"	21°20"	F-B1	OGNVRV	

Geographical coordinates

Chemical Analyses of Groundwater Samples
Plotted on Fig.

From Blanrode Andesite Formation.

CATION AND ANION COMPOSITION IN MG PER LITRE

SAMPLE	NA	K	CA	MG	T CAT	SO ₄	HC0 ₃	CL	F	NO ₃	T ANS	TDS	SI	PH	E COND	DATE	H/mf
35 BT2	41.35	3.68	61.22	30.37	136.61	17.04	291.66	18.68	0.41	76.97	404.76	541.38	18.71	7.60	0.0	10/11/61	8141
36 BT7	49.48	1.31	63.51	32.23	146.52	16.16	311.62	21.48	0.38	86.84	436.48	583.01	20.98	7.50	0.0	10/11/61	2689
37 DK1	6.76	60.43	31.10	168.95	32.67	342.30	44.72	0.32	55.95	475.96	644.92	13.69	8.10	0.0	10/11/61	4969	
38 EM1	35.62	1.11	48.41	29.75	114.88	8.32	131.55	101.25	0.33	52.18	293.63	408.52	19.90	6.90	0.0	10/11/61	4800
39 KN2/1	73.49	4.83	99.58	44.67	222.56	69.73	322.05	111.87	0.56	86.49	590.70	813.27	18.57	7.60	0.0	10/11/61	4876
40 LT9/1	77.25	2.43	207.77	73.58	361.03	32.46	293.50	235.92	0.32	432.30	994.50	1355.53	18.58	7.70	0.0	10/11/61	4810
41 LT16/2	35.62	1.56	53.80	26.67	117.64	15.13	221.09	37.79	0.25	75.59	329.85	447.50	17.49	7.60	0.0	10/11/61	4842
42 LT9/2	58.77	0.97	64.29	31.30	155.33	23.49	343.33	18.57	0.31	68.25	453.95	609.28	17.19	7.70	0.0	10/11/61	4839
43 LT20/1	32.91	1.72	28.86	14.13	77.61	10.16	124.98	16.35	0.37	98.56	220.42	298.04	20.78	6.90	0.0	10/11/61	4745
44 LT39/1	81.65	1.68	44.86	31.43	159.61	53.64	180.40	117.45	0.59	53.03	404.81	564.43	13.97	7.80	0.0	10/11/61	5945
45 LT49/3	28.21	0.85	60.00	18.53	107.59	17.61	229.79	17.79	0.22	66.31	331.35	438.94	16.00	7.20	0.0	10/11/61	4826
46 PA61/6	71.24	3.69	117.94	4.828	241.14	31.01	324.49	178.28	0.31	100.79	634.88	876.03	18.56	7.60	0.0	10/11/61	4923
47 MN10	32.65	1.27	45.22	28.00	107.13	9.13	251.77	19.86	0.39	41.16	322.31	429.45	17.01	7.30	0.0	10/11/61	2620
48 MN15	48.33	3.54	70.51	38.66	161.04	16.69	362.44	29.06	0.56	67.54	476.29	637.33	15.36	7.50	0.0	10/11/61	2610
49 G34307-6	50.30	3.81	56.17	27.56	137.83	25.51	283.40	33.51	0.40	58.38	401.20	539.04	17.43	8.20	0.0	10/11/61	5161
50 NT86/34	63.53	7.98	50.70	25.64	147.85	22.08	327.00	33.87	0.41	24.87	408.23	556.08	12.99	8.00	0.0	10/11/61	6103
51 RT11	52.31	0.95	69.00	29.13	151.39	7.75	415.60	10.49	0.39	16.90	451.13	602.52	15.66	7.60	0.0	10/11/61	4950
52 VT6	35.79	1.48	55.75	28.20	121.22	1.71	235.00	33.07	0.23	91.49	361.50	482.72	17.30	7.90	0.0	10/11/61	6080
53 ZN2	45.82	1.63	66.45	28.20	142.10	11.10	277.60	33.06	0.22	98.49	410.47	552.57	17.02	8.00	0.0	10/11/61	6072

* shows with same conc.

102.00 = max. limits for drinking water TDS

45.00 = max. limit.

102.00 = max. limits for drinking water TDS

HILLI-EQUIVALENTS PER LITRE AND IONIC BALANCE

SAMPLE	NA	K	CA	MG	T CAT	SO ₄	HC0 ₃	CL	F	NO ₃	T ANS	1/10L	NO ₃ /CL	F-B1 OGVRV	F-B1 OGVRV
BT2	1.798	0.094	3.054	2.497	7.445	0.354	4.780	0.526	0.021	1.241	6.925	-2.432	2.36	F-B1 OGVRV	F-B1 OGVRV
BT7	2.152	-0.033	3.169	2.650	8.005	0.336	5.107	0.605	0.019	1.400	7.470	-2.407	2.31	F-B1 OGVRV	F-B1 OGVRV
DK1	0.074	0.172	3.015	2.557	8.820	0.680	5.610	1.261	0.016	0.902	8.470	-1.468	0.72	F-B1 OGVRV	F-B1 OGVRV
EM1	1.549	0.028	2.415	2.446	6.440	0.173	2.156	2.855	0.017	0.841	6.043	-1.980	0.29	F-B1 OGVRV	F-B1 OGVRV
KN2/1	3.196	0.123	4.969	3.673	1.451	0.278	3.154	0.029	1.395	11.309	-2.319	0.44	F-B1 OGVRV	F-B1 OGVRV	
LT9/1	3.360	0.062	1.0367	6.0367	1.9.841	0.675	4.810	6.652	0.016	6.973	1.9.129	-1.767	1.05	F-B1 OGVRV	F-B1 OGVRV
LT16/2	1.549	0.039	2.684	2.193	6.467	0.315	3.623	1.065	0.013	0.896	5.914	-2.790	0.84	F-B1 OGVRV	F-B1 OGVRV
LT19/2	2.556	0.024	3.208	2.574	8.363	0.489	5.627	0.523	0.016	1.101	7.757	-2.673	2.11	F-B1 OGVRV	F-B1 OGVRV
LT20/1	1.431	0.043	1.440	1.162	4.077	0.211	2.048	0.461	0.019	1.106	3.846	-1.392	2.45	F-B1 OGVRV	F-B1 OGVRV
LT39/1	3.551	0.042	2.238	2.584	8.418	1.116	2.956	3.303	0.031	0.855	8.263	-0.658	0.26	F-B1 OGVRV	F-B1 OGVRV
LT49/3	1.227	0.021	2.993	1.523	5.766	0.366	3.766	0.491	0.011	1.069	5.705	-0.315	2.18	F-B1 OGVRV	F-B1 OGVRV
MA61/6	3.098	0.094	5.885	3.970	1.3.049	0.645	5.318	5.027	0.016	1.625	12.633	-1.374	0.32	F-B1 OGVRV	F-B1 OGVRV
MN10	1.420	0.032	2.256	2.302	6.011	0.190	4.126	0.560	0.020	0.664	5.561	-2.339	1.19	F-B1 OGVRV	F-B1 OGVRV
MN15	2.102	0.020	3.518	3.179	8.890	0.347	5.940	0.819	0.029	1.089	8.226	-2.838	1.33	F-B1 OGVRV	F-B1 OGVRV
G34307-6	2.188	0.097	2.802	2.266	7.354	0.531	4.644	0.944	0.021	0.941	7.083	-1.253	1.05	F-B1 OGVRV	F-B1 OGVRV
NTB6/3	2.763	0.204	2.529	2.108	7.606	0.459	5.359	0.955	0.021	0.401	7.197	-1.875	0.42	F-B1 OGVRV	F-B1 OGVRV
RT11	2.275	0.024	3.443	2.395	8.138	0.161	6.811	0.295	0.020	0.272	7.562	-2.576	0.92	F-B1 OGVRV	F-B1 OGVRV
VI6	1.556	0.037	2.781	2.319	6.695	0.035	3.851	0.932	0.012	1.475	6.307	-1.899	1.58	F-B1 OGVRV	F-B1 OGVRV
ZN2	1.993	0.041	3.315	2.319	7.669	0.022	4.549	0.932	0.011	1.588	7.105	-2.606	1.70	F-B1 OGVRV	F-B1 OGVRV

APPENDIX 4

PUMPING TEST DATA

HAMBURG 82 HO

G34310 Pumped borehole

Start 0700, 11/11/1981 Shutdown 1700, 11/11/1981

R.W.L. 2,673 m		Pump intake 22 m		Pumping period 600 min.		
Step	1	2	3	4	5	6
Av. Q 1/s	4,29	8,40	9,87	13,31	16,84	14,54
Mins. elapsed from start of step						
Drawdown (m)						
1	1,547	5,177	7,197	10,322	14,229	14,477
2	1,749	5,382	7,277	10,702	14,373	-
3	1,851	5,489	7,297	-	14,523	-
4	1,930	5,577	7,297	11,002	14,839	15,057
5	1,977	5,597	7,322	11,060	15,027	15,127
6	2,019	5,656	7,387	11,112	15,207	-
7	2,049	5,693	7,427	11,182	-	-
8	2,068	5,757	7,513	11,222	15,182	-
9	2,078	5,769	7,490	11,247	16,102	-
10	2,099	5,822	7,444	11,280		14,987
12	2,116	5,980	7,577	11,299		14,982
14	2,145	6,038	7,705	11,337		14,967
16	2,169	6,074	7,769	11,355	Turbulence	14,885
18	2,537	5,922	7,807	11,400	too high	14,942
20	2,593	6,130	7,797	11,439	for	14,982
25	2,667	6,065	7,777	11,499	readings	15,022
30	2,705	6,083	7,842	11,590		15,017
35	2,730	6,075	7,832	11,612		15,027
40	2,764	6,100	7,947	11,654		15,037
45	2,779	6,248	7,995	11,667		15,027
50	2,799	6,387	8,032	11,752		15,067
55	2,824	6,277	8,030	11,764		15,100
60	2,864	6,257	7,970	11,810		15,090
70	2,877	6,214	8,001	11,842	17,540	14,860
80	2,875	6,317	7,757	11,902	-	14,930
90	2,896	6,467	7,675	11,962	-	-
100	2,930	5,981	7,517	11,980	17,800	14,860

G34310 - Measurements of discharge rates and water quality

Time since start of pumping (min)	Discharge rate (l/s)	pH	T.D.S. mg/l @ 25°C
22	4,34	7,1	220
48	4,26		
85	4,26		219
122	8,35		
148	8,41	6,9	221
175	8,31		
193	8,51		
220	10,00		
242	10,00	6,9	221
285	9,60		
322	13,33		
352	13,33	7,05	230
392	13,26		
425	16,74		
450	17,14	7,05	230
475	16,74		
495	16,74		
522	14,57		
547	14,46	6,90	225
575	14,55		
595	14,57		

Well head water temperature is 20,5 to 21,5°C

HAMBURG 82 HO

G34310 Pumped borehole, G34311 + G34312 Observation boreholes

Start 0900, 17/11/1981 Shutdown 0755, 22/11/81. Pumping period 7 135 min.
 Pump intake 17,7 m

	G34310		G34311	G34312
	R.W.L 2,680 m		R.W.L 2,996 m	R.W.L 2,865 m
			r = 30 m	r = 30 m
Time elapsed				
since pumping				
started, t	s	corrected	s	s
(min.)	(m)	drawdown s'(m)	(m)	(m)
2	-		no change	0,460
3	8,840	5,66	"	-
4	9,167	5,87	"	1,452
5	9,402	6,02	"	-
6	9,582	6,13	"	1,944
7	9,707	6,21	"	-
8	9,802	6,27	"	2,095
9	9,903	6,34	"	-
10	9,982	6,39	"	2,195
12	10,145	6,49	0,002	2,375 (13 min)
15	10,300	6,59	0,005	2,400
20	10,452	6,69	0,015	2,605
25	-	6,83	0,033	2,805 (28)
30	10,676	6,83	0,046	2,975 (36)
40	10,875	6,96	0,078	3,065 (43)
50	10,984	7,03	0,120	3,160
60	11,100	7,10	0,157	3,235 (58)
70	11,195	7,16	0,191	3,305 (65)
80	11,240	7,19	0,239	3,420
90	-	-	0,281	3,465 (88)
100	11,440	7,32	0,318	3,495

G34310

R.W.L 2,680 m

G34311

R.W.L 2,996 m

G34312

r = 30 m

r = 30 m

Time elapsed

since pumping

started, t
(min.)s
(m)corrected
drawdown
s' (m)s
(m)s
(m)

120	11,568	7,40	0,420	3,597
150	11,708	7,49	0,554	3,715
180	-	-	0,670	3,805 -
200	11,90	7,64	-	-
220	-	-	0,838	3,920
260(250)	12,086	7,74	0,960	4,010
300	12,295	7,87	1,097	4,075
350	-	-	1,255	4,145
400	12,440	7,96	1,385	4,215
450	-	-	1,499	4,285
500	12,655	8,10	1,604	4,390
600	12,775	8,18	1,766	4,480
700	12,890	8,25	1,886	4,530
800	13,040	8,35	1,976	4,620
900	13,125	8,40	2,017	4,670
1 000	-	-	2,057	4,705
1 200	13,250	8,48	2,134	4,805
1 500	13,473	8,62	2,219	4,975 -
1 800	13,698	8,77	2,258	5,110
2 200 (2 400)	13,892	8,89	2,310	5,215
2 600 (2 730)	14,160	9,06	2,340	5,255
3 000	14,303	9,15	2,382	5,355
3 500 (3 400)	14,395	9,21	2,419	5,422
4 000 (4 200)	14,805	9,48	2,464	5,612
4 500	-	-	2,476	5,671
5 000	14,420	9,23	2,511	5,641
6 000	14,300	9,15	2,531	5,701
6 500 (6300)	14,120	9,04	2,546	5,696
7 000	-	-	2,556	5,396 -
7 135	13,320	8,52	2,577	5,290

G34310 - Discharge and water quality measurements

t (min)	Q l/sec	pH	T.D.S. mg/l @ 25°C
65	13,14		
240	13,04		
370	12,97	6,90	235
485	12,97	7,05	232
660	12,90		
1 260	12,74	7,15	217
1 515	12,81		
1 910	12,74	6,95	224
2 220	12,67		
2 700	12,77	7,05	221
3 170	12,67	7,00	228
3 465	12,67	6,90	221
4 160	12,59		
4 320	12,63	7,10	221
4 680	12,63		
4 990	12,54	7,10	226
5 610	12,41		220
5 990	12,31		
6 120	12,24		
6 315	12,20		219
7 080	10,91*		

*pump cylinder screen partially blocked

Mean discharge 12,65 l sec, 1 093 m³/d

Well head water temperature 19 to 21°C

Observation boreholes

G34313		Hg 1		Hg 2	
R.W.L. 3,610 m		R.W.L. 7,935 m		R.W.L. 4,415 m	
r = 480 m		r = 315 m		r = 118 m	
t (min)	s (m)	t (min)	s (m)	t (min)	s (m)
1 295	0,005	26	0,005	17	
2 700	+ 0,025	108	0,013	to	fluctuation
4 200	0,047	158	0,010	7 135	of 0,005 m -
5 640	0,072	210	0,035		no response
7 135	0,055	260	0,013		to pumping
		307	0,020		of G34310
	No clear	410	0,023		
	response	510	0,025		
	to pumping	590	0,038		
	of G34310	710	0,050		
		820	0,065		
		905	0,055		
		1 195	0,040		
		1 510	0,060		
		1 810	0,040		
		2 740	0,060		
		3 005	0,060		
		3 410	0,060		
		4 190	0,085		
		4 410	0,088		
		4 840	0,070		

Recovery measurements

G34310			G34311			G34312		
t'	t/t'	s'	R.W.L. 2,680 m	R.W.L. 2,996 m	R.W.L. 2,865 m	F.D.P.W.L. 16,000 m	F.D.P.W.L. 5,573	F.D.P.W.L. 8,155
			Recovery from			Recovery from		
t'	t/t'	s'	F.D.P.W.L.	s'	F.D.P.W.L.	s'	F.D.P.W.L.	Recovery from
(min)		(m)	(m)	(m)	(m)	(m)	(m)	
5	1	428	8,835	7,165	2,566	0,007	3,910	4,245
7	1	020	probe jammed		2,566	0,007	3,690	4,465
10	715	-	-	2,566	0,007	3,490	4,665	
15	477	-		2,566	0,007	3,330	4,825	
25	286	-		2,566	0,007	3,205	4,950	
35	204	-		not measurable		3,020	5,135	
45	160	-		no change		2,900	5,255	
60	120	-		no change		2,885	5,270	
71,4	101	-		no change		2,885	5,270	
80	90	-		no change		2,885	5,270	
90	80	-		no change		2,820	5,335	
103	70	-		no change		2,710	5,445	
121	60	-		no change		2,635	5,520	
146	50	-		no change		2,510	5,645	
161	45	2,315	13,685	-	no change		2,460	5,695
182	40	2,230	13,770	-	no change		2,382	5,773
210	35	2,145	13,855	-	no change		2,287	5,868
245	30	2,055	13,945	-	no change		2,192	5,963
285	26	1,970	14,030	-	no change		2,112	6,043
340	22	1,865	14,135	-	no change		1,747	6,408
418	18	1,745	14,255	-	2,545	3,028	1,627	6,528
510	15	1,645	14,355	-	2,471	3,102	1,527	6,628
600	12,9	1,580	14,420	-	2,396	3,177	1,457	6,698
700	11,2	1,490	14,510	-	2,246	3,327	1,372	6,783
800	9,9	1,425	14,575	-	2,181	3,392	1,312	6,843
900	8,9	1,370	14,630	-	2,091	3,482	1,262	6,893
1 000	8,1	1,315	14,685	-	2,021	3,552	1,207	6,948
1 200	6,95	1,230	14,770	-	1,796	3,777	1,127	7,028
1 500	5,8	1,195	14,805	-	1,528	4,045	1,047	7,108
1 800	4,96	1,058	14,942	-	1,329	4,244	0,923	7,232
2 200	4,24	0,993	15,007	-	1,148	4,425	0,853	7,302
2 600	3,74	0,918	15,082	-	1,017	4,556	0,808	7,347
3 000	3,4	0,857	15,143	-	0,937	4,636	0,800	7,355
4 000	2,8	0,747	15,253	-	0,794	4,779	0,707	7,448
5 000	2,4	-	-	-	0,707	4,866	0,642	7,513
5 800	2,23	0,592	15,408	-	0,584	4,989	0,552	7,603

PALACHOEMA 64 HO

Pa 1 (B/no 136769/1) Pumping borehole

Pa 2 (B/no 136770/6) Observation borehole

Start 0705, 04/11/1981 Shutdown 0705, 05/11/1981

Rest water levels: Pa 1 - 4,512 m; Pa 2 - 4,625 m

Pump intake 35 m Pumping period 1 440 min, $r = 71,5$ m

Discharge and water quality measurements - Pa 1

Time since pumping started, t (min)	Discharge rate (1/s)	T.D.S. mg/l @ 25°C
---------------------------------------	----------------------	--------------------

17	12,87	235
40	12,87	
62	12,87	
110	12,75	235
140	12,87	
197	12,87	
265	12,87	232
315	12,87	234
375	12,64	237
435	12,64	234
545	12,64	238
625	12,64	234
685	12,64	239
775	12,64	232
1 335	12,64	240
1 425	12,64	237

Well head water temperature is 19,2 to 21°C

Mean discharge 12,69 1/s - 1 110 m^3/d

No water level measurements could be made in Pa 1

Pa 2 - Drawdown data

t (min)	drawdown (s) (m)	t (min)	s (m)
7	0,002	200	0,955
8	0,012	250	1,030
10	0,037	300	1,103
12	0,068	350	1,180
15	0,120	405	1,225
22	0,229	450	1,287
30	0,333	550	1,367
40	0,438	600	1,407
50	0,518	700	1,487
60	0,578	800	1,558
70	0,628	1 000	1,684
80	0,671	1 200	1,799
90	0,707	1 440	1,910
100	0,747		
120	0,801		
135	0,845		
170	0,927		

Recovery measurements

t' (min)	t'/t'	Pa 1	Pa 2	Recovery from F.D.P.W.L. (m)
		R.W.L. 4,512 m	R.W.L. 4,625 m	
		F.D.P.W.L.	F.D.P.W.L. 6,535 m	
		not measurable		
s'		s'		
(m)		(m)		
20	73	3,198	1,795	0,115
30	49	2,822	1,675	0,235
40	37	2,627	1,585	0,325
50	29,8	2,469	1,535	0,375
60	25	-	1,485	0,425
70	21,6	-	1,440	0,470
80	19	-	1,395	0,515
100	15,4	2,013	1,345	0,565
120	13	-	1,295	0,615
130	12,1	1,771	-	-
150	10,6	-	1,255	0,655
170	9,5	1,683	-	-
200	8,2	1,590	1,180	0,730
250	6,8	1,483	1,106	0,804
300	5,8	1,350	1,075	0,835
400	4,6	-	0,965	0,945
450	4,2	1,126	0,900	1,010
500	3,9	1,145	0,833	1,077
600	3,4	1,055	0,815	1,095
700	3,06	1,165	0,765	1,145
800	2,8	0,930	0,718	1,192
1 000	2,44	0,830	0,640	1,270
1 200	2,2	0,745	0,580	1,330
1 500	1,96	0,665	0,508	1,402
2 000	1,72	0,520	0,422	1,488
2 500	1,58	0,500	0,387	1,523
2 800	1,51	0,415	0,347	1,563
3 500	1,41	0,333	0,345	1,565
4 000	1,36	0,328	0,295	1,615
4 500	1,32	0,303	0,255	1,655

SCHWEIZER-RENEKE TOWNLANDS 62 HO

G34315 Pumped borehole

25/11/1981 Phase 1 Start 0825 Shutdown 1125 Pumping period 180 min.
 Phase 2 Start 1450 Shutdown 1720 Pumping period 150 min.
 R.W.L. prior to Phase 1: 6,282 m, R.W.L. prior to Phase 2 6,550 m.
 Pump intake at 64 m. Cascading water entering at 16 m made water level measurements taken during pumping unreliable.

DISCHARGE MEASUREMENTS

Phase 1	Time since pumping started (min)	Discharge (1/s)	Phase 2	t (min)	Discharge (1/s)
12	6,02		5	7,13	
30	5,95		15	6,21	
50	5,85		35	6,05	
70	5,81		65	5,90	
100	5,71		95	5,45*	
125	5,58		135	5,37	
165	5,54		148	5,29	

* engine revs decreased

Recovery measurements

Phase 1		Phase 2		
Residual drawdown s'	t/t'	Time since pumping stopped t' (min)	t/t'	s'
(m)				(m)
5,833	91	2	76	4,053
3,913	61	3	51	2,940
2,883	46	4	38,5	2,392
2,368	37	5		-
2,028	31	6	26,4	1,702
1,843	26,7	7	22,4	1,552
1,723	23,5	8	19,8	1,447
-	-	9	17,7	1,365
1,558	19	10	16	1,300
1,453	16	12	13,5	1,193
1,333	13	15	11	1,077
1,193	10	20	8,5	0,955
-	-	25	7	0,848
0,986	7	30	6	0,771
0,853	5,5	40	4,8	0,650
0,758	4,6	50	4	0,555
0,688	4	60	3,5	0,485
0,623	3,6	70	3,14	0,425
0,578	3,25	80	2,9	0,397
0,533	3	90	2,7	0,337
0,498	2,8	100	2,5	0,303
-	-	110	2,36	0,280
-	-	120	2,25	0,250
0,420	2,4	127	-	-
0,356	2,2	150	2	0,182
0,298	2	180	1,8	0,142
0,268	1,9	205	-	-
-	-	840	1,18 (+)	0,090

MARAETCHESFONTEIN 54 HO

G34307 Pumped borehole

Phase 1 Start 1 000, 27/10/1981 Shutdown 1 800, 27/10/1981

R.W.L. 7,80 m Pump intake 62 m Pumping period 480 min.

t (min)	s (m)	Drawdown		Recovery
		t' (min)	s' (m)	s-s' (from F.D.P.W.L.) (m)
0,1	2,14	1	9,33	10,66
0,12	2,45	2	9,03	10,96
0,15	3,14	3	8,80	11,19
0,2	3,60	4	8,57	11,42
0,3	4,74	5	8,42	11,57
0,4	5,51	6	8,26	11,73
0,5	5,81	7	8,19	11,80
0,6	6,24	8	8,11	11,88
0,7	6,58	9	8,03	11,96
0,8	6,73	10	7,93	12,06
0,9	6,98	12	7,76	12,23
1,0	7,16	15	7,54	12,45
1,2	7,04	20	7,34	12,65
1,5	7,04	25	6,96	13,03
1,8	7,42	30	6,81	13,18
2,2	8,11	40	6,43	13,56
2,6	8,42	50	6,04	13,95
3,0	9,18	60	5,78	14,21
3,5	9,82	70	5,51	14,48
4,0	10,25	80	5,32	14,67
4,5	10,40	90	5,05	14,94
5,0	10,63	100	4,90	15,09
6,0	10,86	120	4,59	15,40
7,0	10,99	150	4,13	15,86
8,0	11,02	180	3,83	16,16

G34307		Drawdown	Recovery	
			F.D.P.W.L. = 19,99 m	
t (min)	s (m)	t' (min)	s' (m)	s-s' (from F.D.P.W.L.) (m)
<hr/>				
9,0	11,09	220	3,44	16,55
10	11,25	260	3,06	16,93
12	11,35	300	2,75	17,24
15	11,55	350	2,48	17,51
18	11,66	400	2,30	17,69
22	11,84	480	1,99	18,00
26	11,74	600	1,65	18,34
30	11,93	700	1,38	18,61
35	12,47*	833	1,07	18,92
40	12,70			
45	12,82		Discharge measurements	
50	12,93			
60	13,01	t		Q
70	13,16	(min)		1/s
80	13,39			
<hr/>				
100	13,62	15		6,89
120	13,92	45		7,27
150	14,15	80		7,06
180	14,31	120		6,67
220	14,46	180		6,49
260	18,05**	240		6,37
300	18,97	300		7,50
350	19,05	360		7,58
400	19,35	450		7,42
480	19,99		Av Q =	7 1/s

Water quality 465 mg/l (at 25°C)

* increase in Q

T°C = 21° to 23°

** pump adjusted

MARAETCHESFONTEIN 54 HO

G34307 Pumped borehole

G34307 - Recovery

Phase 2

"R.W.L."

8,870 m

s'
(m)

2

3

4

5

6

7,5

10

13

15

13,000

12,550

12,135

11,460

10,902

10,750

10,047

9,632

8,942

8,335

7,575

6,600

5,788

-

-

1,808

Phase 3

"R.W.L."

10,678 m

s'
(m)

9,265

8,683

8,149

7,952

7,570

7,079

6,827

6,283

6,154

5,869

-

5,207

4,822

4,309

3,900

3,664

3,410

2,989

2,626

2,179

1,627

-

0,643

0,097

6,18

6,49

7,11

660

G34309 - Observation borehole

R.W.L. 8,018 m

F.D.P.W.L. 15,128 m

Phase 2 - drawdown and recovery

s

s'
(m)t and t'
(min)s'
(m)s-s'
(m)

	2	9,265				
	3	8,683	0,25	5	6,96	0,15
	4	8,149	0,51	10	6,91	0,20
	5	7,952	0,89	15	6,83	0,28
	6	7,570	1,11	20	6,73	0,38
	7,5	7,079	1,39	25	6,61	0,50
	10	6,827	1,56	30	6,51	0,60
	13	6,283	1,69	35	6,36	0,75
	15	6,154	1,80	40	6,26	0,85
13,000	20	5,869	1,90	45	6,15	0,96
12,550	25	-	2,01	50	6,06	1,05
12,135	30	5,207	2,25	60	5,95	1,16
11,460	40	4,822	2,44	70	5,73	1,38
10,902	50	4,309	2,65	80	5,57	1,54
10,750	60	3,900	2,83	90	5,39	1,72
10,047	70	3,664	3,00	100	5,23	1,88
9,632	80	3,410	3,26	120	4,94	2,17
8,942	100	2,989	3,67	150	4,54	2,57
8,335	120	2,626	4,08	180	4,10	3,01
7,575	150	2,179	4,53	220	3,69	3,42
6,600	200	1,627	4,89	260	3,34	3,77
5,788	250	-	5,20	300	3,05	4,06
-	335	0,643	5,58	350	2,75	4,36
-	460	0,097	5,91	400	2,47	4,64
1,808	855				2,26	4,85
			6,49	500	2,09	5,02
			7,11	600	-	-
					1,49	5,62

Discharge and water quality measurements

Phase 2

Phase 3

t (min.)	Q (1/s)	T.D.S.		Q (1/s)
		mg/l at 25°C	t	
20	11,8		10	15,82
90	11,8		20	17,17
150	11,6		27	15,82
220	11,43		47	12,5
310	11,10		55	14,08
435	10,90	518	73	13,71
460	11,60*		105	12,1
530	11,33	474	110	shutdown
540	11,42*			Av. Q = 14 l/s
580	12,30*			
600	shutdown			

Average Q = 11,5 l/s

* pump adjusted

MARAETCHESFONTEIN 54 HO

G34307 Pumped borehole

Phase 1 Start 1 000 27/10/1981 Shutdown 1 800, 27/10/1981

Phase 2 Start 0 830 28/10/1981 Shutdown 1 830, 28/10/1981

Phase 3 Start 0 845 29/10/1981 Shutdown 1 035, 29/10/1981

Pumping periods Phase 1: 480 min. Phase 2: 600 min Phase 3: 110 min.

G34308 - Observation b.h.

R.W.L. = 7,418 m r = 230 m

G34309 - Observation b.h.

R.W.L. = 7,048 m F.D.P.W.L. 14,158 m
r = 243 m

Continuous time

from start of

Phase 1: Drawdown and recovery

Phase 1 (min)	Drawdown (s) (m)	t (min)	s (m)	t' (min)	s' (m)	s-s' (m)
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5 - 180	rise in W.L. by 0,005 m	5 10	0,20 0,40		4,21 4,02	0,11 0,30
220	0,002	15	0,60		3,92	0,40
260	0,005	20	0,72	as	3,80	0,52
300	0,010	25	0,83	for	3,70	0,62
350	0,020	30	0,93	column	3,61	0,71
400	0,040	35	1,03	t	3,54	0,78
480	0,070	40	1,12		3,48	0,84
490	0,075	45	1,19		3,41	0,91
520	0,085	50	1,25		3,34	0,98
540	0,092	60	1,37		3,24	1,08
570	0,100	70	1,49		3,15	1,17
610	0,112	80	1,61		3,04	1,28
650	0,125	90	1,71		2,97	1,35
690	0,130	100	1,80		2,91	1,41
740	0,135	120	1,97		2,76	1,56
790	0,145	150	2,18		2,54	1,78
880	0,145	180	2,38		2,39	1,93
980	0,145	220	2,63		2,22	2,10
1 080	0,145	260	2,98		2,06	2,26

G34308 - Observation b.h.

R.W.L. = 7,418 m r = 230 m

G34309 - Observation b.h.

R.W.L. = 7,048 m F.D.P.W.L. 14,158 m
r = 243 m

Continuous time

from start of

Phase 1: Drawdown and recovery

Phase 1 (min)	Drawdown (s) (m)	t (min)	s (m)	t' (min)	s' (m)	s-s' (m)
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1 180	0,150	300	3,29		1,91	2,41
1 340	0,178	350	3,60		1,76	2,56
1 500	0,185	400	3,90		1,65	2,67
1 650	0,218	480	4,32		1,47	2,85
1 800	0,238			600	1,26	3,06
1 900	0,253			700	1,14	3,18
2 000	0,313			830	0,97	3,35
2 200	0,353					
2 400	0,360					
2 600	0,366					
2 875	0,412					
3 000	0,427					
3 200	0,430					
3 400	0,422					
3 600	0,427					
3 800	0,437					
4 000	0,437					
4 330	0,434					

N.B Phase 2 pumping 1 350 to 1 950 min.

Phase 3 pumping 2 895 to 3 005 min.

MARAETCHESFONTEIN 54 HO

G34307 Pumped borehole

Phase 4 Start 0 706, 01/12/1981 Shutdown 0 906, 03/12/1981

Pump intake 64 m Pumping period 3 000 min.

G34307 - Recovery

R.W.L. 7,795 m

t'	s'	t'	s'
(min)	(m)	(min)	(m)

329	14,360		
1 368	6,120	4 000	1,980
1 700	5,060	5 000	1,480
2 000	4,350	6 000	1,140
2 500	3,460	7 000	0,860
3 000	2,815	7 680	0,630

Discharge and water quality measurements

t	Q	pH	T.D.S.	t	Q	pH	T.D.S.
(min)	(1/s)		mg/l	(min)	1/s		mg/l

35	15,29			1 494	11,39		
64	14,94			1 554	11,34		
124	14,37	7,2	464	1 614	11,43		
184	14,06			1 674	11,43		
244	13,64			1 734	11,34		476
304	13,33			1 854	11,25		
364	13,31			1 914	11,25		
424	12,86			1 974	11,16		
484	12,81			2 034	11,16		488
534	12,86			2 094	11,25		
594	12,63	7,45	474	2 154	11,11		

t (min)	Q (l/s)	pH	T.D.S. mg/l @ 25°C	t (min)	Q l/s	pH	T.D.S. mg/l @ 25°C
654	12,41			2 814	10,91		
714	12,46			2 874	10,86	7,2	441
789	*			2 904	10,91		455
804	12,00			2 934	10,91		
1 374	11,43			2 994	10,91		
1 434	11,46					Av Q = 11,7 l/s	

* Engine revs reduced

G34318 r = 1 250 m
R.W.L. = 5,735 m

No change during pumping of G34307

MARAETCHESFONTEIN 54 HO

G34307 Pumped borehole

Phase 4 Start 0 706, 01/12/1981 Shutdown 0 906, 03/12/1981

Pump intake 64 m Pumping period 3 000 mins.

G34308

G34309

R.W.L. = 7,590 m

R.W.L. = 6,092 m F.D.P.W.L. = 21,225 m

r = 230 m

r = 243 m

t (min)	s (m)	t (min)	s (m)	t' (min)	s' (m)	s-s' (m)
90	0,005	3	0,180	20	15,100	0,035
100	0,010	4	0,268	25	15,095	0,040
120	0,015	5	0,373	30	15,085	0,050
150	0,025	6	0,500	35	15,070	0,065
180	0,040	7	0,628	40	15,055	0,080
220	0,060	8	0,753	45	15,030	0,105
260	0,080	9	0,863	50	14,995	0,140
300	0,090	10	0,968	60	14,940	0,195
350	0,105	12	1,152	70	14,830	0,305
400	0,120	15	1,455	80	14,685	0,450
450	0,135	20	1,753	90	14,590	0,545
500	0,145	25	1,953	100	14,380	0,755
600	0,165	30	2,203	120	13,840	1,295
700	0,210	40	2,516	150	12,500	2,635
800	0,255	50	2,831	180	10,900	4,235
900	0,285	60	3,128	220	9,920	5,215
1 000	0,315	70	3,358	260	9,450	5,730
1 200	0,395	80	3,608	300	9,130	6,005
1 500	0,485	90	3,828	350	8,840	6,295
1 800	0,545	100	4,058	400	8,600	6,535
2 200	0,625	120	4,443	450	8,300	6,835
2 600	0,705	150	4,973	500	8,070	7,065
3 000	0,800*	180	5,478	600	7,570	7,565

G34308

R.W.L. = 7,590 m
 r = 230 m

G34309

R.W.L. = 6,092 m F.D.P.W.L. = 21,225 m
 r = 243 m

t (min)	s (m)	t (min)	s (m)	t' (min)	s' (m)	s-s' (m)
3 500	0,880	210	5,943	700	7,030	8,085
4 000	0,930	240	6,328	800	6,560	8,575
4 300	0,945	300	7,033	900	6,145	8,990
4 500	0,950	400	7,903	1 000	5,745	9,390
5 000	0,935	500	8,451	1 200	5,135	10,000
6 000	0 910	600	9,030	1 500	4,345	10,790
7 000	0,860	740	9,960	1 800	3,785	11,350
8 000	0,770	840	10,721	2 200	3,225	11,910
9 000	0,700	900	13,193	2 600	2,805	12,330
10 000	0,640	1 000	14,543	3 000	2,480	12,655
12 000	0,510	1 200	15,943	3 500	2,125	13,010
		1 350	16,310	4 000	1,880	13,255
*pump shutdown		1 500	16,621	4 500	1 665	13,470
		1 680	16,948	5 000	1,510	13,625
		1 800	17,098	6 000	1,195	13,940
		2 100	16,810	7 000	1,005	14,130
		2 400	15,723	8 000	0,860	14,275
		3 000	15,135	9 000	0,680	14,455